

# The Atmospheric Circulation as a Key Mediator of Aerosol-driven Climate Impacts

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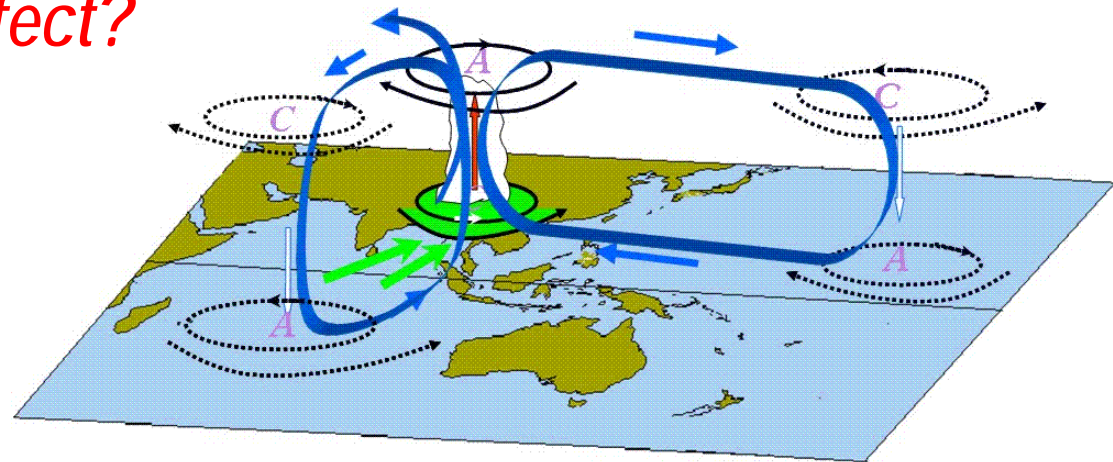
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## Summer Broad-Scale Circulations

*Is there a downstream effect?*

*Why?*

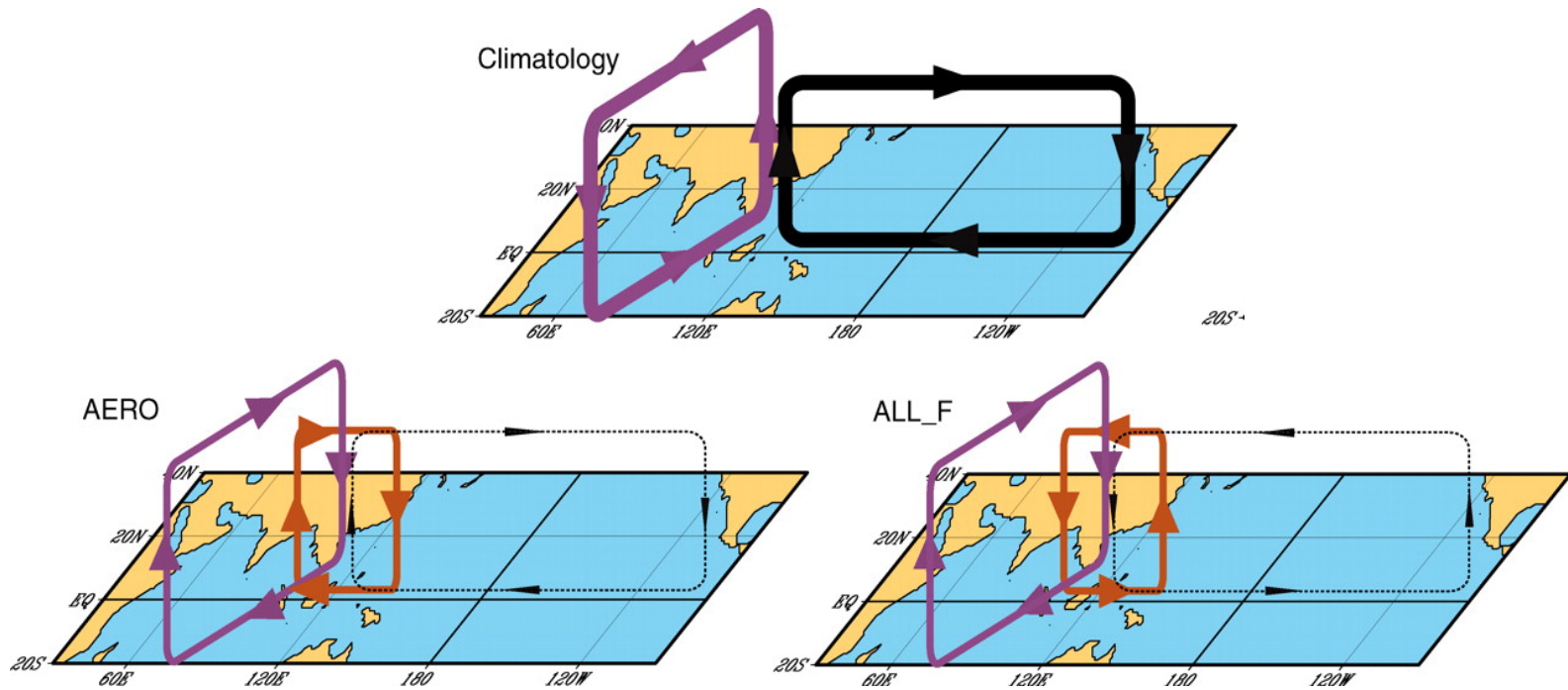


# Beyond the local response ...

*Aerosol increase cause a decline in summer precipitation over SA*

**Aerosol-precipitation-circulation** interactions are poorly known. Atmospheric circulation changes may extend the aerosol impact beyond the source region.

