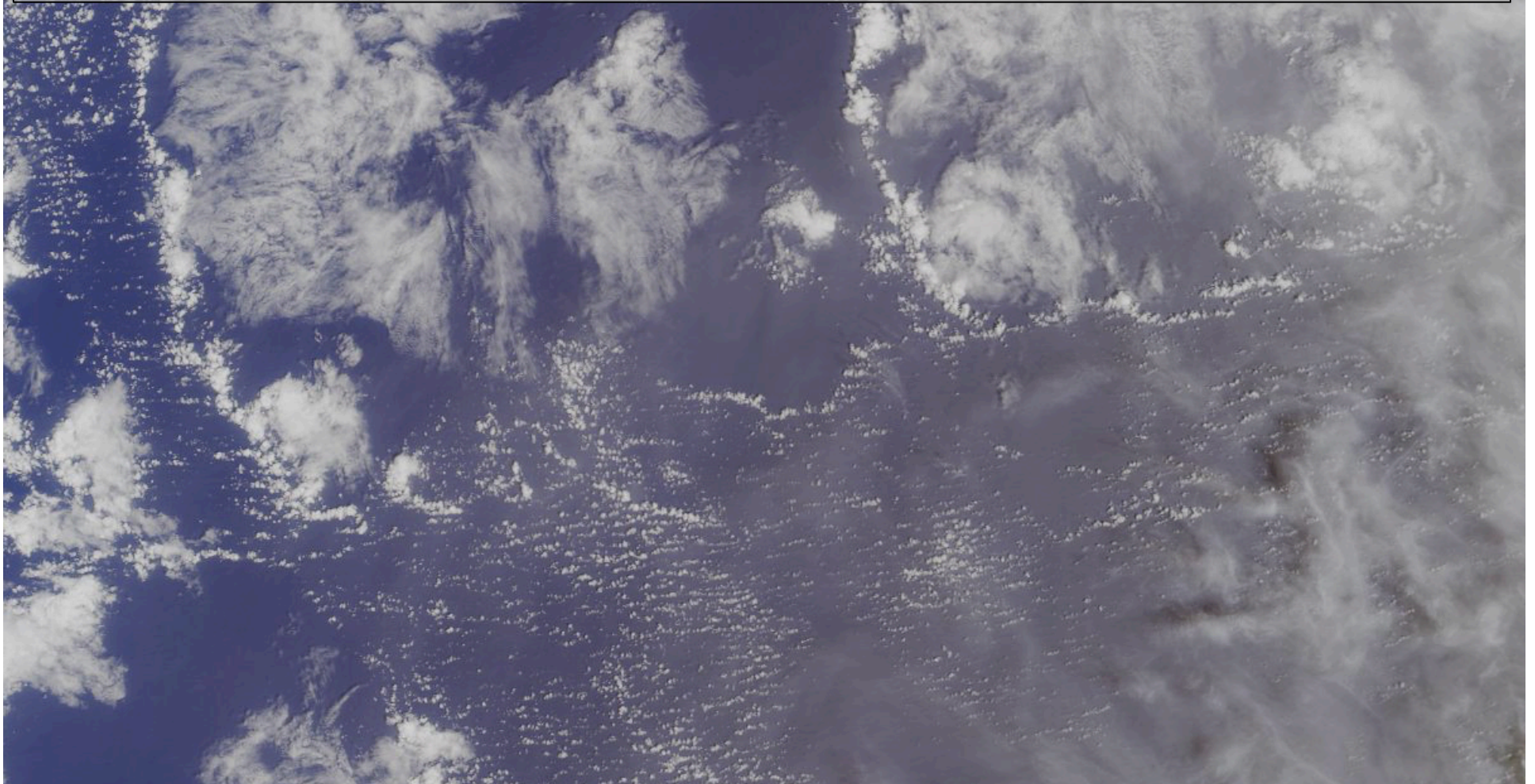


How to Support Aerosol-Cloud Interaction Studies

Insights from ACPC

Ralph Kahn

NASA/Goddard Space Flight Center

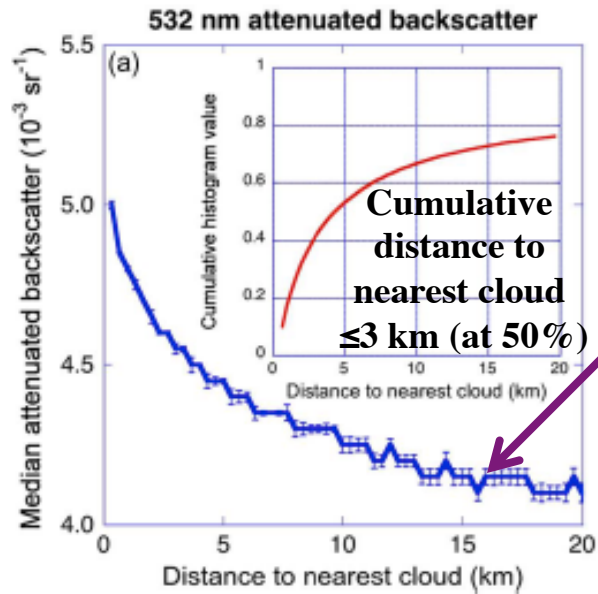


Satellite Aerosol Retrieval *Limitations*

- Difficult to retrieve aerosols that are *collocated with cloud*
 - *Cloud-scattered light* & cloud “contamination” can affect near-cloud aerosol retrievals
 - Aerosols can also affect the *retrieval of cloud properties*
- Rarely can detect aerosol in *droplet-formation region* below clouds – need cloud & aerosol *vertical distributions*
- Aerosols smaller than about *0.1 micron diameter* look like atmospheric gas molecules – must *infer CCN* number
