

# Trans-Atlantic Dust Transport and Deposition - Models vs. Observations

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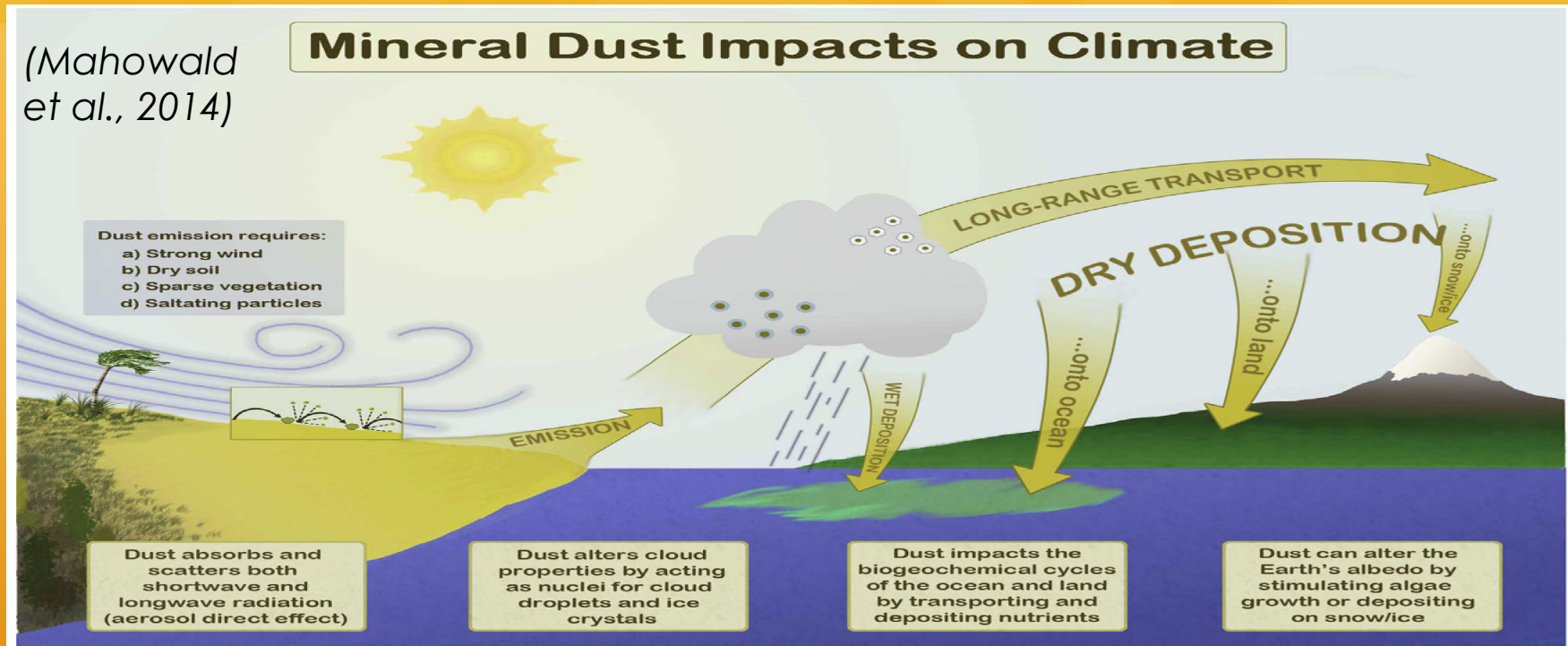
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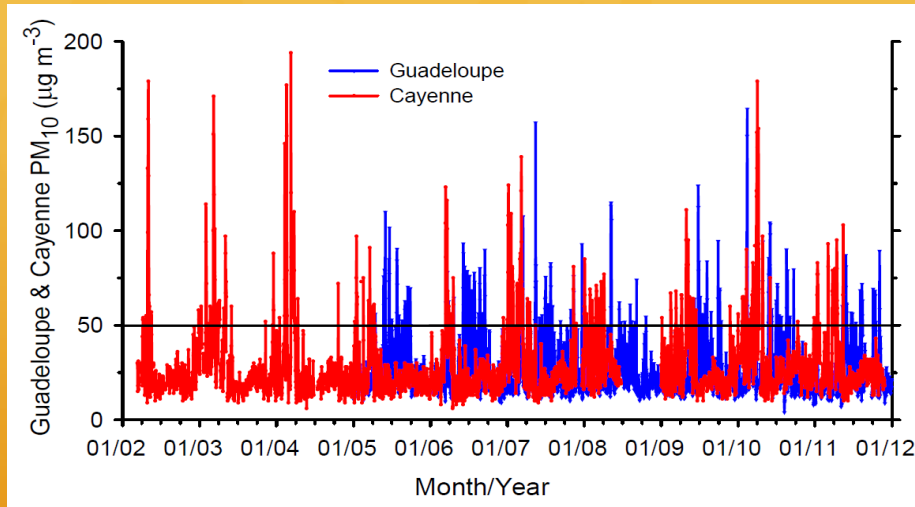
# Dust Interactions with Climate



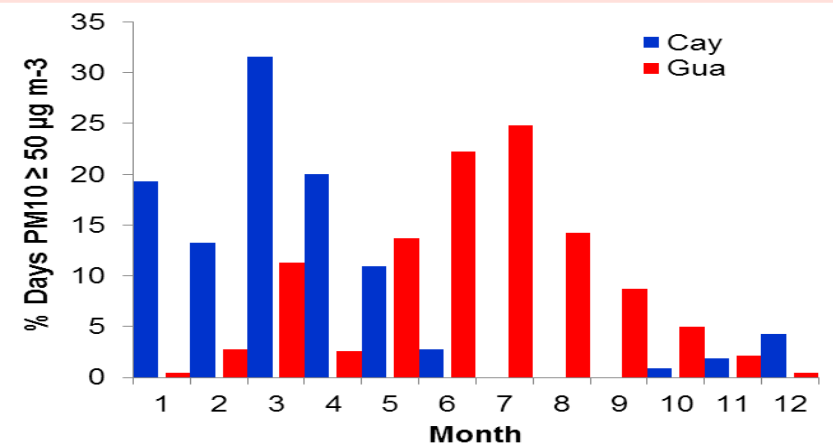
- Dust affects climate (and weather) in many ways (e.g., radiation, clouds, biogeochemical cycle, snow albedo)
- Dust emissions and transport are affected by climate.

# Dust Impacts on Air Quality

(Prospero et al., GBC, 2014)



Percent of Days Exceeding  $50 \mu\text{g m}^{-3}$



- African dust elevates the surface PM<sub>10</sub> level exceeding WHO guideline in the Caribbean Basin.
- The rate of PM<sub>10</sub> non-compliance is comparable to those observed in major urban areas in Europe and the U.S.

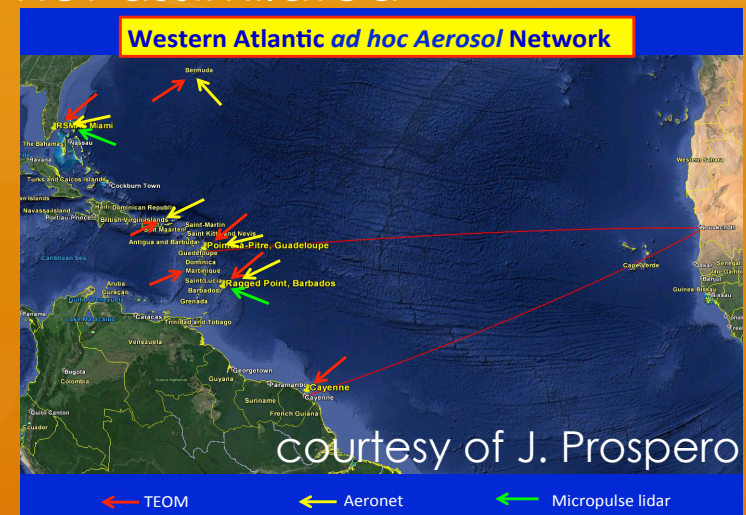
# Objectives

- To quantify altitude-resolved trans-Atlantic transport and deposition of dust from satellite observations
- To compare observations and models

# DATA & MODELS

- CALIOP/CALIPSO observations of 3-D distribution of aerosols
- Model simulations
  - GOCART (1980-2009)
    - offline CTM driven by MERRA met.
    - 2.5x2 resolution
  - MERRA2 aerosol reanalysis (1979-2014)
    - GOCART module
    - 0.625x0.5 resolution
    - total AOD at single- $\lambda$  is constrained by satellite obs.
    - composition and profile are not assimilated

- Ground-based measurements
  - Dust concentration
  - AOD
  - PM10, PM2.5
  - Vertical profile



# CALIOP/CALIPSO – 3D View of Dust Transport

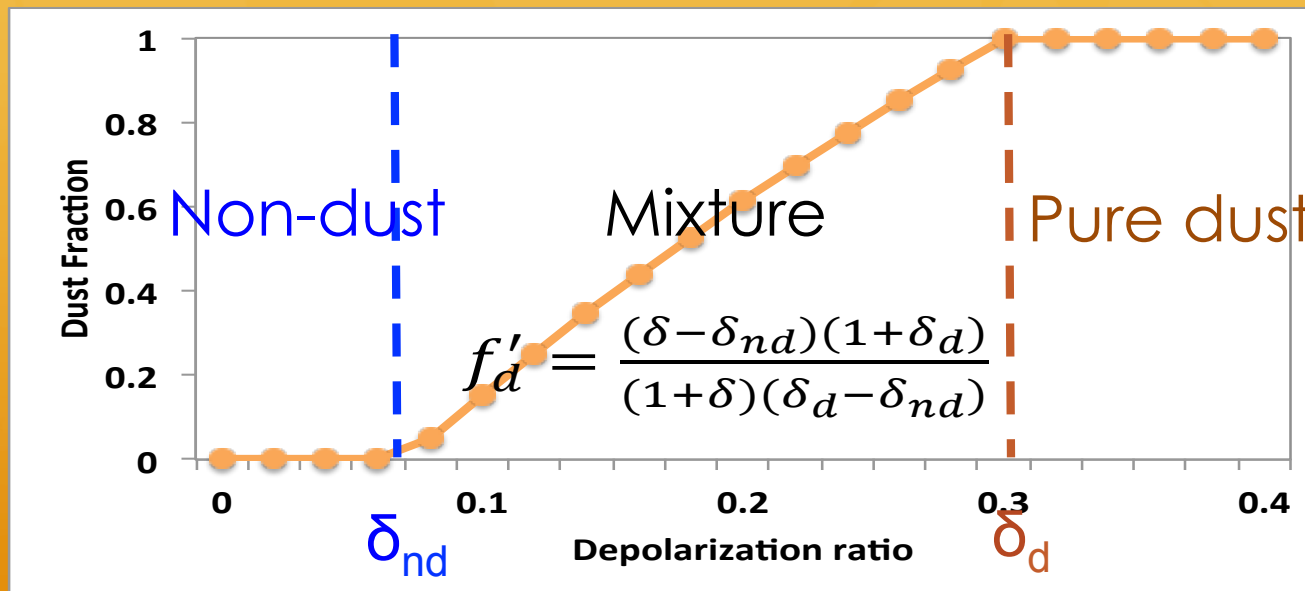
CALIOP is a polarization lidar onboard CALIPSO since June 2006

- 3D distributions of aerosol backscatter and extinction – realistic transport height
- Particulate depolarization ratio – separating dust from non-dust
- Above-cloud aerosol (ACA) profiles – new information



# Separating Dust from Non-dust Aerosol

Dust particles are large in size and non-spherical in shape   
They have significantly larger depolarization ratio than smoke and marine aerosol.