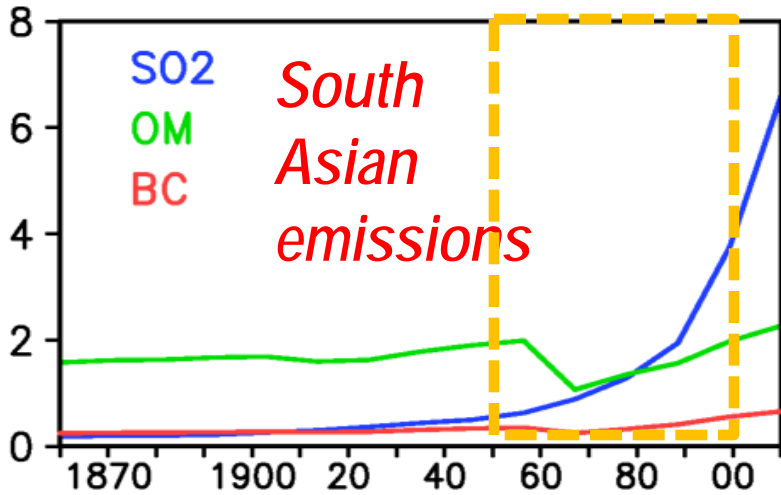
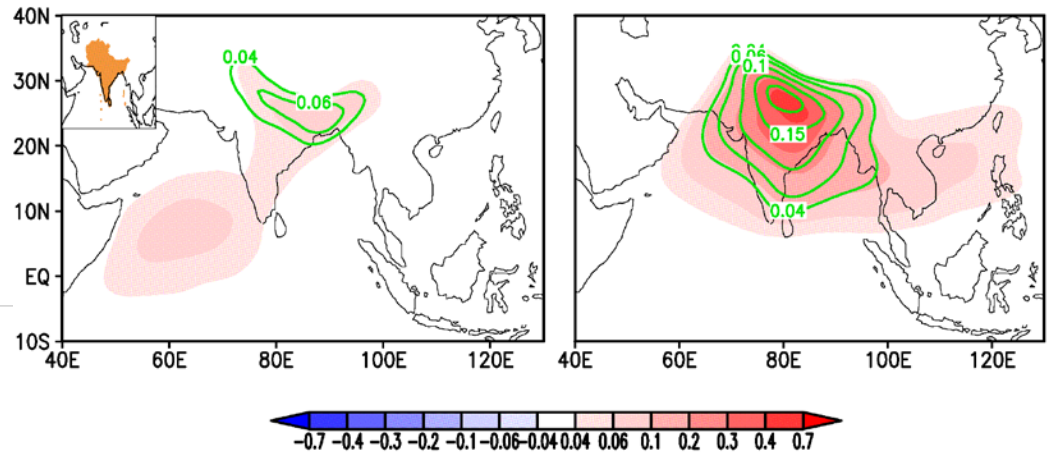
Changes in deep convection lead to strong feedbacks with the dynamics



# Remarkable increase in anthropogenic aerosols from 1950s

Emission increase (x-fold)		
	South Asia	East Asia
SO <sub>2</sub>	10	16
BC	2	5

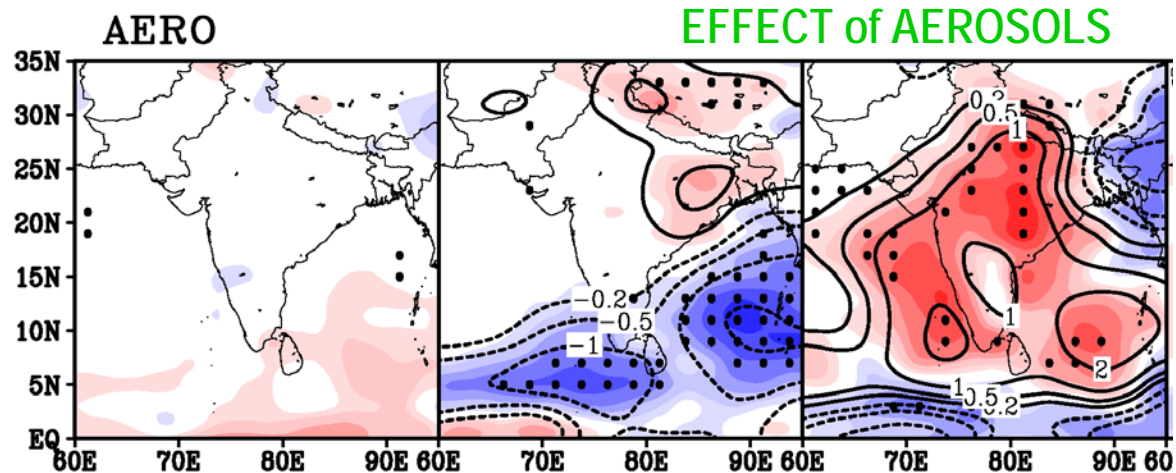
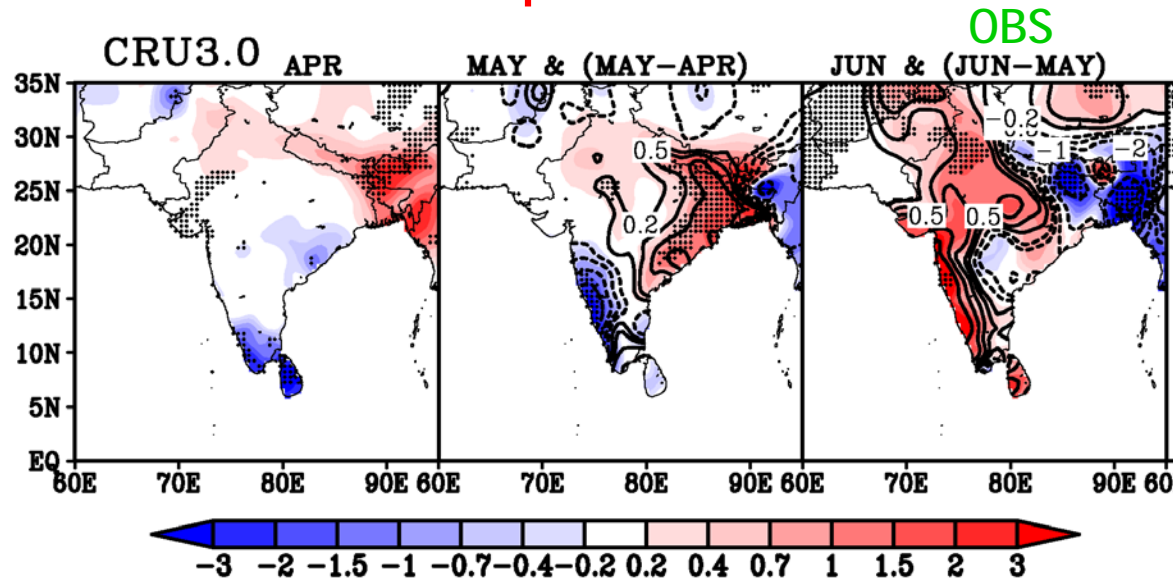
**DJF** **JJAS**  
*AOD : Total (shades) & absorbing (contours, x10)*