- Must deduce aerosol *hygroscopicity* (composition) from qualitative “type” – size, shape, and SSA constraints
- Environmental (Meteorological) Coupling – Factors can *co-vary*
 - LWP can decrease as aerosol number concentration increases (also depends on atm. stability)
- Many aerosol-cloud interaction time & spatial scales do not match *satellite sampling* (horizontal & vertical res., snapshots, coverage)

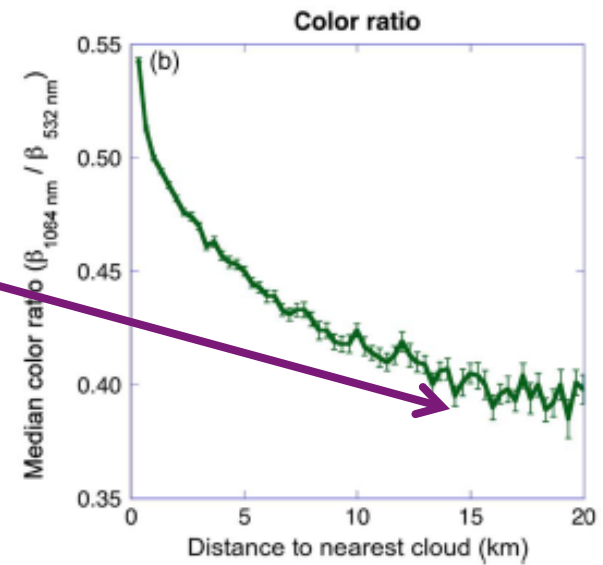
*Satellites are fairly blunt instruments
for studying aerosol-cloud interactions!!*

Aerosol Properties Near Cloud



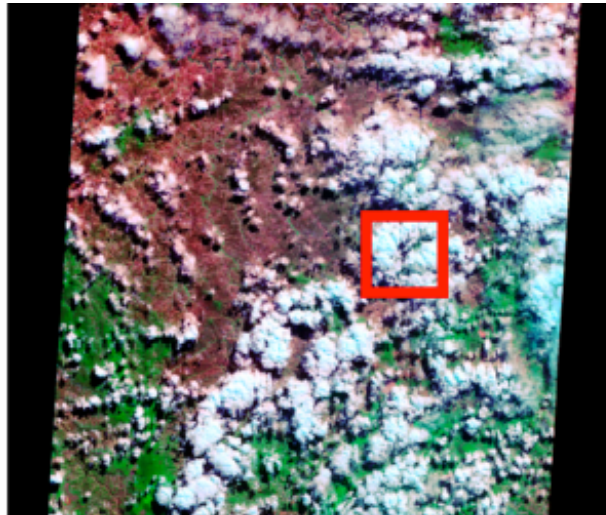
Cumulative distance to nearest cloud ≤ 3 km (at 50%)