- ❑ Lower-bound Dust Fraction (LDF):  $\delta_d = 0.30, \delta_{nd} = 0.07$
- ❑ Upper-bound Dust Fraction (UDF):  $\delta_d = 0.20, \delta_{nd} = 0.02$

We use the average of LDF and UDF as the best estimate

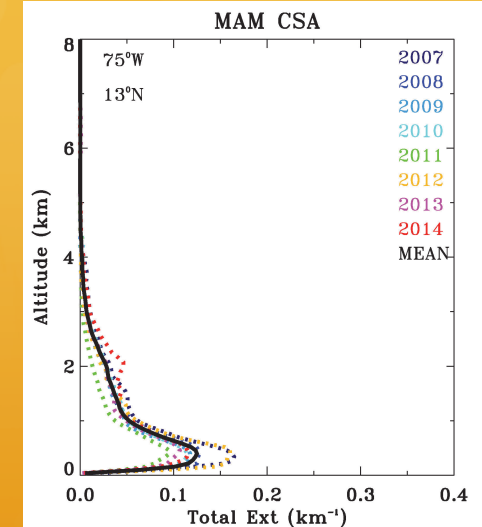
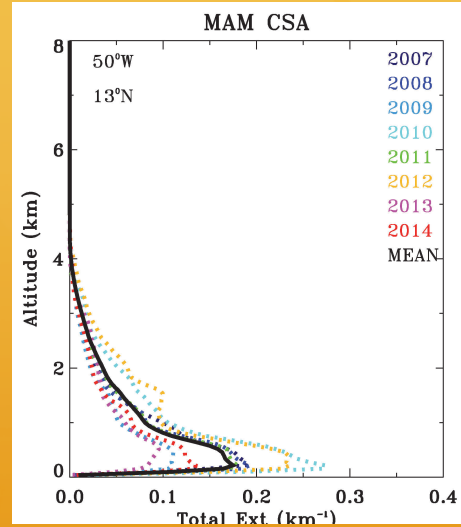
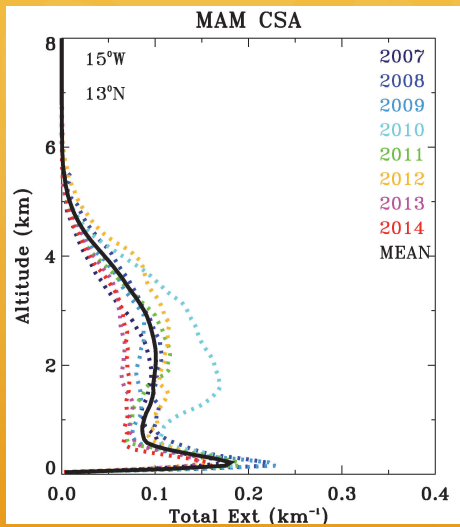
# 2007-2014 Extinction Profiles – Spring

## African Coast

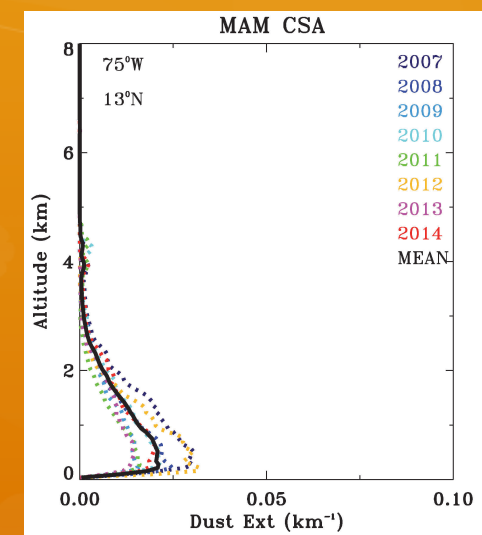
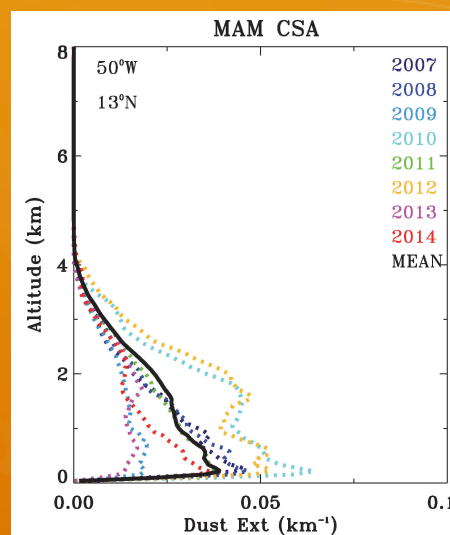
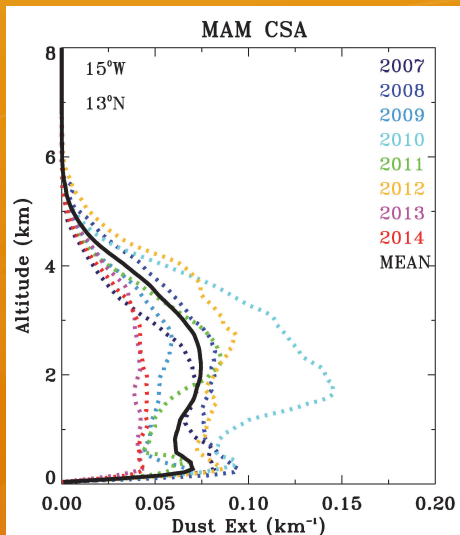
## Upwind of Barbados

## West Caribbean

Total



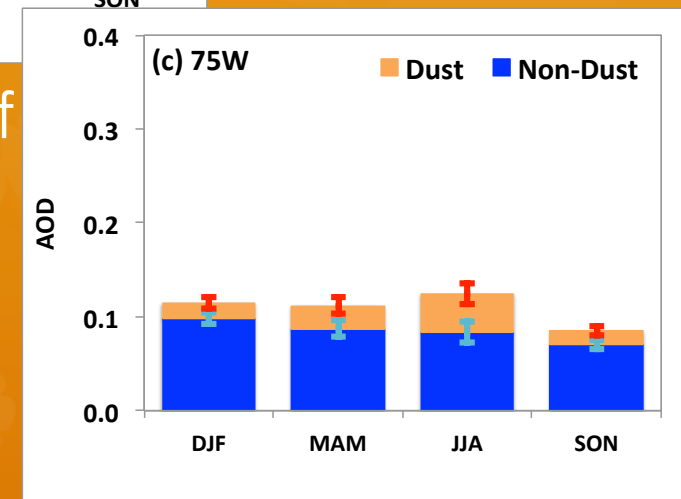
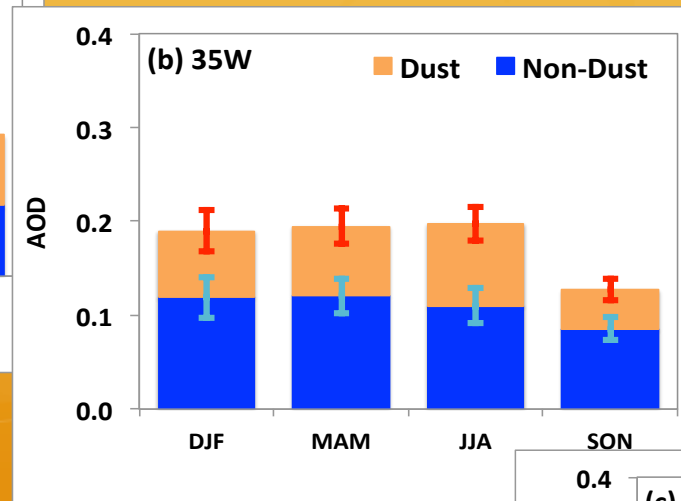
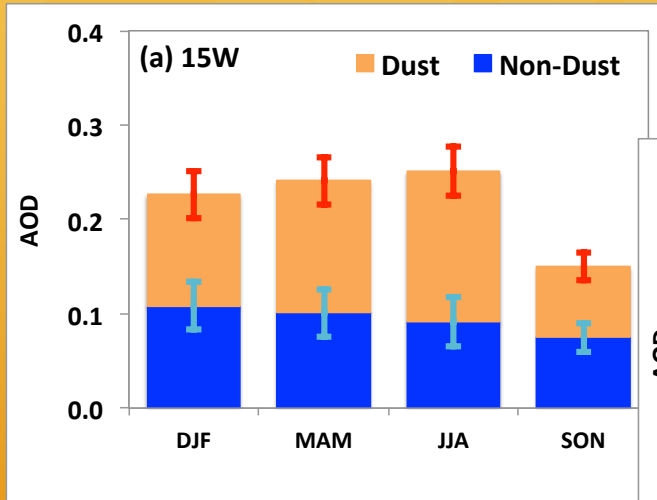
Dust





# Partition of Dust & Non-Dust

2007-2014 & 10S - 30N Average



Sensitivity of DOD to the thresholds of depol. ratio (UDF & LDF scenarios):

- 15W:  $\pm 19\%$
- 35W:  $\pm 26\%$
- 75W:  $\pm 32\%$

# Estimate of Dust Mass Flux from CALIOP

Dust backscatter/  
extinction profile from  
CALIOP

MEE (Mass  
Extinction  
Efficiency)

Extinction (1/m) =  
Mass Conc. (g/m<sup>3</sup>) \* MEE (=0.37 m<sup>2</sup>/g)

Profile of Dust Mass  
Concentration (m)

MERRA  
reanalysis  
wind field

Dust Mass Flux  
 $F = \int m(z)u(z)dz$



$$DD = -(\partial F_x / \partial x + \partial F_y / \partial y)$$

Dust deposition (DD) is estimated as the convergence of zonal & meridional fluxes (assuming no leak from the top)

# Sources of Uncertainty

- Dust and non-dust separation:  **$\pm (15\sim34)\%$**
- Unchanged size:  $\pm 15\%$  (Maring et al., 2003)
- CALIOP AOD low-bias: **-30%**  
(Omar et al., 2013; Liu et al., 2014)
- Low dust MEE/ high mass bias: **+30%**  
(Ansmann et al., 2012)
- Shape of vertical profile: *probably small*,  **$\pm 10\%$**

