Aerosol Distributions and Direct Radiative Forcing – Estimates from a Global Aerosol Analysis with MODIS Assimilation

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Comprehensive Exam Part II March 5, 2004 Program in Atmospheric and Oceanic Sciences University of Colorado, Boulder

Introduction

Aerosols are particles in the atmosphere that can absorb and reflect solar radiation.

Aerosol Radiative Forcing \longrightarrow Climate Response

"It is unlikely that satellites can directly supply the required aerosol information; rather success will depend upon appropriate combinations of satellite data, models, field measurements and surface monitoring. These considerations apply to investigation of the effect of aerosols on clouds, that is, the 'indirect' aerosol climate forcing, as well as the direct aerosol forcing."

Executive Summary of the NASA Global Aerosol Climatology Project

Aerosol Optical Depth (AOD) Assimilation is a method of blending model simulated aerosol fields with satellite observations.

Constraining a global transport model in this way m ay lead to improved estimates of aerosol radiative forcing.

Chemical Transport Models Dust Aerosol Example

Winds fro m Weather Forecasts/Analyses advect tracers in the transport m odel.

 $\overline{\text{kg m}}$ ⁻² day⁻¹

Optimal Interpolation AOD λ = 630 nm *OR* 550 nm

AOD AssimilationAVHRR Advanced Very High Resolution Radiometer*OR***MODIS** Moderate Resolution Imaging Spectrometer

1o by 1o gridded aerosol product Stowe et al 1997Kaufman et al 1998

Meteorological fields **NCEP/NCAR Reanalysis** resolution $T62 \sim 1.9$ ^o, 28 levels *OR* **NCEP Aviation Analysis** resolution T126, 42 levels *OR* **CAM** (NCAR Community Atmosphere Model) resolution T42, 28 levels

MATCH

Model for Atmospheric Transport and Chemistry Rasch et al 1997

SO2/DMS/Carbon Aerosol Emission Inventoriesmonthly climatologies Benkovitz et al 1996Cooke et al 1999Liousse et al 1996

MATCH Configuration **Dust Aerosol**

Sulfur Cycle/ Sulfate Aerosol

Gas phase/aqueous chemistry Barth et al 2000tracers DMS, $\mathrm{SO}_2,$ $\mathrm{SO}_4,$ $\mathrm{H}_2\mathrm{O}_2$ monthly climatologies for O_3 , OH, HO_2 , NO₃ from MOZART (Model for Ozone and its Precursors in the Troposphere)

Hydrological Cycle Prognostic cloud water Rasch and Kristjansson 1997 Vertical convectionZhang and McFarlane 1995 Precipitation - bulk microphysical Flatau 1989

Mobilization and deposition Zender et al 2003 Mahowald et al 2003 4 size categories 0.005 – 0.5 µ m (radius), 0.5 – 1.25 µm, 1.25 – 2.5 µ ^m, 2.5 – 5.0 µ m

> Diagnosed sea-salt aerosol Blanchard and Woodcock 1980 No nitrate aerosol

Carbon AerosolBlack Carbon (Soot) Organic Carbon hydrophobic \rightarrow hydrophilic Cooke and Wilson 1996

> **Aerosol Optics** Sulfate*, Sea-Salt, Organic Carbon, Soot *Optical Properties of Aerosols and Clouds* Hess et al 1998DustZender et al 2003*Currently based on $\rm (NH_4)_2SO_4$

How does AOD assimilation work?

Assimilation adjusts model aerosol mass so that model AOD more closely matches satellite observed AOD.

> $\bm{\tau}_{\bm{\lambda}} = \Sigma_{\mathbf{s}}$ $\Sigma_{\rm k}$ [$\Delta p_{\rm k}$ / ${\rm g}$ ${\rm k}_{\lambda}$ (RH)] ${\rm q}_{\rm sk}$

Single wavelength assimilation scales aerosol mass mixing ratios independent of vertical level and species $\mathbf{q}_{\rm sk} \to \alpha \; \mathbf{q}_{\rm sk}$ through *Optimal Interpolation*, with a spatial correlation length of \sim 100 km.

An example illustrates the subsequent model propagation of this mass correction …

Aerosol Assimilation Example

Saharan Dust Storm March 2, 2003

MODIS AOD at MATCH 1.9^o resolution

March 2

MATCH AOD

AOD Difference

MATCH with MODIS Assimilation (on March 2 only)

MATCH

1.8

0.9

0

 $\rm 0.3$

0

March 3

March 4

1.8

0.9

 $\pmb{0}$

 $\rm 0.3$

 $\pmb{0}$

March 5

March 6

1.8

MODIS Sampling 2001

0.6

0.3

0

MODIS AOD 2001

Aerosol Optical Depth 2001

MATCH

MATCH with MODIS A s similation

Aerosol Optical Depth MATCH/MODIS Correlation 2001

5-day running mean

with Assimilation

-0.3

 $\pmb{0}$

AOD Assimilation Correction 2001

AOD Difference

MATCH with MODIS Assi milation **MATCH**

December - January - February 2001

March - April - May 2001

-0.2

June - July - August 2001

September - October - November 2001

Dust Mass Budget MATCH

Mass \sim 18.6 Tg Emissions \sim 2.7 Tg day⁻¹ Dry Deposition ~ 1.2 Tg day⁻¹ Wet Deposition ~ 1.5 Tg day⁻¹