Backscatter & color ratio enhanced to ~ 15 km



Global data
Sept. – Oct. 2008

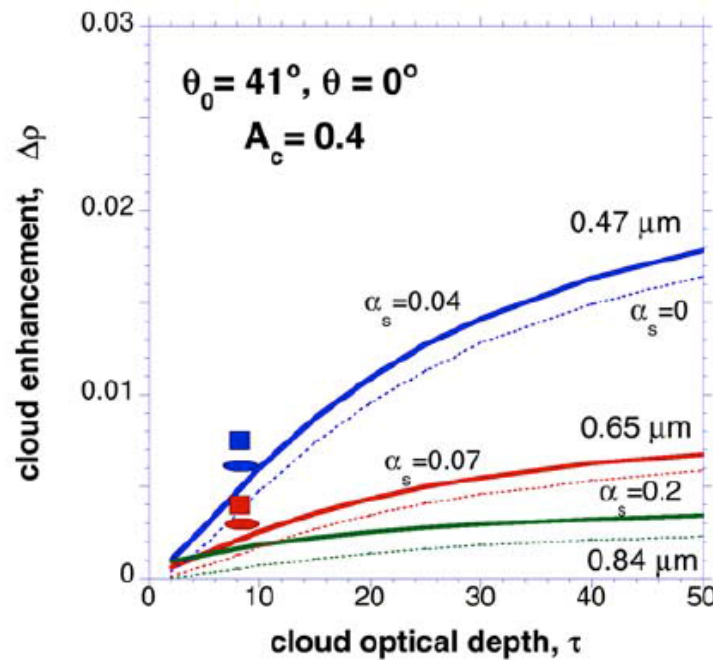
3-D Light Scattering Effects on Remote Sensing