Overall known uncertainty:  **$\pm (45\sim70)\%$**

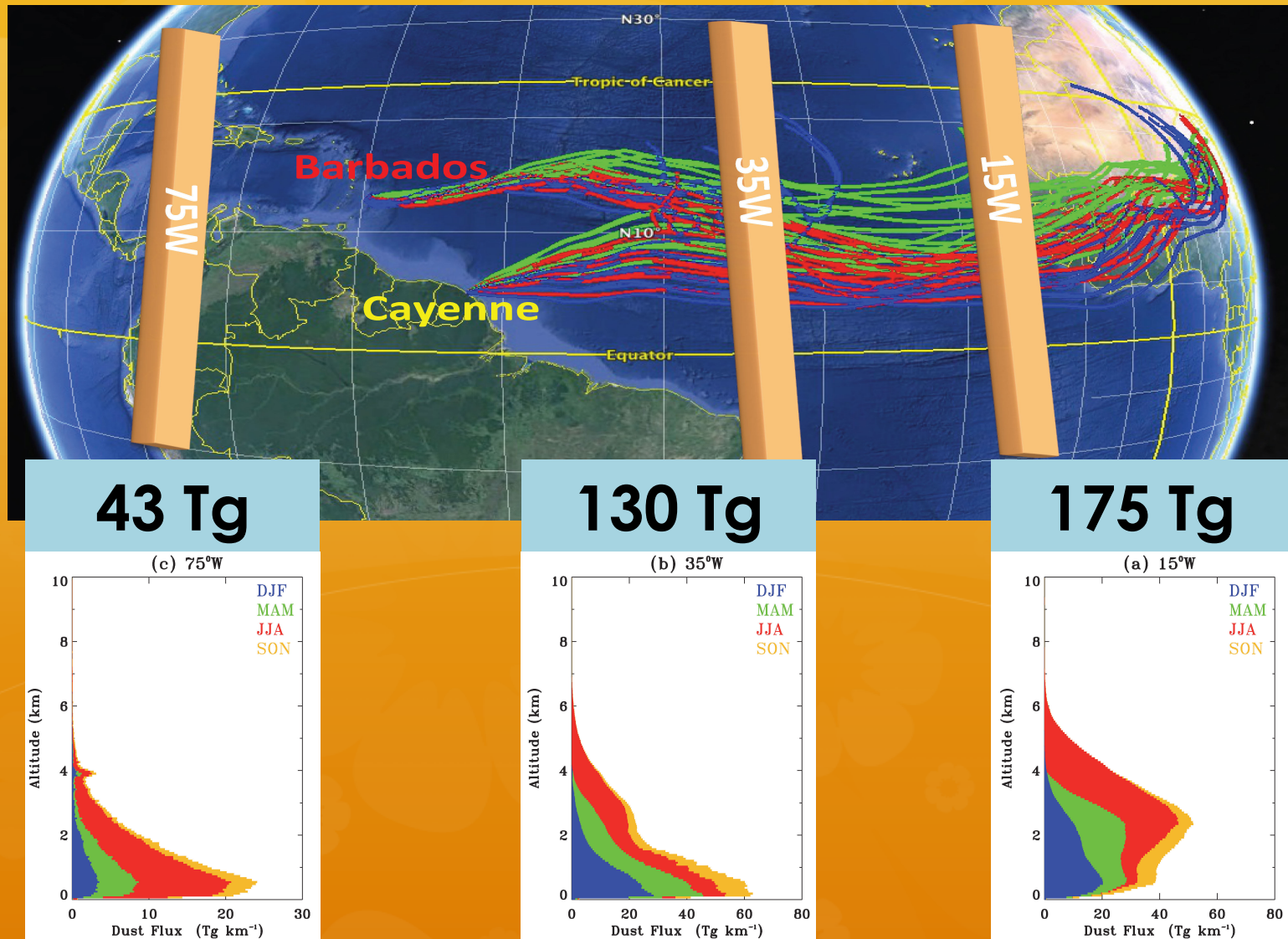
*Other uncertainties:*

Below-cloud dust

Diurnal variation

*Yu et al. (RSE, 2015)*

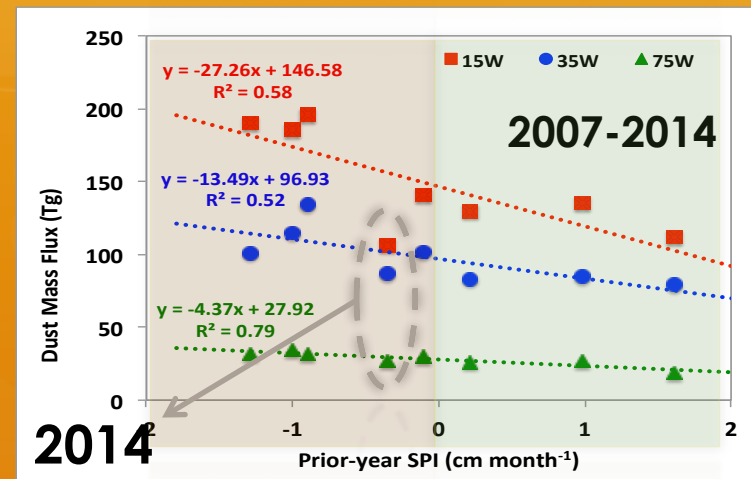
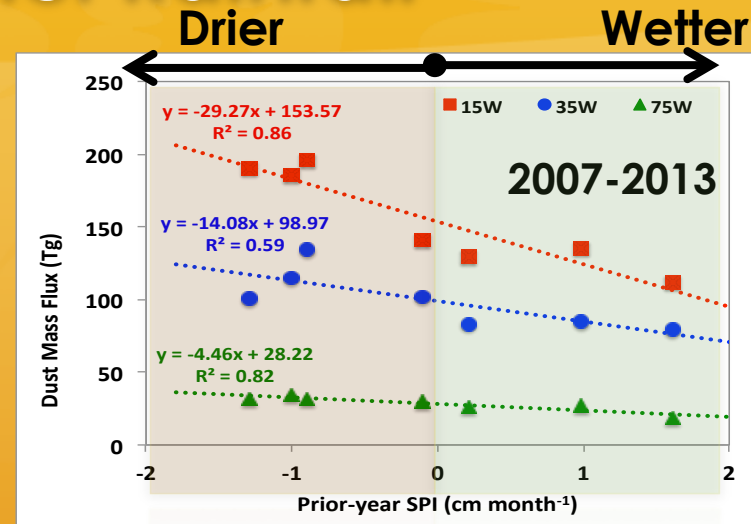
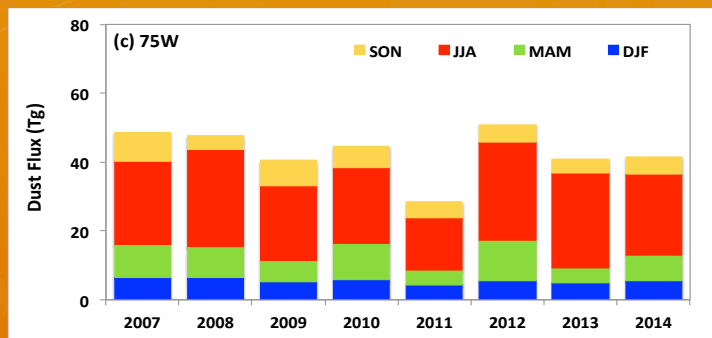
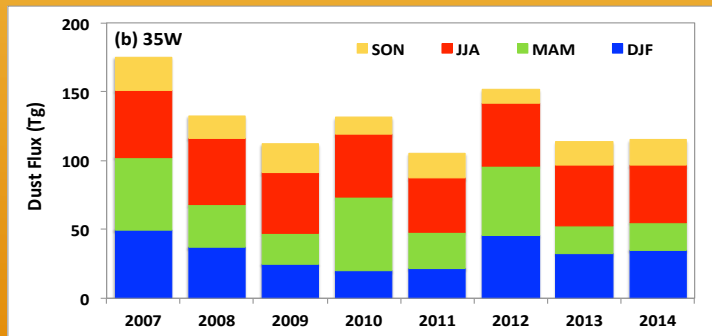
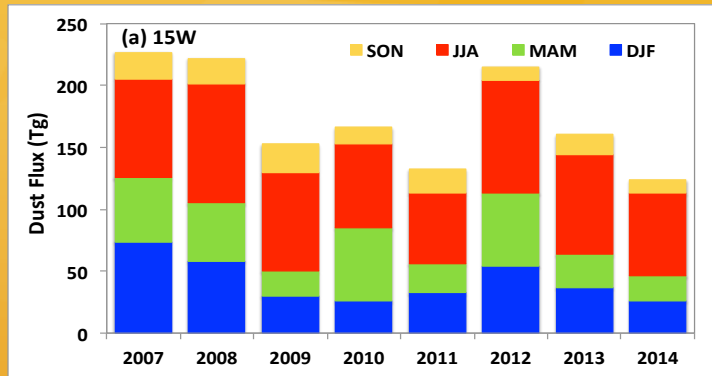
# 8-year Average Dust Flux & Vertical Profile



10°S-30°N integrated flux

1 Tg = 10<sup>12</sup> g

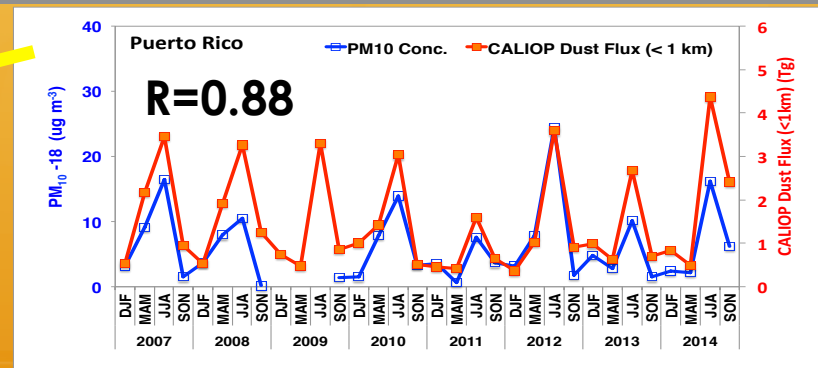
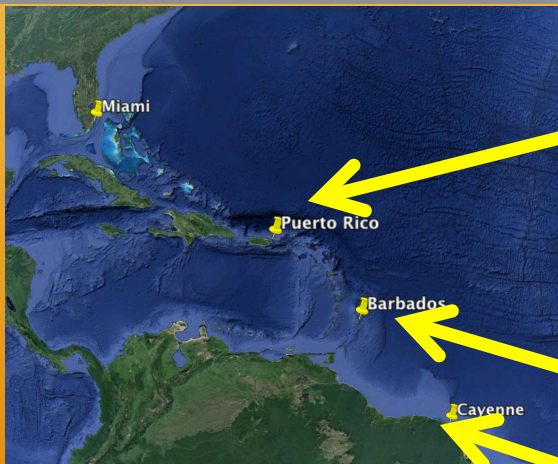
# Interannual variation of dust transport vs. prior-year Sahel Rainfall