# Aerosol-induced changes in the atmospheric circulation during the onset phase

## Precipitation trend

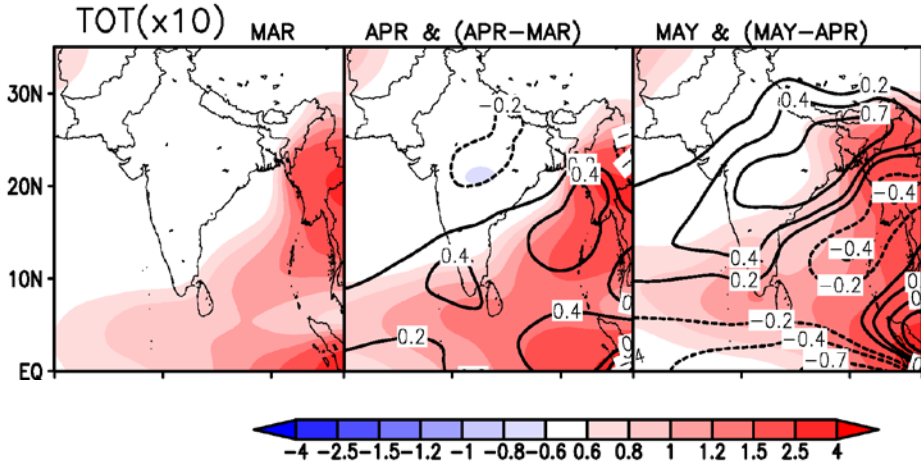
Why the onset?



# Increased aerosols lead to large regional energy deficit

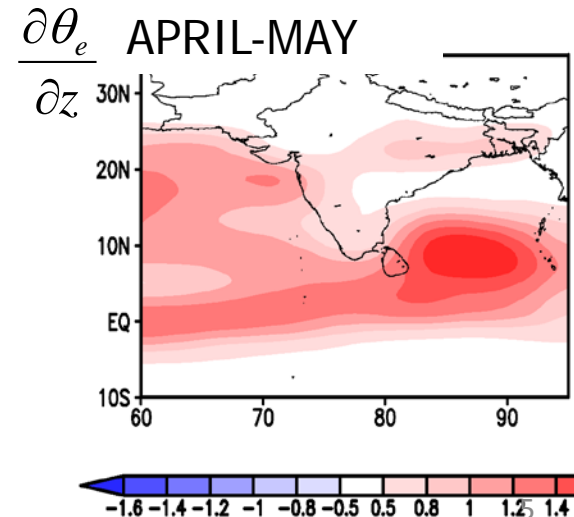
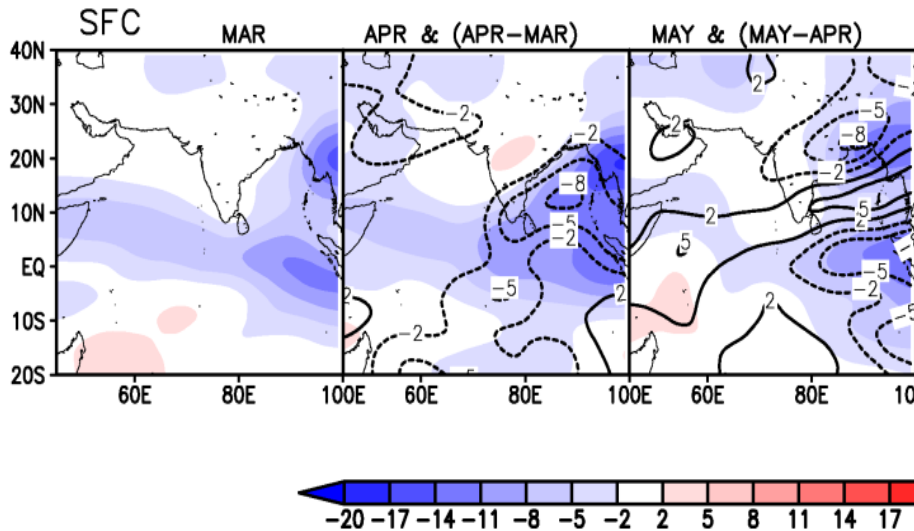
ALL AEROSOLS

AOD trend

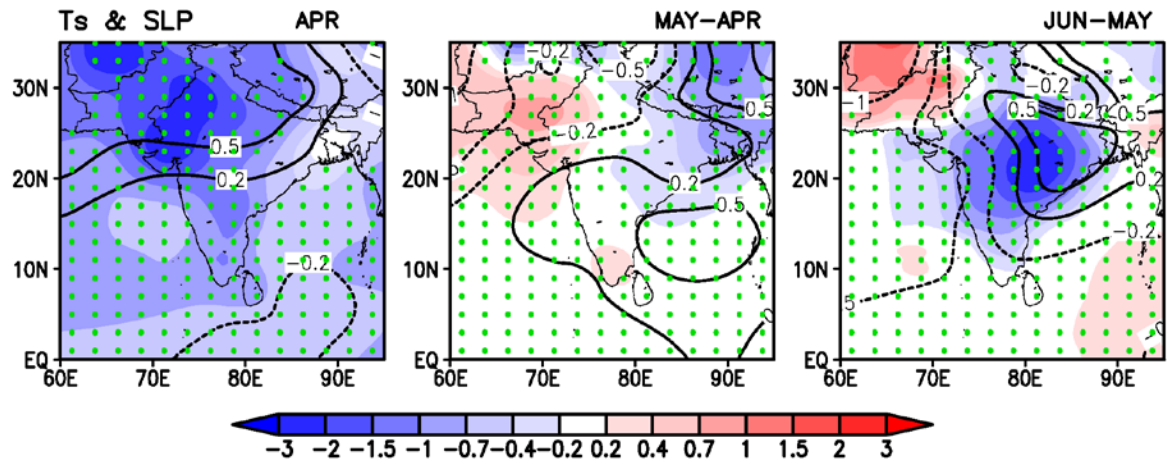


Radiative Flux Perturbation ( $\text{W m}^{-2} (50 \text{ y})^{-1}$ )