 τ \sim 7.0 days

MATCH with MODIS Assimilation

 τ \sim 7.0 days

 $\overline{0.5}$ and $\overline{0.5}$

 $g m^{-2}$

0

Sulfate Mass Budget MATCH

Mass \sim 0.6 Tg(S) Emissions \sim 0.005 Tg(S) day⁻¹ Gas Phase \sim 0.02 Tg(S) Aqueous Phase ~ 0.125 Tg(S) Dry Deposition ~ 0.02 Tg(S) day⁻¹ Wet Deposition ~ 0.13 Tg(S) day⁻¹

τ ~ 3.9 days

MATCH with MODIS Assimilation

$Mass \sim 0.73 \text{ Tg(S)}$ Emissions \sim 0.005 Tg(S) day⁻¹ Gas Phase \sim 0.02 Tg(S) $\overline{\text{A}}$ queous Phase ~ 0.125 Tg(S) Dry Deposition ~ 0.025 Tg(S) day⁻¹ Wet Deposition ~ 0.145 Tg(S) day⁻¹

Sulfate Mass

0

 $\mathtt{\tau}\sim 4.3~\mathrm{days}$

 0.008 $g(S) m^{-2}$

Organic Carbon Mass Budget

MATCH

 $Mass \sim 1.7$ Tg Emissions \sim 0.24 Tg day⁻¹ Dry Deposition ~ 0.06 Tg day⁻¹ Wet Deposition ~ 0.18 Tg day⁻¹

 τ \sim 7.2 days

MATCH with MODIS Assimilation

0

 $Mass \sim 2.2$ Tg Emissions \sim 0.24 Tg day⁻¹ Dry Deposition ~ 0.06 Tg day⁻¹ Wet Deposition ~ 0.22 Tg day⁻¹

 τ \sim 7.6 days

Organic Carbon Mass

 g m⁻²

 0.03

Black Carbon Mass Budget

MATCH

 $Mass \sim 0.19$ Tg Emissions \sim 0.03 Tg day⁻¹ Dry Deposition ~ 0.01 Tg day⁻¹ Wet Deposition ~ 0.02 Tg day⁻¹

τ ~ 6.6 days

MATCH with MODIS Assimilation

0

 $Mass \sim 0.25$ Tg Emissions \sim 0.03 Tg day⁻¹ Dry Deposition \sim 0.01 Tg day⁻¹ Wet Deposition ~ 0.025 Tg day⁻¹

 $\mathtt{\tau} \sim 7.1 \; \mathrm{days}$

 g m⁻²

 0.003

Aerosol Radiative Forcing

Top-of-Atmosphere Net Shortwave Flux (Clear-Sky)

TOA Net SW Flux Bias over Oceans

Diffuse Land Surface Albedo Differences CAM

 $0.1\,$

 $\pmb{0}$

-0.1

with MODIS Surface Albedo

CAM

${\rm SW}$

TOA Net SW Flux Bias over Land

CAM -CERES

CAM with MODIS Surface Albedo

Aerosol TOA SW Radiative Forcing (Clear-Sky) fro m CAM with MATCH/MODIS Aerosol Climatology

Dust

Carbon

Sea-Salt

 $\rm 0$

Aerosol Atmospheric Absorption

Change in Forcing CAM Surface Albedo \Rightarrow MODIS/CAM Surface Albedo

Dust

Sulfate

Carbon

Dust TOA Forcing > 0 over Sahara?

Change in Dust Absorption

 $W m⁻²$

2.5

0.5

 -0.5

Forcing

with MODIS

 $W m⁻²$

Global Mean TOA Aerosol Forcing

Conclusions/Future Work

- \Box The MATCH/MODIS Aerosol Dataset provides estimates of aerosol mass and radiative forcing directly constrained by satellite observations.
- \Box Validation of this Dataset will occur in the next few weeks through the AEROCOM (Aerosol Inter-Comparison) project.
- □ The CERES (ERBE-Like Algorithm) for satellite measured fluxes does not detect regional aerosol signatures over oceans.

 However, this Dataset will be an integral part of the NASA CERES-SARB (Clouds and the Earth's Radiant Energy System – Surface Atmosphere Radiation Budget) project, which will estimate top and in-atmosphere fluxes constrained by the CERES measurements using improved CERES/MODIS cloud and scene algorithms.

 MISR (Multi-Angle Imaging Spectro-Radiometer) flux and radiance assimilation.

Thanks to

W. Collins, P. Rasch & N. Mahowald, A. Conley & D. Bundy

This project is funded through NASA ACMAP (Atmospheric Chemistry Modeling and Analysis Project) Grant W-199412.