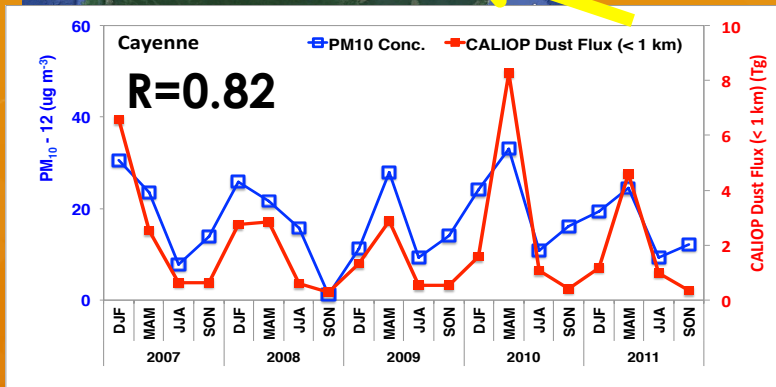
SPI - Sahel Precipitation Index (data from JISAO of NOAA & U. Washington)

# Correlation of Dust Flux with Surface Dust Concentration

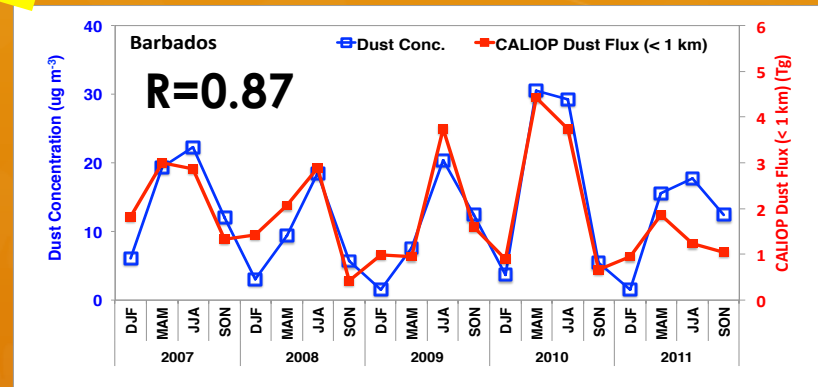
- Dust flux is calculated along the 550 km meridional line just upwind of the interested regions.
- Assume dust flux in the lowest 1 km layer is responsible for dust detected in the surface.



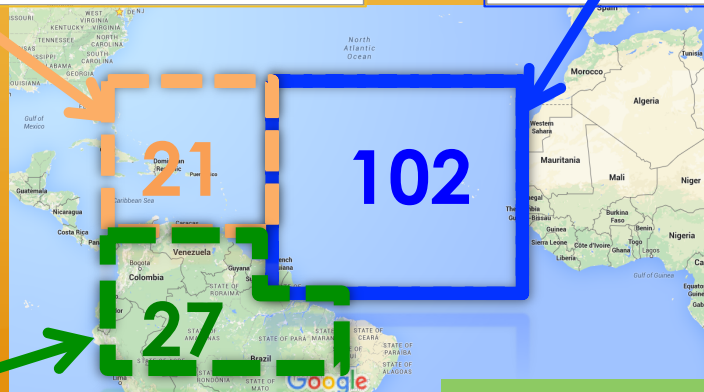
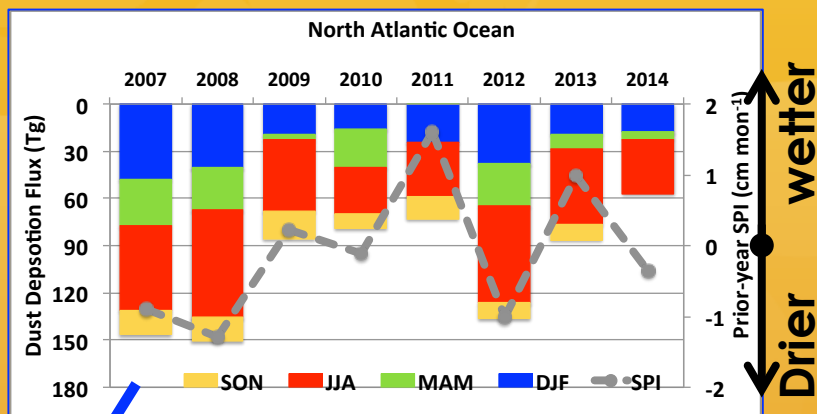
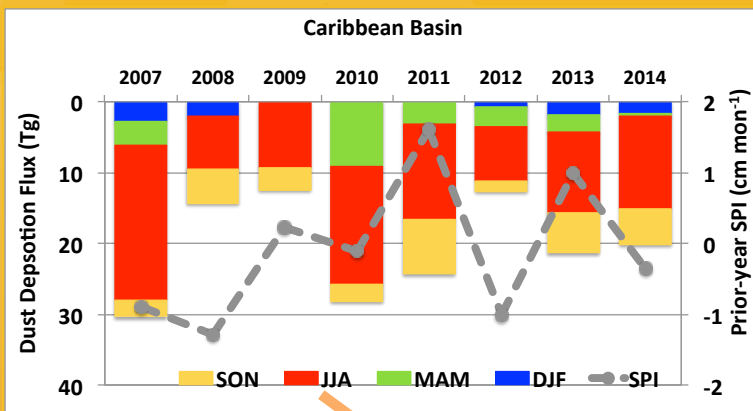
18 ug/m<sup>3</sup> is the lowest seasonal mean



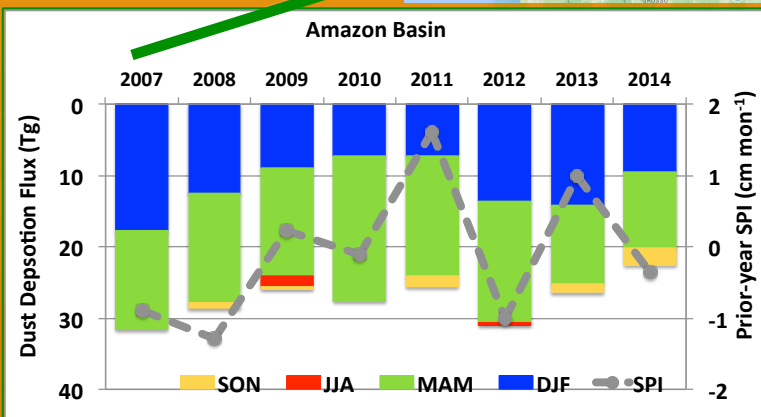
12 ug/m<sup>3</sup> is the lowest seasonal mean



# Dust Deposition into Ocean and Amazon

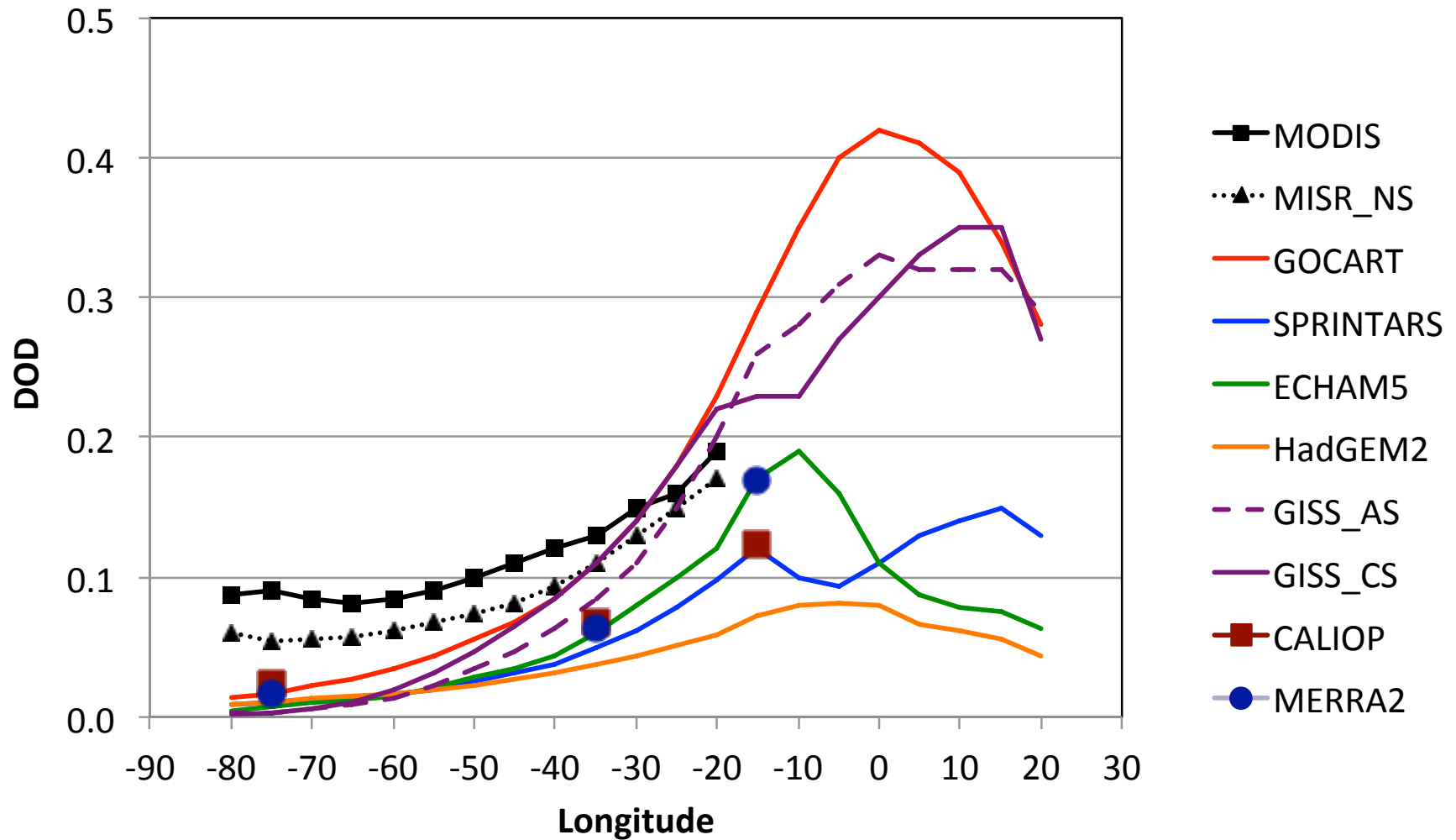


DD vs. SPI:  $R = -0.71$  (07-14)  
 $R = -0.88$  (07-13)



- African dust adds 23 (7-39)  $\text{g ha}^{-1} \text{a}^{-1}$  phosphorus (P) to P-deficient Amazon Basin, which replenishes the loss of P by rains and flood.
  - African dust may play an important fertilizing role in the Amazon rainforests on the long term
- (Yu et al., GRL, 2015)