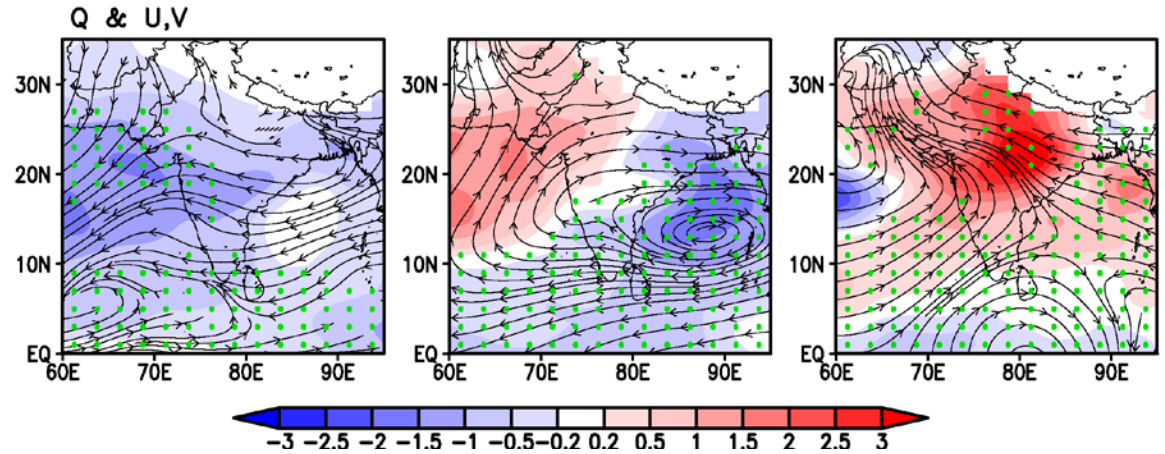
⇒ Increased atmos. stability



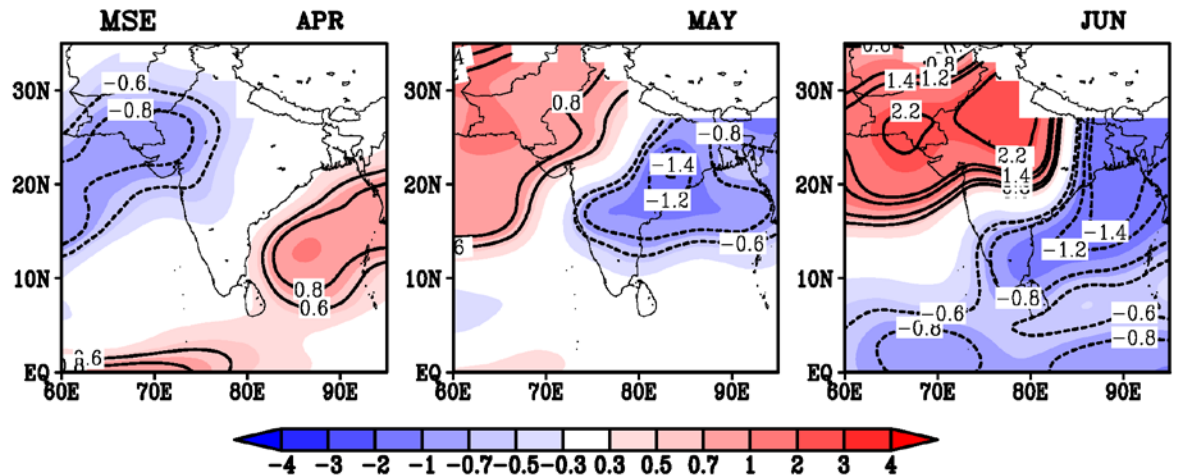
# Surface Temperature & Sea Level Pressure