ASTER false-color image
Brazil, 09 August, 2001

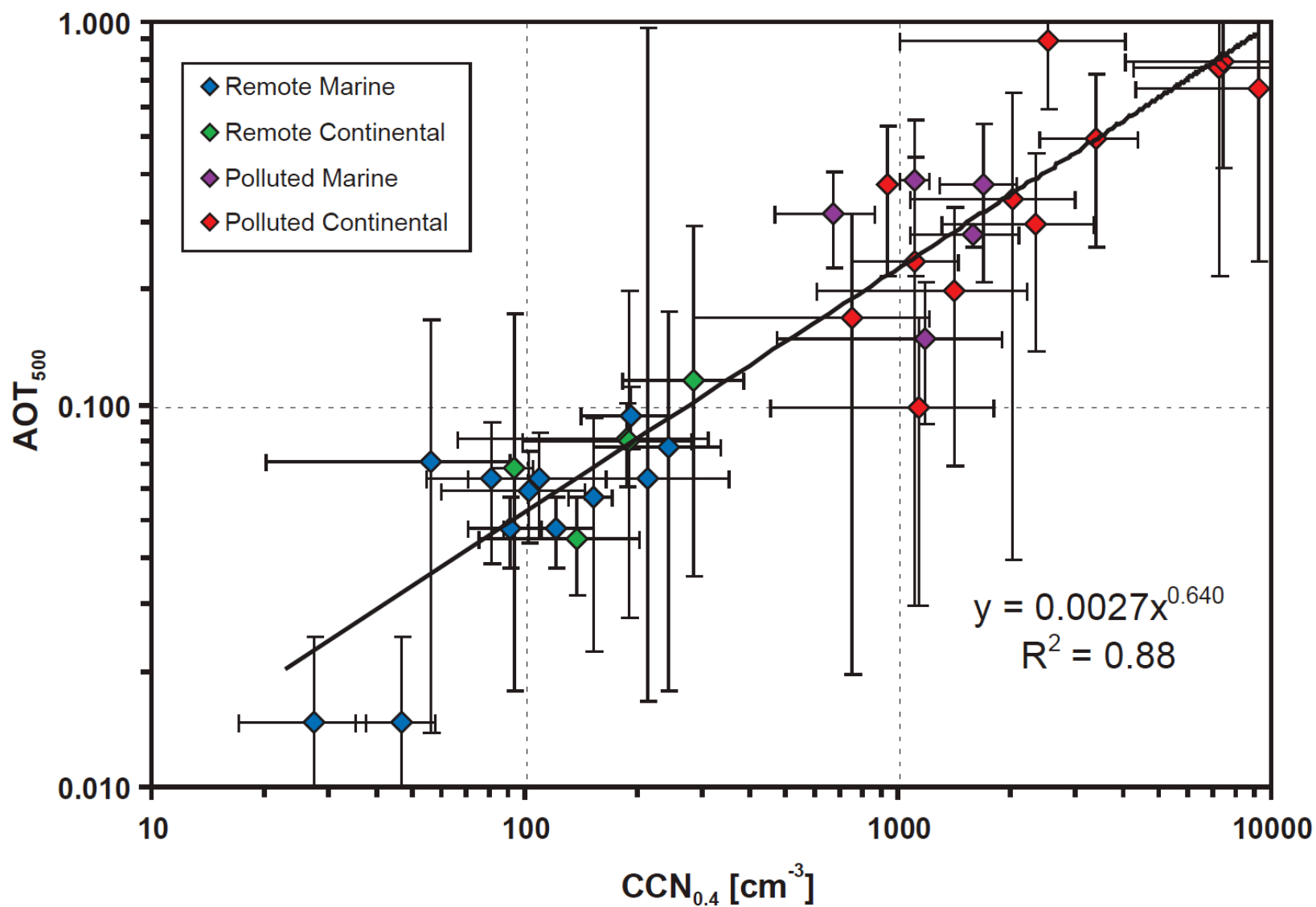
Refl. in “clear” pixels
used for MODIS AOD
Retrievals (*squares*)

Refl. in pixels 3 km
away from cloud (*ovals*)
[Wen et al. 2007]



- Simulated cloud \rightarrow Rayleigh
scattered light enhancement vs. τ_c
- Using the image geometry
 - For three wavelengths
 - For different surf. reflectances (α_s)

Correlation Between AOD from Space and CCN in Remote & Polluted Regions



Using $AI (= \tau_a \times ANG)$ to Estimate CCN

Kapustin, Clarke, et al., JGR 2006

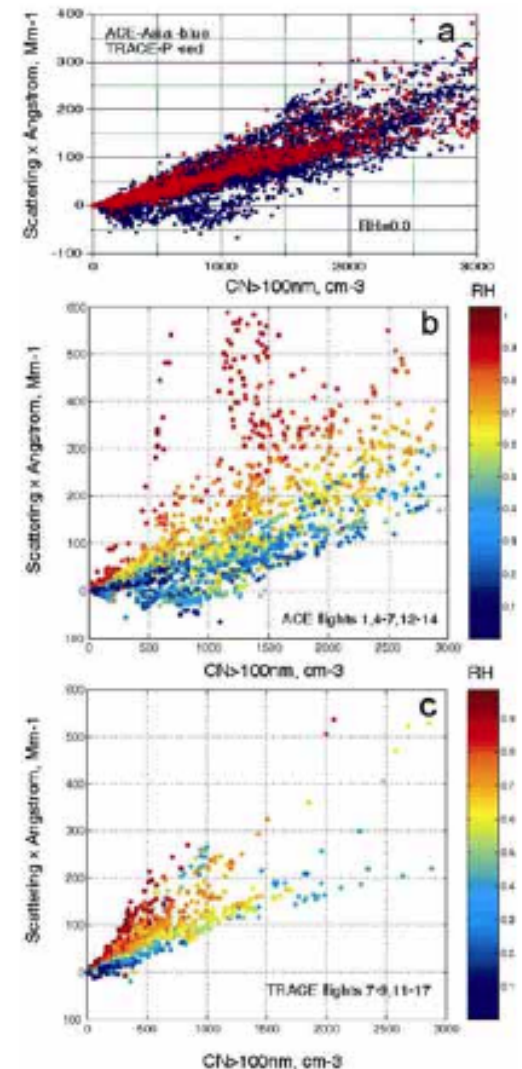
- Test Idea: Smaller particles more likely to become CCN ; Ang is a smaller quantity for larger particles
- ACE-Asia, Trace-P *in situ* **field data** – CCN proxy
- **AI does not work quantitatively in general**, but can **if the data are stratified** by:

- **RH** in the aerosol layer(s) observed by satellites
- **Aerosol Type** (hygroscopicity; pollution, BB, dust)
- **Aerosol Size** (Ang is not unique for bi-modal dist.)

Practically, in addition to τ_a and Ang , this requires:

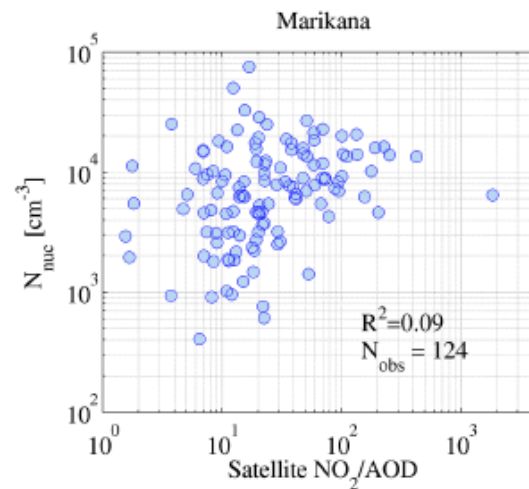
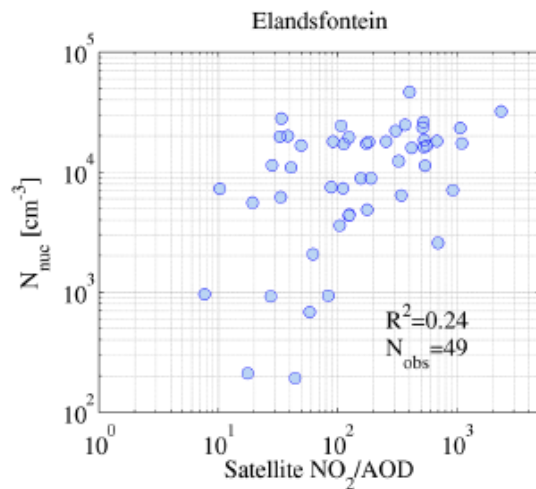
- Vertical **humidity structure**
- **Height-resolved aerosol type**
- **Height-resolved size** dist.
[extrapolated to small sizes(?)]

This study includes enough detail to assess $AI \sim N_a$ and $AI \sim CCN$



AI vs. *in situ* CCN proxy
(a) all ACE (blue) & Trace-P, **dry**
(b) ACE - OPC-only, amb. RH
(c) TP - OPC-only, amb. RH

Satellite-Derived Proxies for CCN



Satellite
 NO_2/AOD proxy
vs.
in situ nucleation
particle concentration

- OMI NO_2 Column
- OMI SO_2 Column (mainly near-surface)
- OMI UVB (310 nm) Surface noontime irradiance to form secondary sulfate
- MODIS AOD [*attempt* to represent the condensation *sink* for nucleation particles]