# CALIOP vs. MERRA2



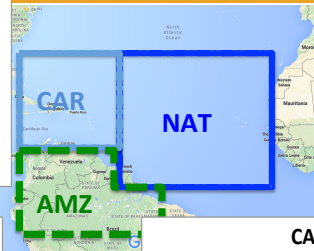
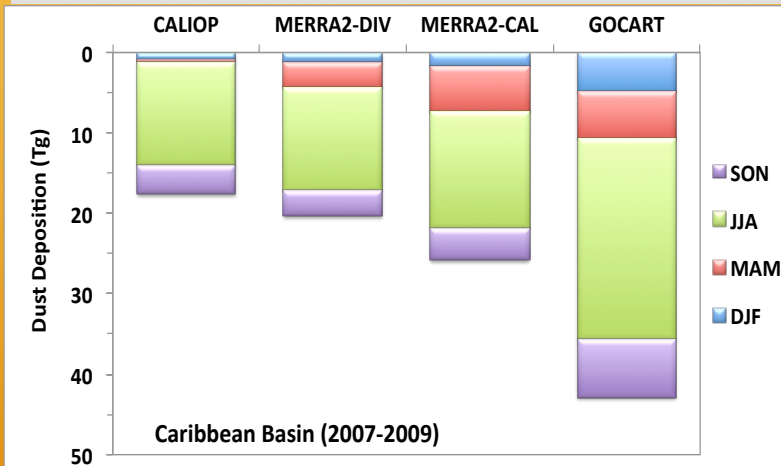


# Dust Deposition: CALIOP vs. MODELS

## Two types of dust deposition estimates

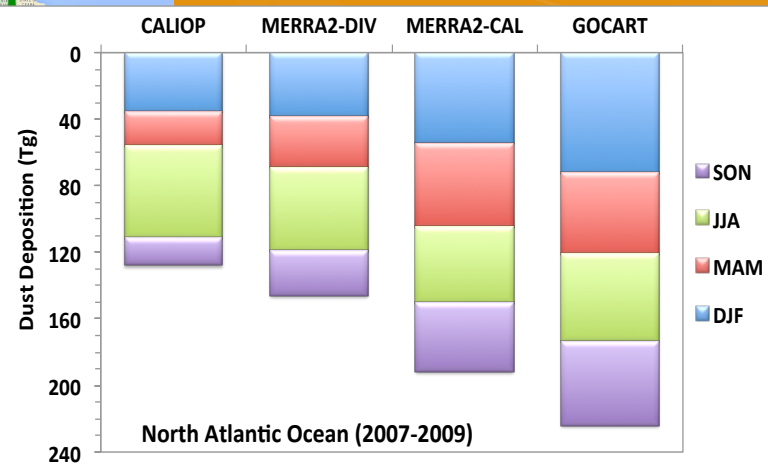
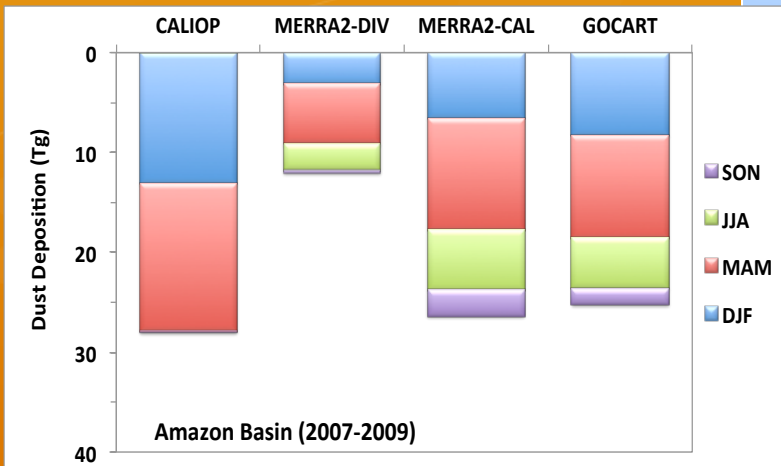
- CAL – based on para. of dry & wet removals
- DIV – diagnosed from flux divergence

	CAL	DIV
<b>CALIOP</b>		<b>X</b>
<b>MERRA2</b>	<b>X</b>	<b>X</b>
<b>GOCART</b>	<b>X</b>	



GOCART, CAL = DIV  
(mass balance)

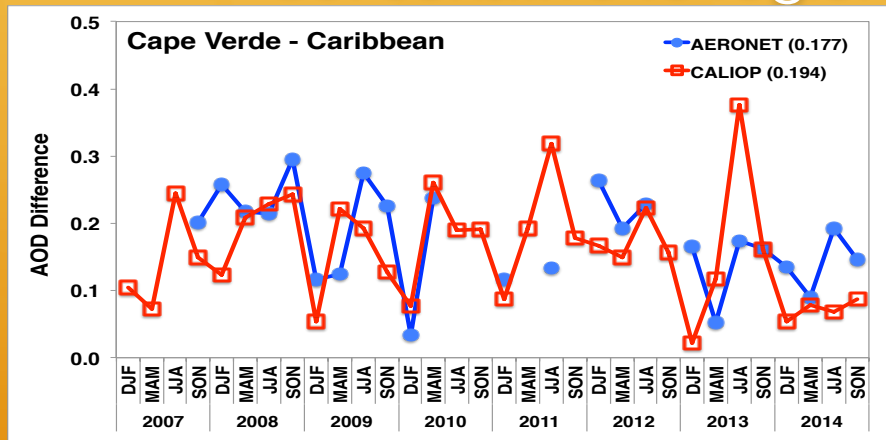
MERRA2, CAL > DIV if its  
removal mechanisms are  
too efficient (adtl. source)



# Dust Lifetime Estimate

$$\text{Dust Lifetime (d)} = \frac{[\text{DOD}/\text{MEE}] (\text{gm}^{-2})}{[\text{Dust Deposition Flux}] (\text{gm}^{-2}\text{d}^{-1})}$$

- Not sensitive to MEE
- But Sensitive to region-mean DOD to DOD gradient

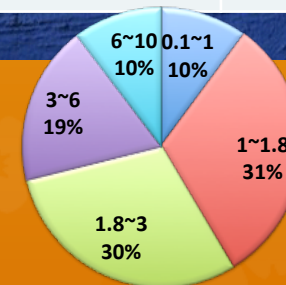


- Dust deposition into NAT & CAR is 123 Tg a<sup>-1</sup> or 6.7 gm<sup>-2</sup> a<sup>-1</sup>
- Regional mean DOD = 0.058
- Dust lifetime = ~ 9 days

## Estimated lifetime (d) from AeroCom models (Kim et al., 2014)

GOCART	GISS	SPRINTARS	ECHAM5	HadGEM2
4.9	5.0	4.2	5.3	5.7

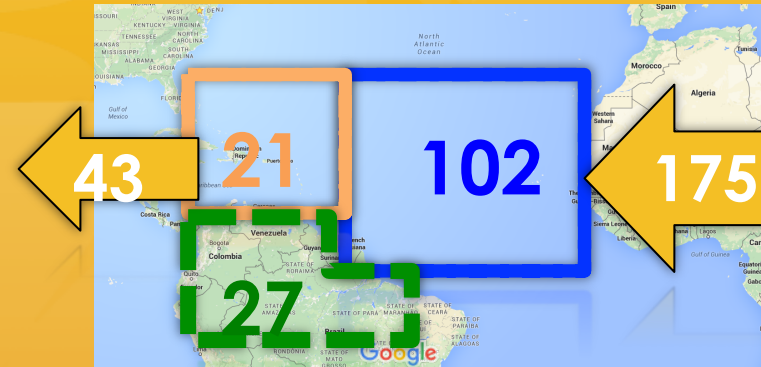
Caveat: CALIOP-based dust deposition is underestimated because of non-sensitivity to coarse particles



GOCART: 29% of dust dep. By particles with radius >3 um

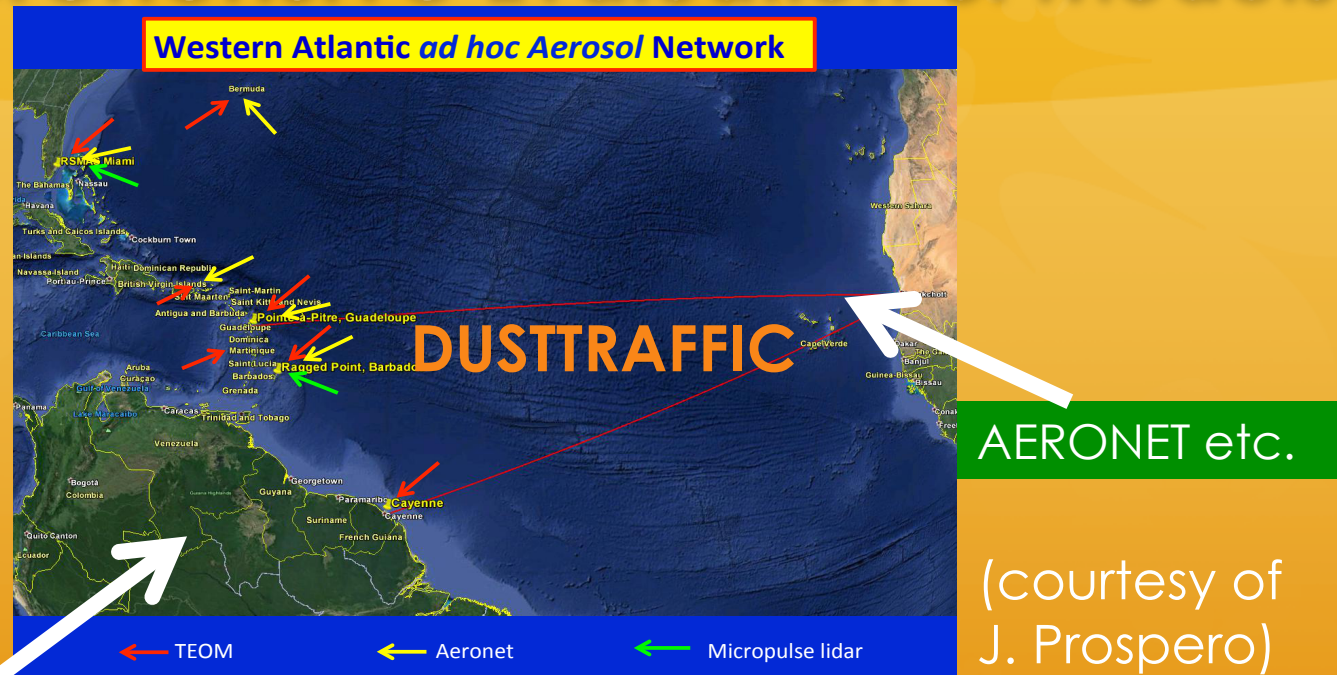
# CONCLUSIONS

- dust transport and deposition based on 8-yr CALIOP data record



- The CALIOP-based dust flux in the lowest 1km layer correlates with surface dust conc. in the Caribbean Basin ( $R=0.82-0.88$ )
- The CALIOP-based dust transport and deposition show modest correlation with prior-year Sahel rainfall. *Longer data record is needed for getting a robust relationship.*
- Significant differences in dust transport and deposition exist between CALIOP and MERRA2. MERRA2 shows larger east-west gradient than CALIOP. Further comparison of vertical profile is needed (on going).
- Dust lifetime based on CALIOP observations is estimated at ~9 days, significantly longer than AeroCom models. But caveats exist.