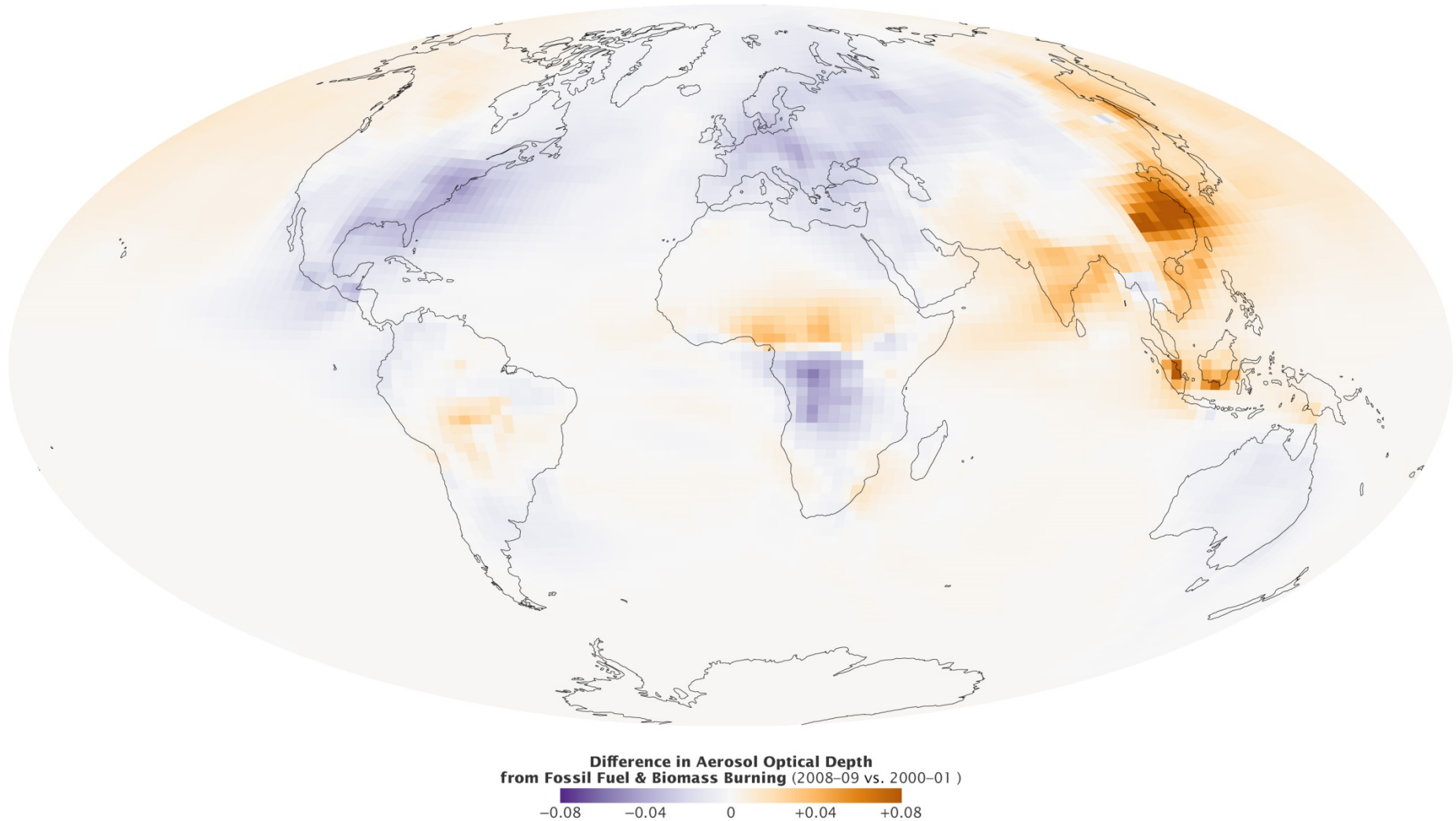
# Spec. Humidity & 850-hPa Winds



# Sub-cloud Moist Static Energy

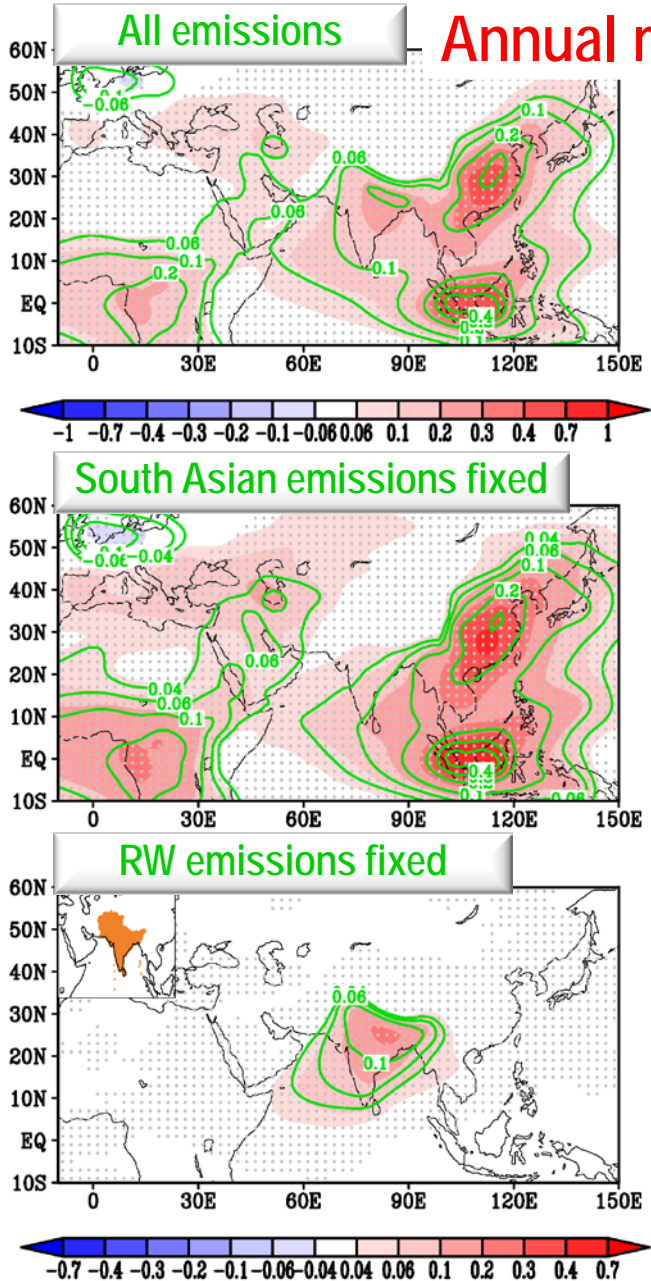


# A contrasting pattern of aerosol changes



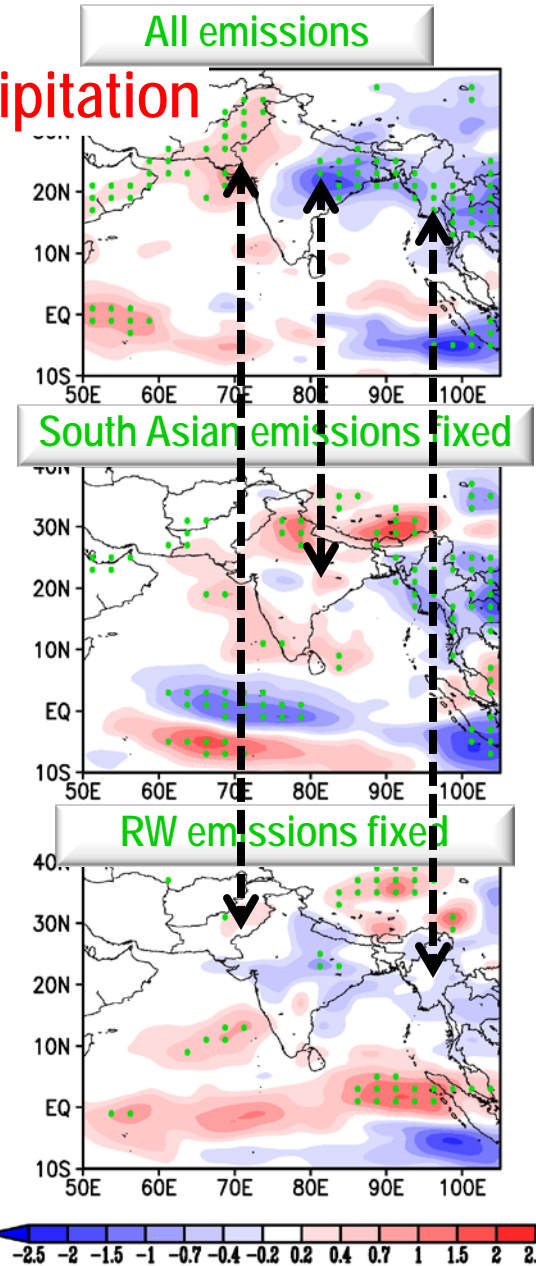
Changes in anthropogenic aerosols between 2000-01 and 2008-09 (Chin et al. 2014)