Sundstrom et al., ACP 2015

***These are quantities we can retrieve from satellites,
though they are not necessarily the ones we really want***

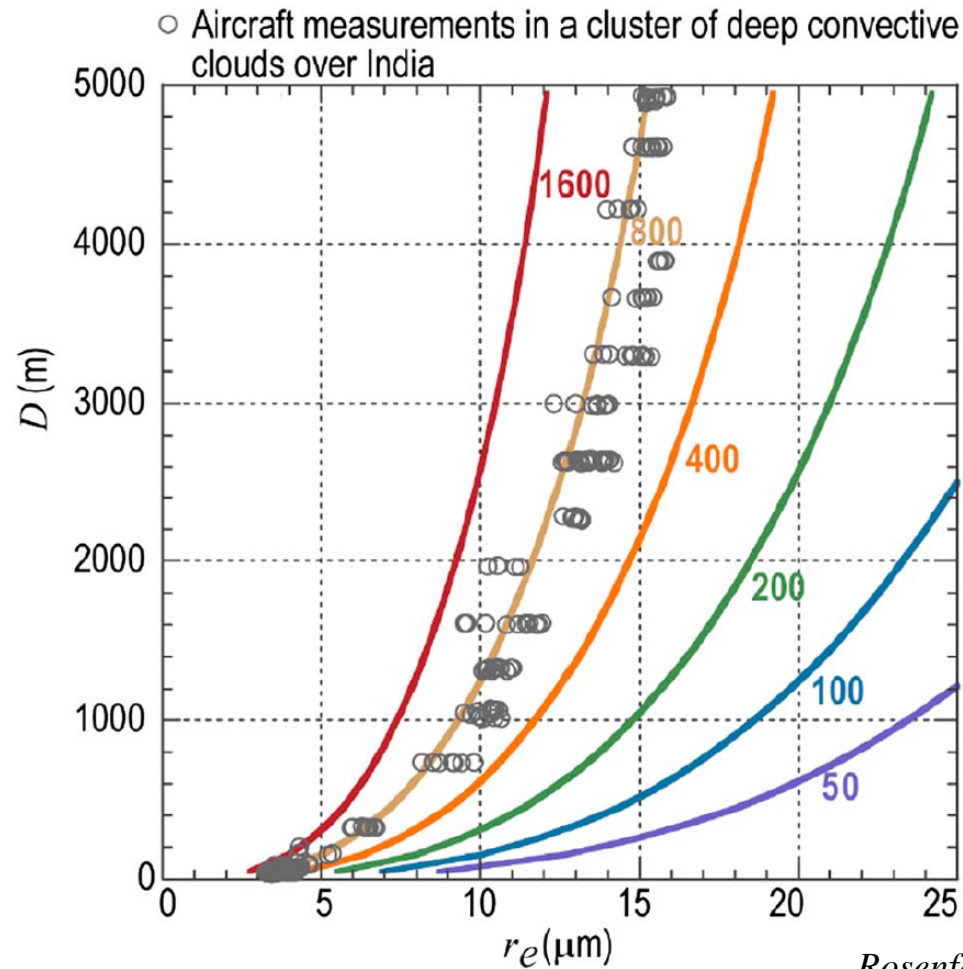
- Ambiguity in vertical distributions of formation areas and sinks
- Lack of information about diurnal variation from satellites
- The 2-D spatial distribution of proxies compares ~ better with *in situ* observations for S. Africa, except where gas column concentrations are low

Satellite Capabilities

- *Polar orbiting imagers* provide **frequent, global coverage**
- *Geostationary platforms* offer **high temporal resolution**
- *Multi-angle imagers* offer **aerosol plume height & cloud-top mapping**
- *Passive instruments* can retrieve total-column **aerosol amount (AOD)**
- *Active instruments* determine aerosol & some cloud **vertical structure**
- *UV imagers* and *active sensors* can retrieve **aerosol above cloud**
- *Multi-angle, spectral, polarized* imagers obtain **some aerosol type info.**
- *Active sensors* can obtain **some aerosol type info., day & night**
- *Satellite trace-gas retrievals* offer **clues about aerosol type**
- *Vis-IR imagers* can retrieve **cloud phase, r_c , T_c , p_c , τ_c , α_c , C_f , LWP**