# Comprehensive Evaluation of Models



Long-term measurements  
(filter, lidar, AERONET) in the  
Amazon [Paulo Artaxo]

A suite of satellite measurements and derived  
products (e.g., passive & active, polar orbiting &  
geostationary)

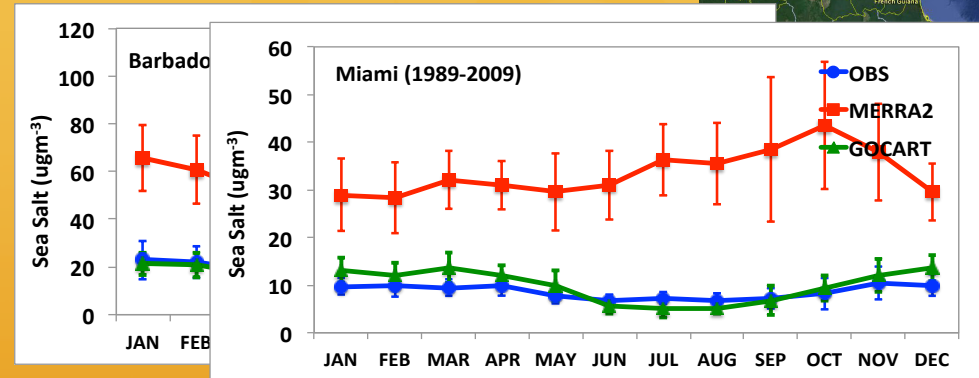
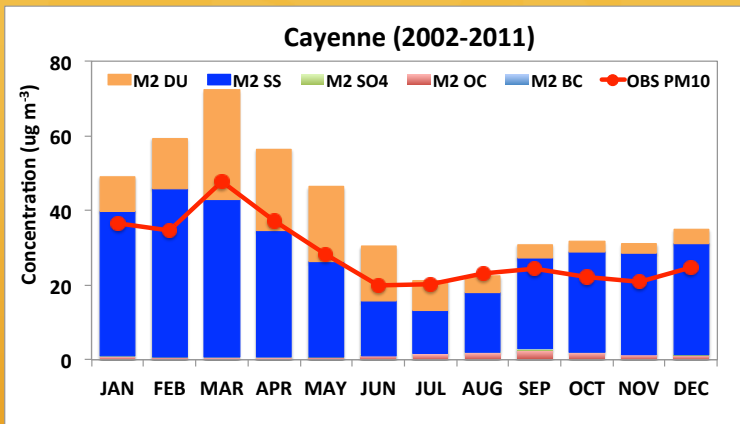
Thank you for your  
attention!

*Questions?*

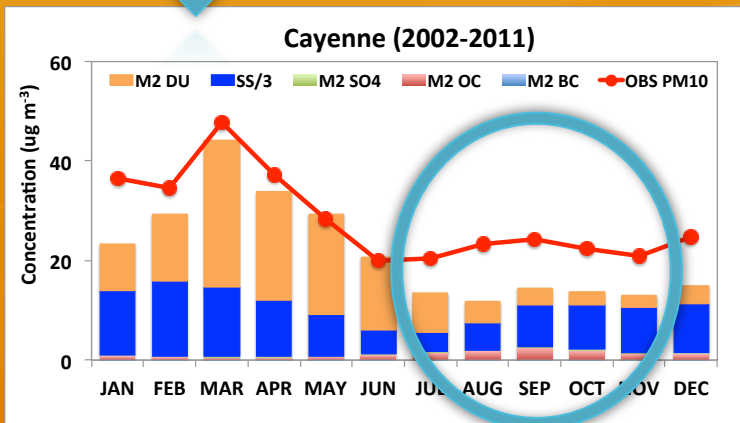
# Model Simulation of Dust Import into Amazon



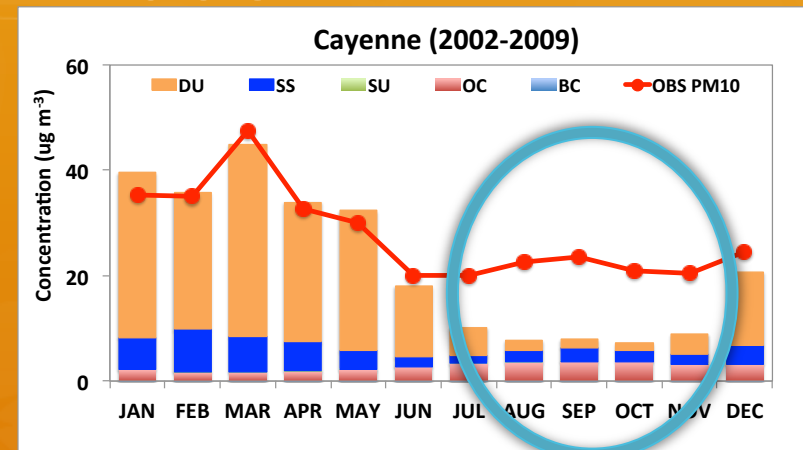
## MERRA2



Scaling sea-salt by 1/3



## GOCART



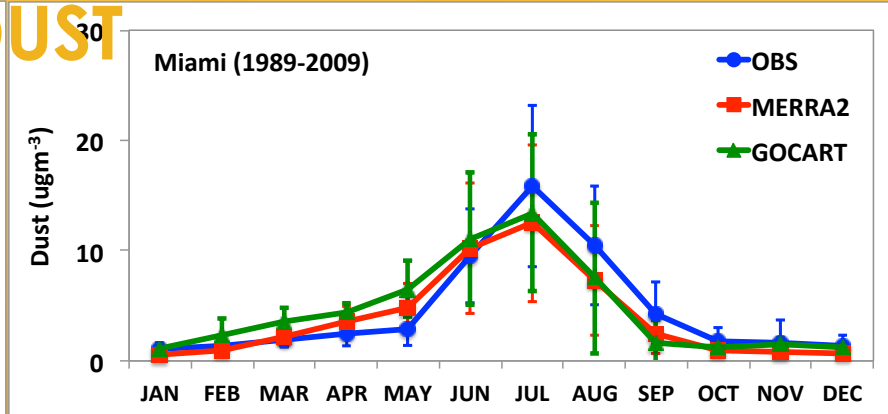
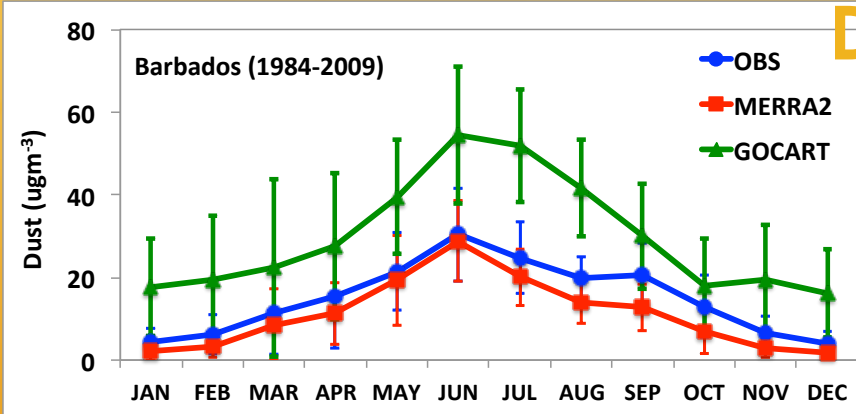
Missing sources in non-dust season (e.g., transport of smoke from Brazil)?

# MODELS vs. OBS.

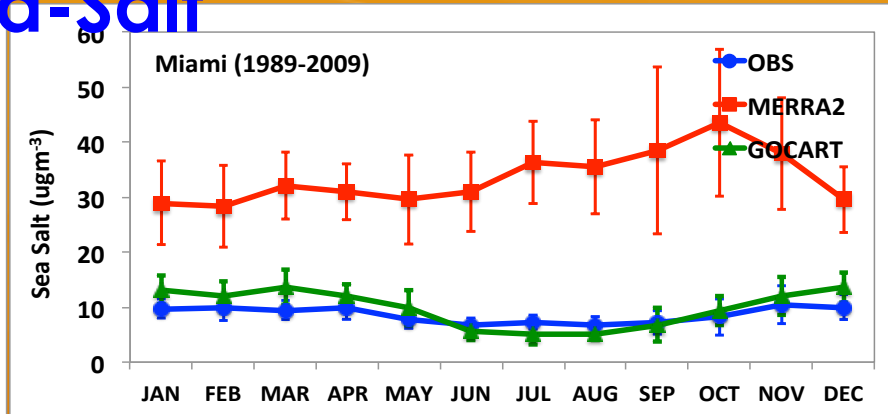
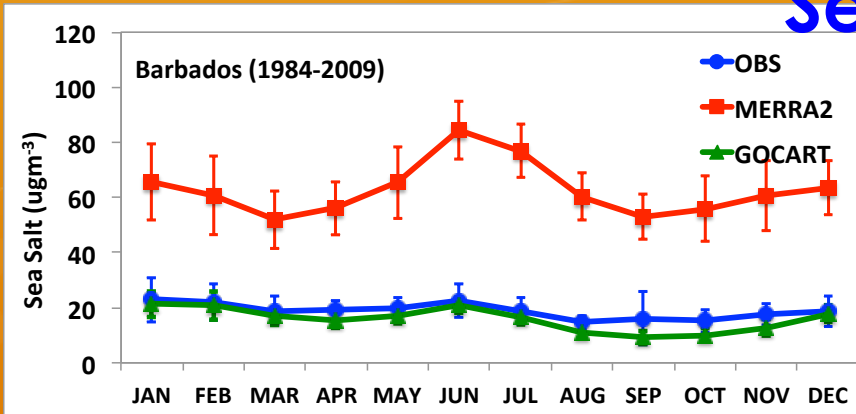
## Barbados