# Contrasting local and remote aerosol forcing



## JJAS Precipitation

GFDL CM3  
All-Forcing experiments



### Core monsoon India

$$ALLF = -0.81$$

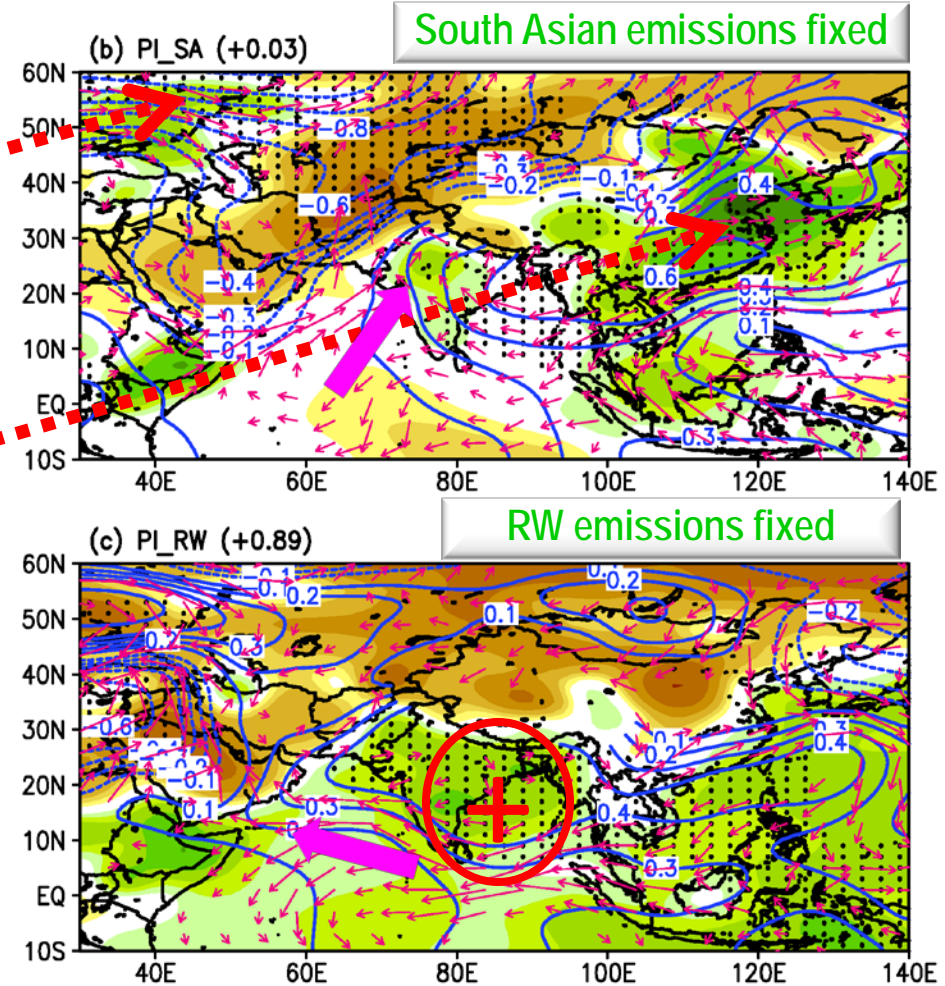
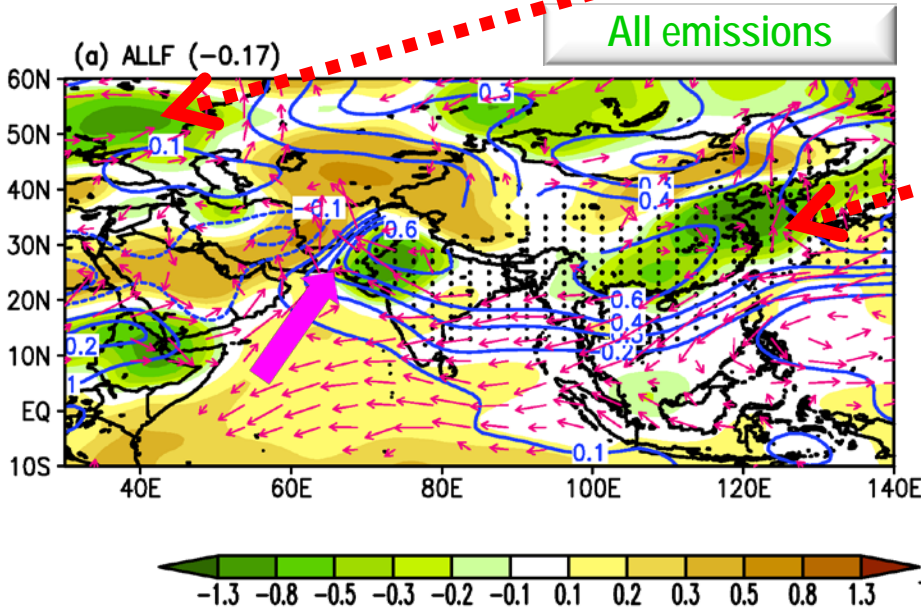
$$PI_{SA} = +0.24$$

