Optimizing Black Carbon Emissions Using a Kalman Filter

> Jason B. Cohen and Chien Wang

jasonbc@smart.mit.edu



Singapore-MIT Alliance for Research and Technology

Introduction

- > BC heats the atmosphere while it cools the surface.
- This can alter the large scale dynamics of the atmosphere and thus the energy and hydrological cycles.
- > BC distribution is heterogeneous (one week lifetime).
- The emissions are regarded as one of the most important uncertain factors in modeling BC.
- ➤ The current bottom-up emissions estimation is entitled to a factor ≥2 range, adding in model uncertainty raises a concern in estimating the radiative forcing and climate response.

Optimizing for BC Emissions



Observations: AERONET (Holben et al. 1998)

-AAOD instead of AOD

- -Level 2 data (AOD is data, SSA is from an inversion) only
- -Data corresponding to AOD<0.4 not used
- -Data used is from 2001-2006
- -A total of 85 AERONET sites have been used meeting these restrictions

Interactive Aerosol-Climate Model

- 2-moment (mass and number) aerosol module
- 3 sizes of sulfate; Primary OC & BC; aged BC [core/shell]; aged OC [internally mixed]
- Processing includes: condensation of H₂SO₄, nucleation, coagulation, water/cloud interactions, wet and dry deposition.
- Dust Climatology: CCSM Dust Model [N. Mahowald]
- Inversion Runs used CTM mode driven by NCEP Reanalysis
 2001-2006
- Climate Runs in GCM Mode, using a slab ocean model.
- Effects Include Urban Processing Metamodel

Dust Climatology



Sensitivity Plot for lat=55,lon=24

Sensitivity Matrix for Selected Stations



Sensitivity Plot for lat=80,lon=26



18



Optimization is Robust to Various Assumptions

All Zones



Optimization over Specific Continents

Northern East Asia



- CERTEAR

Optimized Emissions of BC



Derived Climatological AAOD



Column Absorption is Increased both in regions with increased emissions as well as in other regions. Some regions experience a large seasonal variation.

Winter

Summer

-10.0	-8.0	-6.0	-4.0	-2.0	0.0	2.0	4.0	6.0	8.0	
Optimized-Urban AAOD*1000										

Impact on Atmospheric **Radiative Forcing**



180W 135W 90W 45W 45E 90E 135E 180E 0

Spring

Winter

ATMOS = TOA - SURFACE

This is among the key factors to determine BC's climate effect.

Most of the change of the global energy budget, due to radiative forcing, is because of atmospheric absorption by BC.

Forcing is not uniform (eg: North to South Gradient) and may induce a non-linear climate response.

and		a all and	and the second
and and	Carin	Contra Contra	V Contraction
7 7	Jos	10.	Q,
	s de .		



-4.0	-2.0 U	0.0 Jrban-Optimized FSNA	2.0 A	
ABS d[W/m2]	Global	SE Asia	India	
Annual	0.75	1.1	2.0	
Winter	0.59	1.2	2.0	
Spring	0.78	1.1	2.0	
Summer	0.94	0.82	1.7	
Autumn	0.69	1.2	2.1	

Difference in Climate Response: Convective Precipitation



South East Asian Response: Precipitation



Conclusions

- Using the MIT-CAM (Urban) model and AERONET measurements, leads to an optimized global BC emissions of 17Tg/yr.
- Optimized minus A-priori emissions are distributed non-uniformly.
- The optimized emissions are quite robust to uncertainties in the error magnitude of the AERONET data, however to obtain better spatial or temporal disaggregation, new data sources are required.
- Since these changes are large in amount and unequally distributed, they are expected (via the direct effect) to induce changes in large scale dynamics, clouds, and precipitation.
- The regional-scale climate response is more variable and stronger in magnitude than the global scale average response.

Thank You

Any Questions?