***Need to be creative &
Play to the strengths of what satellites offer!!***

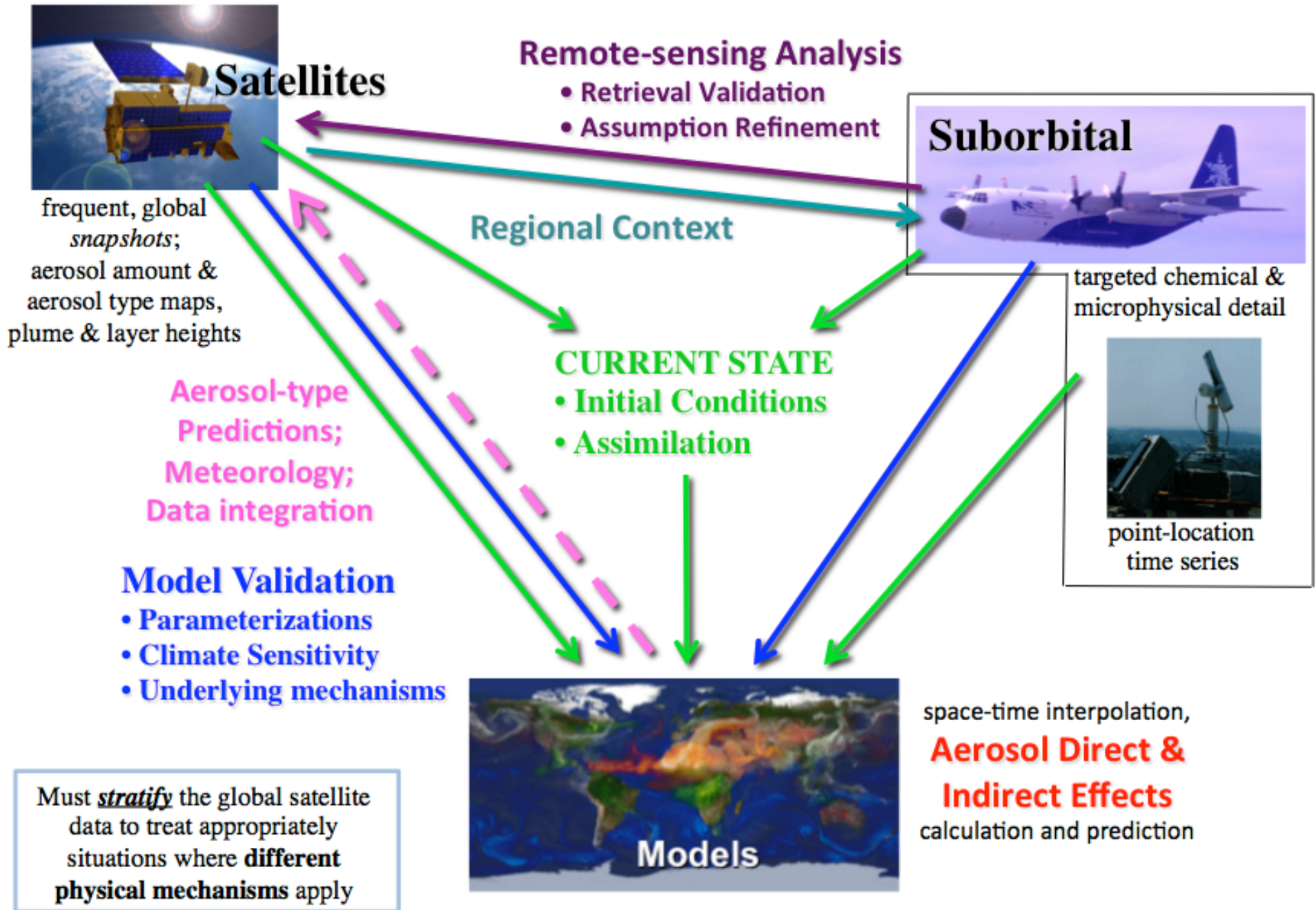
“Cloud-Chamber” Satellite CCN Retrieval



Knowing $r_e(D)$, n_c
at cloud base can
be deduced for an
adiabatic cloud
with limited
entrainment

Rosenfeld et al., Rev. Geophys. 2014

Figure 9. The adiabatic effective radius as a function of height above cloud base (D) for clouds with various number of cloud drop concentrations at their base, for cloud base at a height of 2 km and temperature of 15°C. Actual aircraft measurements of $1 \text{ s } r_e$ are shown for a cloud that had $610 \text{ drops cm}^{-3}$ at its base. From