$$PI_{RW} = -0.41$$



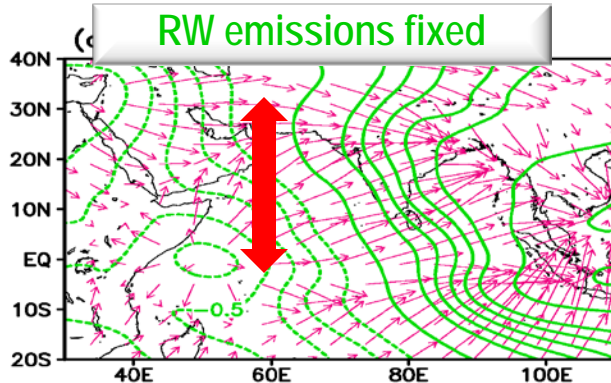
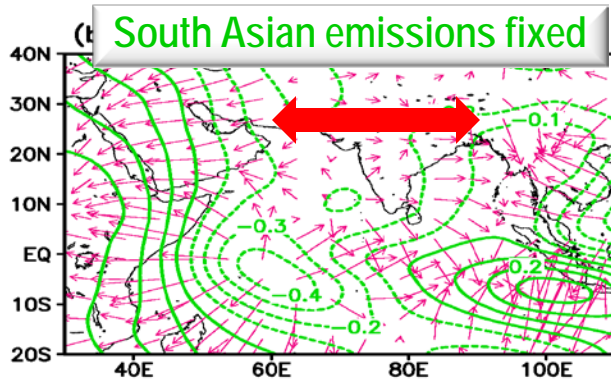
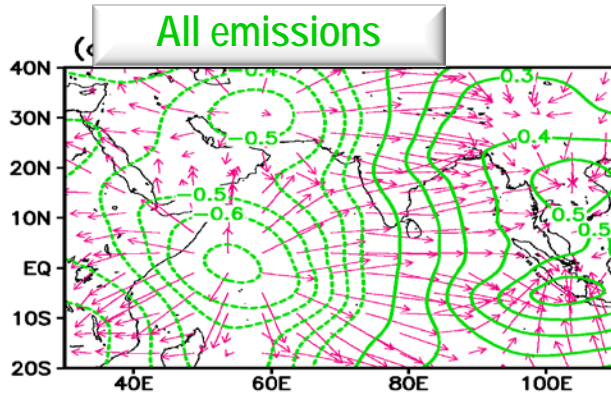
# Different patterns in near-surface changes

Ts (shades), SLP (contours), 850-hPa winds  
(minus domain-average change)



# Implications for the large-scale circulation

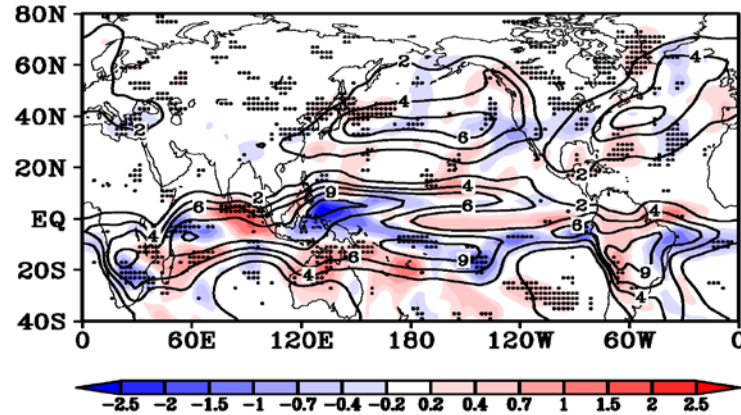
## JJAS 200-hPa velocity potential & divergent circulation



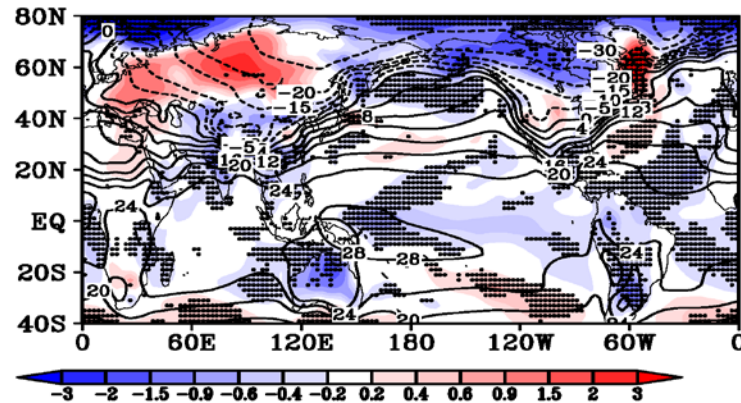
# Potential large-scale impact of SA aerosols

Precipitation (mm/day)

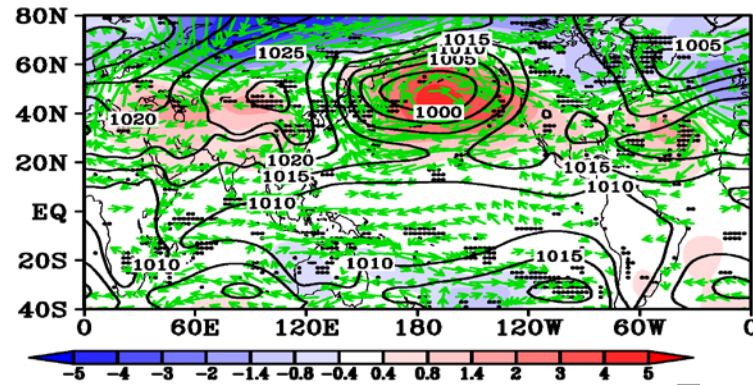
DJF



Sfc. Temperature (K)



SLP (hPa) &  
850-hPa winds (m/s)