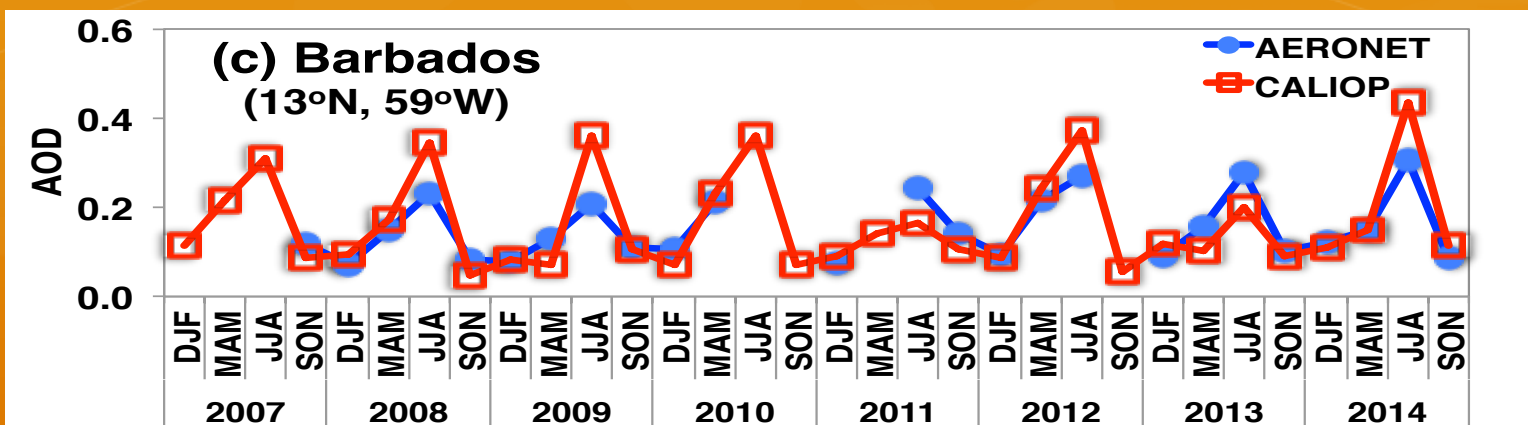
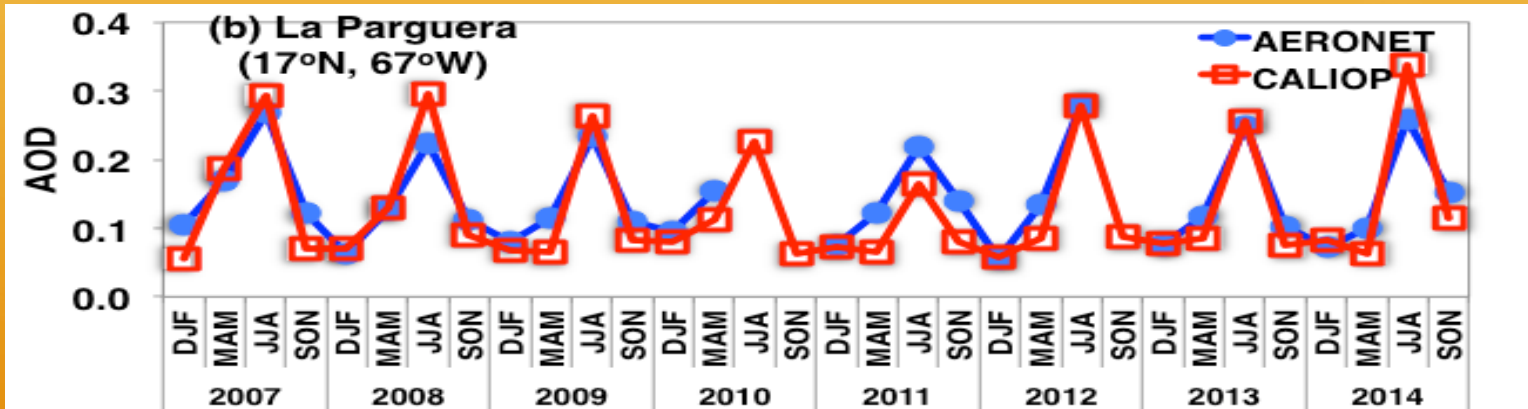
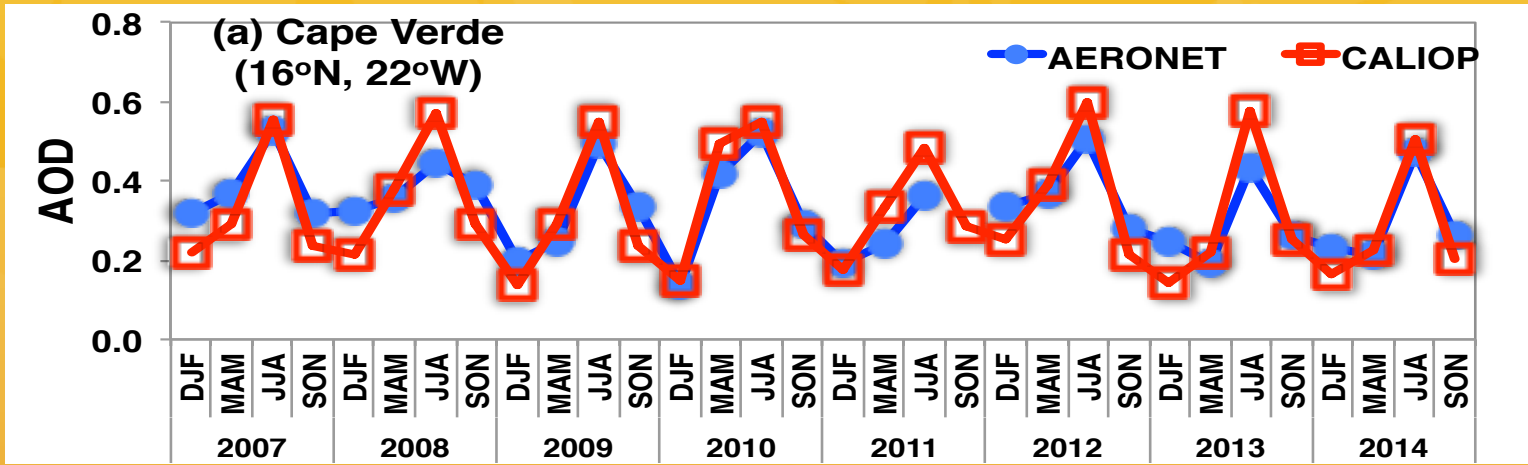
## Miami

### DUST



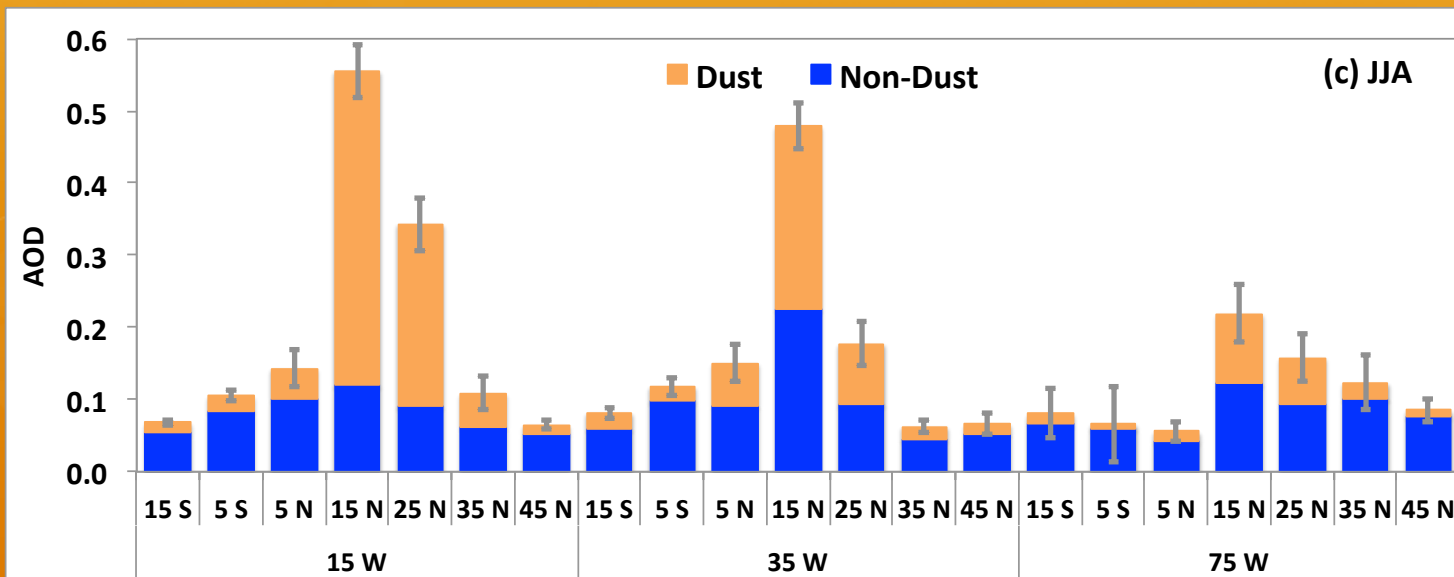
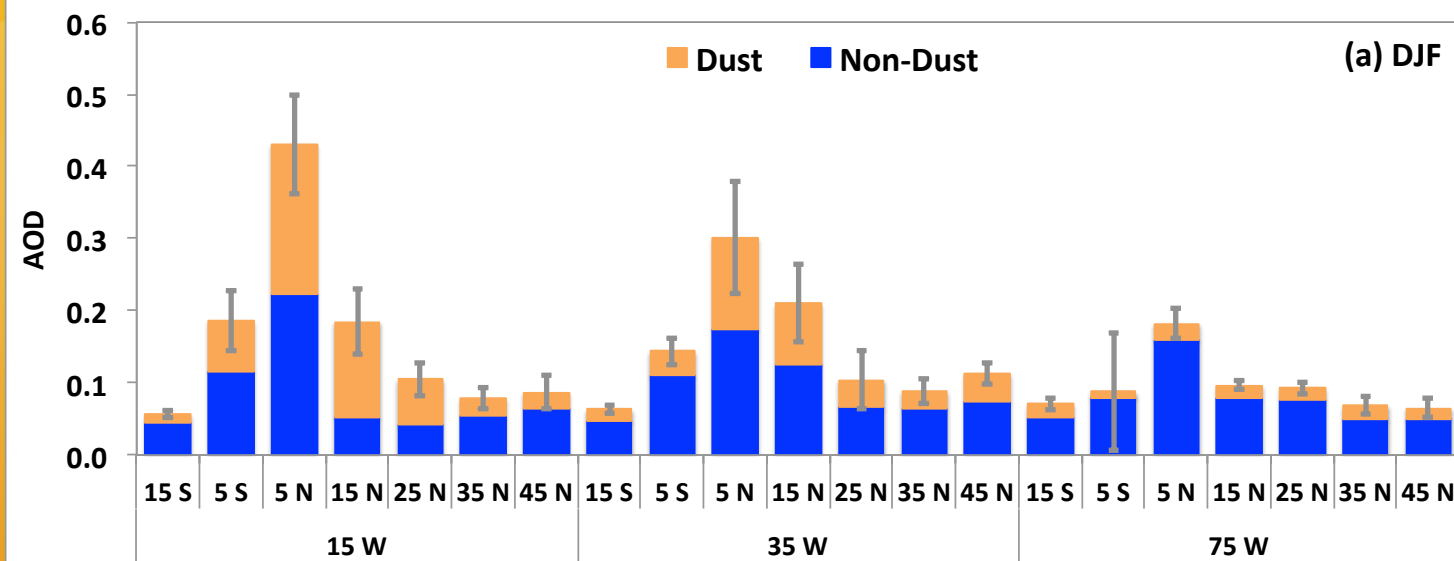
### Sea-Salt