Adapted from: Kahn, Survy. *Geophys.* 2012

Backup Slides

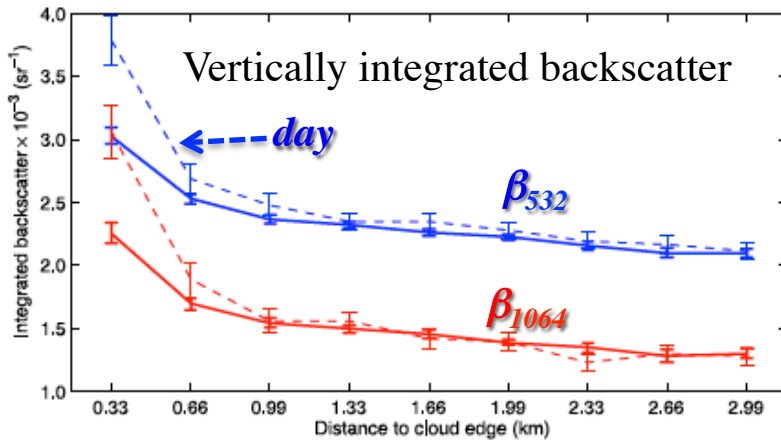
For Aerosol-Cloud Interactions – Overall Satellite *Limitations*

- *Polar orbiters* provide *snapshots only*
- Difficult to probe *cloud base*
- Typically ~100s of meters or poorer *horizontal resolution*
- *Passive instruments (imagers)* offer little *vertical information*
- *Active instruments (e.g., lidar)* offer little *spatial coverage*
- Little information about aerosol *particle microphysical properties*
- Bigger issues retrieving aerosols *in the presence of clouds!*
- Cloud property retrievals can be aliased *by the presence of aerosols*

Aerosol Properties Near Cloud

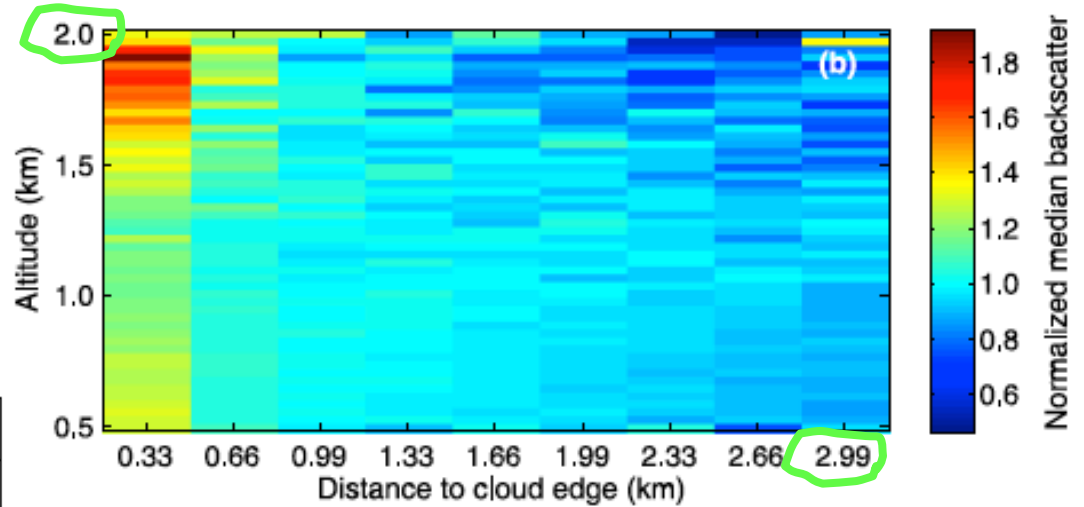
CALIPSO nighttime 532 nm backscatter, normalized over 2.99 km.

Enhanced aerosol opacity near cloud edge, especially at **cloud top and bottom**.



CALIPSO median nighttime 1064/532 nm color ratio.

Larger particles near cloud edge, especially at cloud top and bottom.



Detrainment at cloud top??
Hygroscopic growth at cloud bottom??
Collision Coalescence ($R\uparrow$; $N, \sigma\downarrow$)??