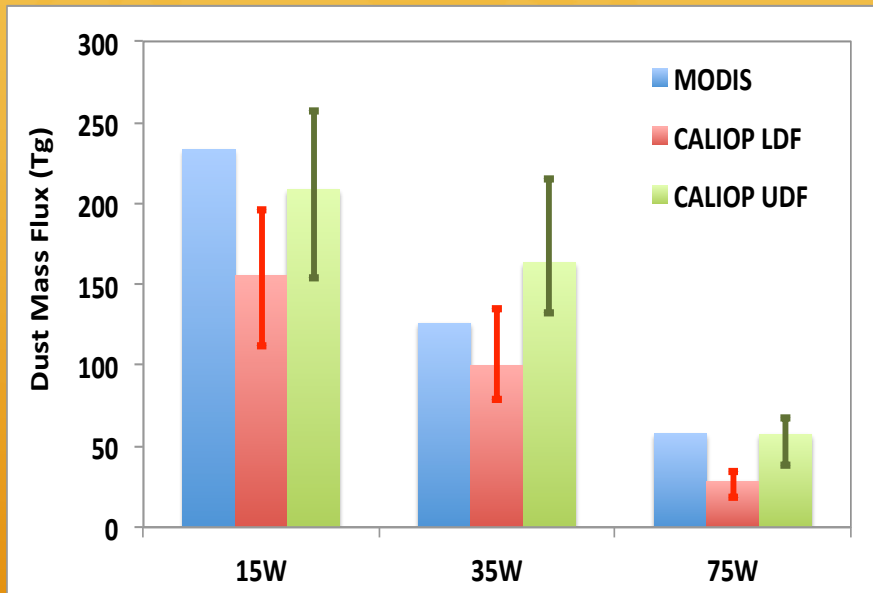


# Partition of Dust & Non-Dust



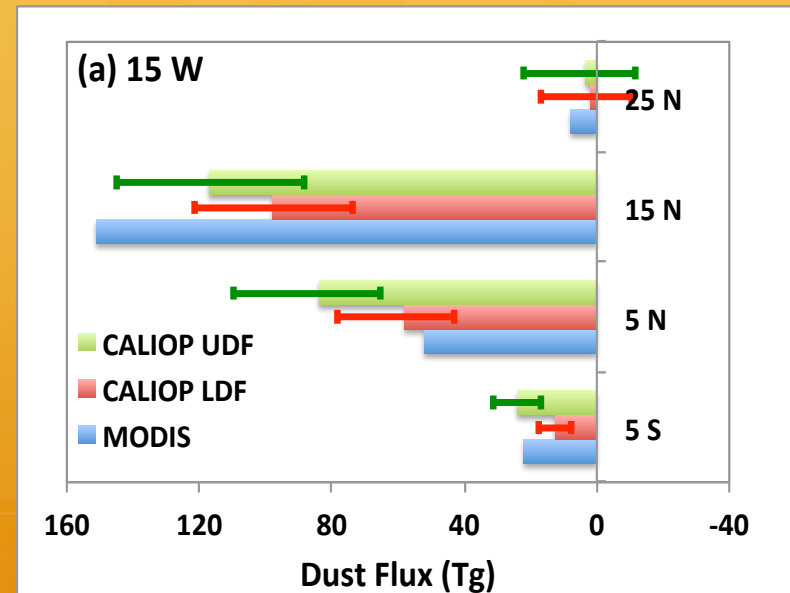
# CALIOP vs. MODIS

## 10°S-30°N Integrated



MODIS-based dust mass flux (Kaufman et al., 2005) generally agrees well with CALIOP upper-bound estimate.

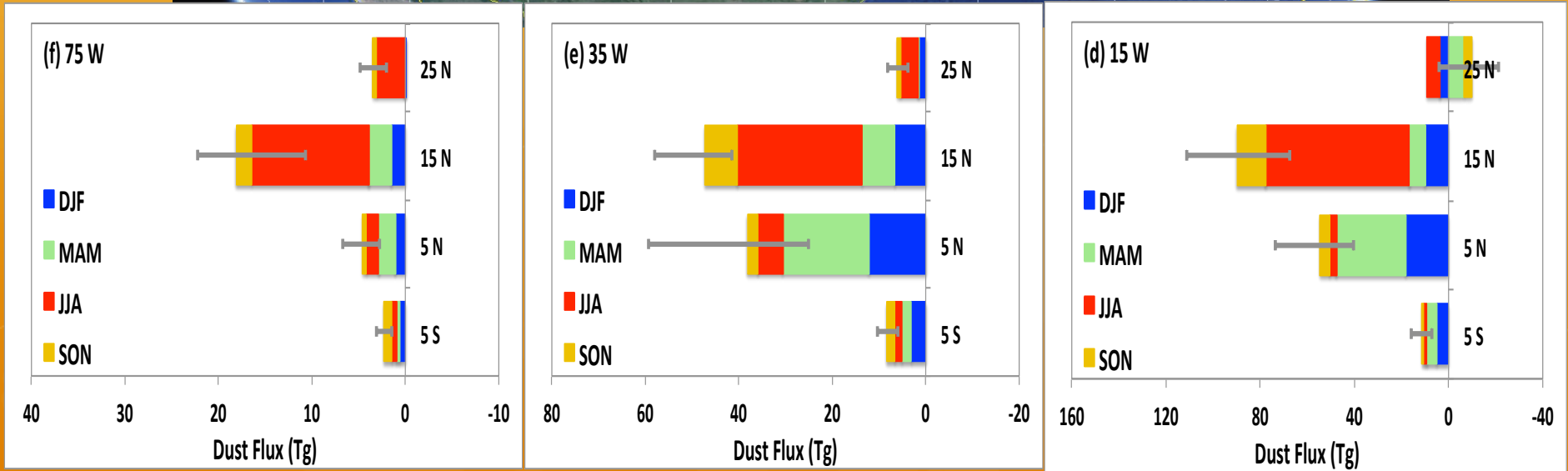
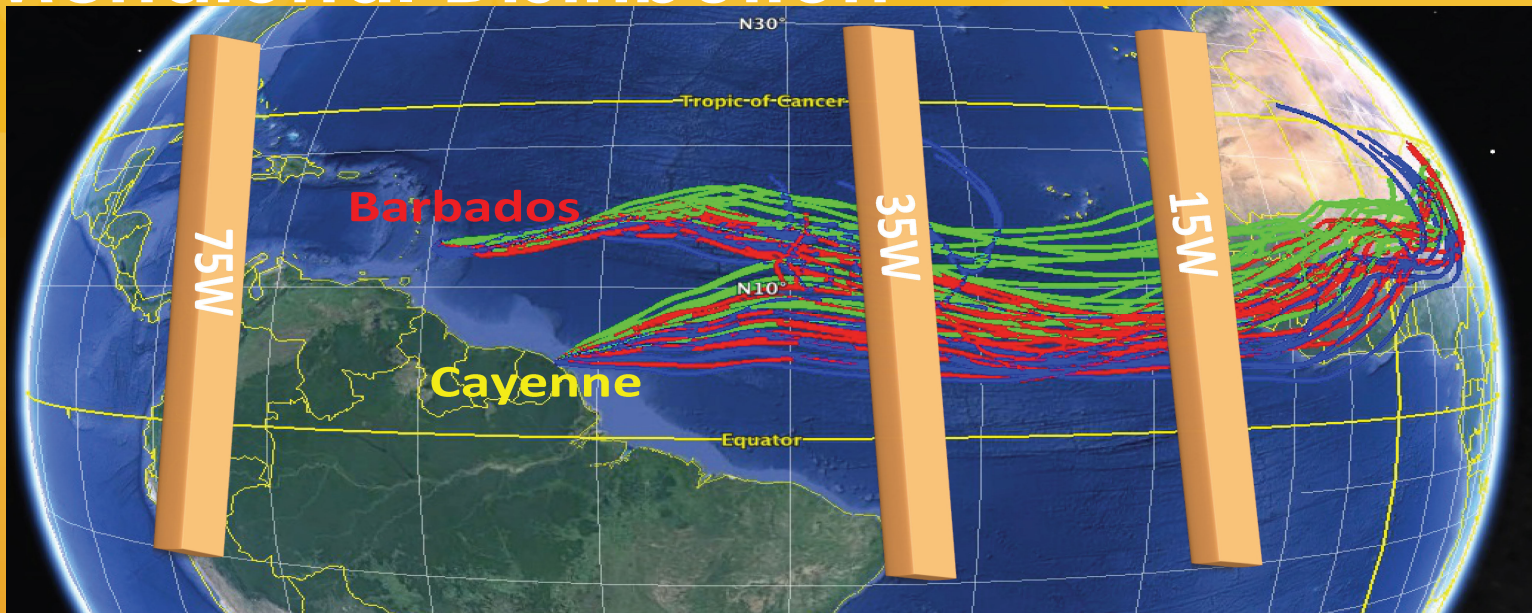
## Meridional Distribution



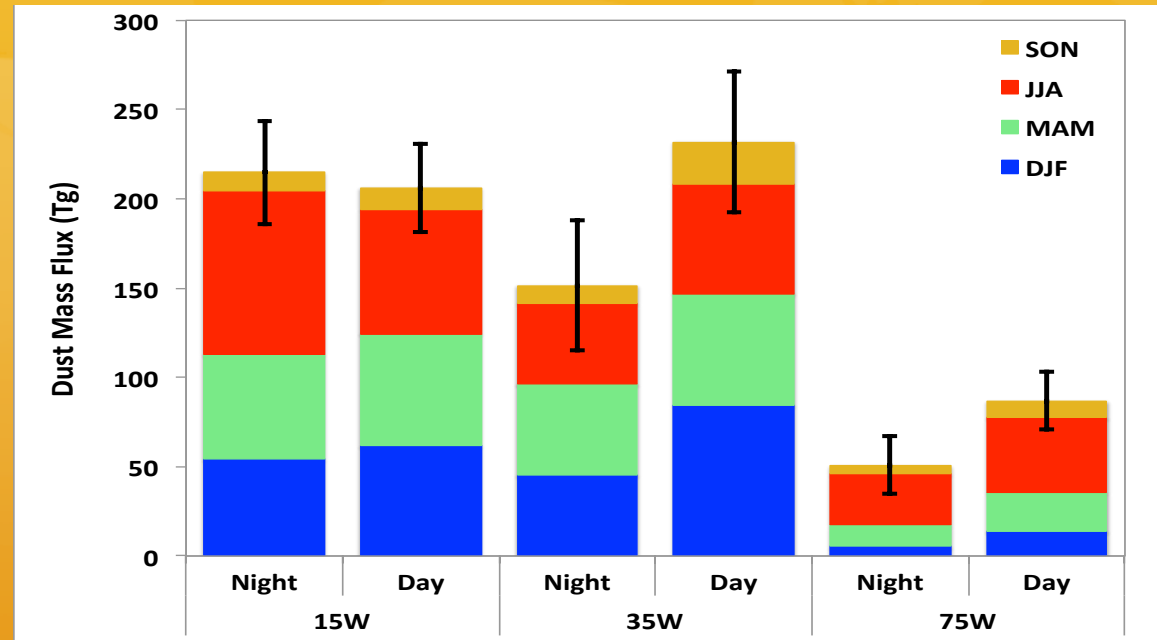
Larger differences in the meridional distribution.

*Error bar indicates the range over the 7-year period.*

# Meridional Distribution



# Daytime vs nighttime



- ❑ But we can't attribute the day-night difference to physical processes, because CALIOP daytime and nighttime data have different quality.
- ❑ AERONET AOD in trans-Atlantic dust route shows small daytime variations (*Smirnov et al, 2002; Zhang et al., 2012*)