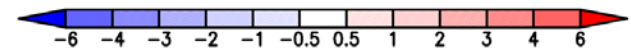
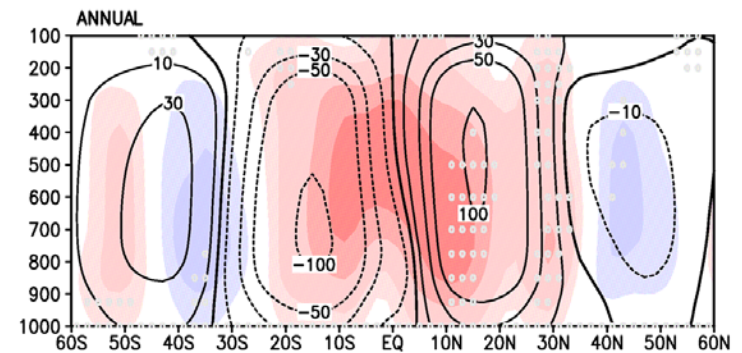
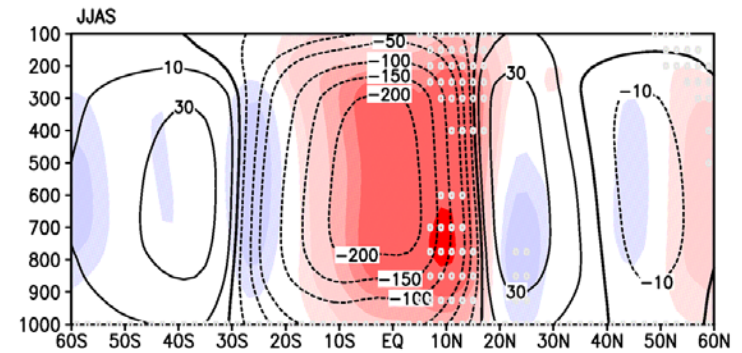
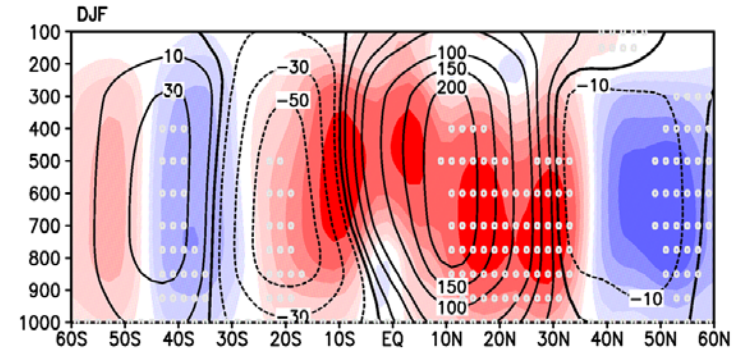
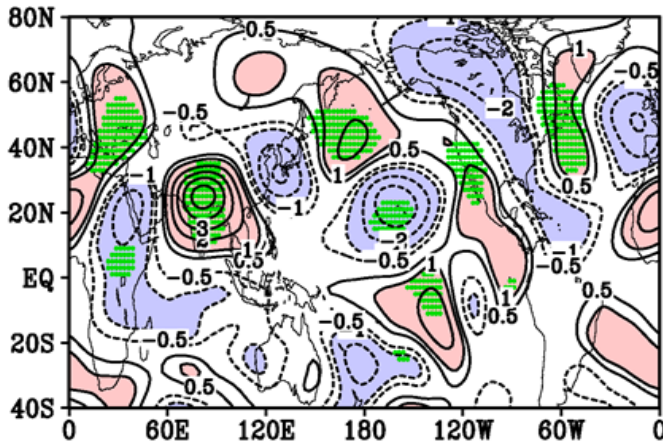
# Atmospheric circulation response

Mass meridional streamfunction ( $\times 10^9 \text{ Kg/s}$ ):  
1950s values (black) and changes (shades)

*Wave-like patterns across the  
Northern Hemisphere*

DJF

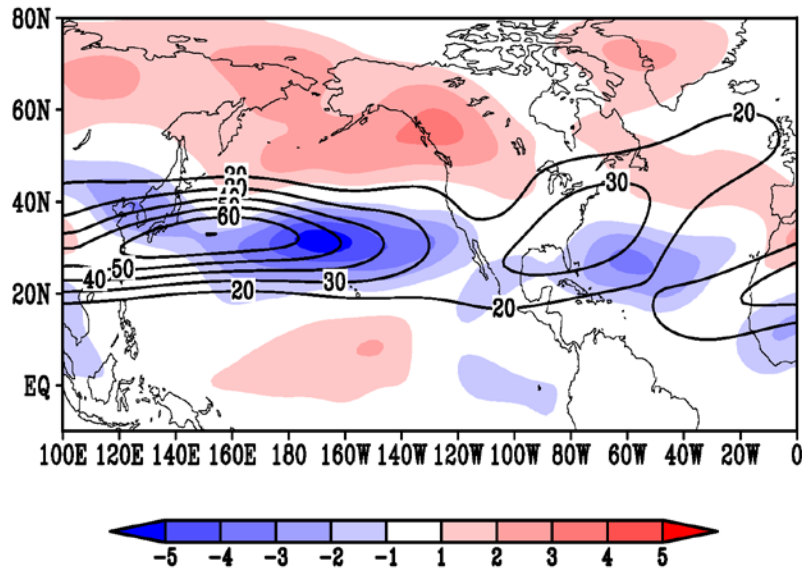
*200-hPa Stat. eddy streamfunction ( $10^6 \text{ m}^2/\text{s}$ )*



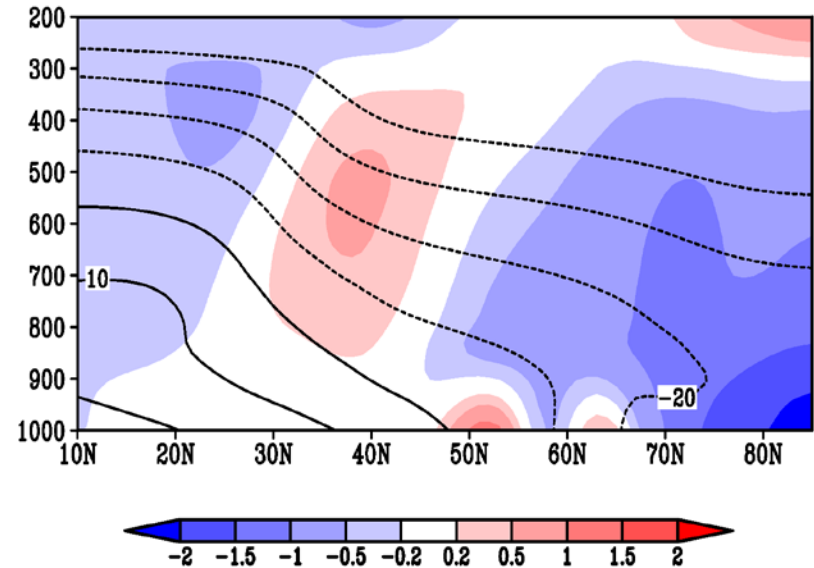
# Boreal winter jet

*Northward shift of the Atlantic-Pacific jet associated with changes in the tropospheric meridional temperature gradient*

**300-hPa zonal wind (m/s)**



**Temperature (K)**



# Concluding remarks

- ❑ Circulation adjustments *following* local aerosol forcing are key players
- ❑ Can extend the aerosol impact downwind (teleconnections) & lead to further +/- feedbacks
- ❑ Complex and varied interplay among local, regional, remote aerosols

## Next step (Fall/Winter)

Contrasting **Europe + North America** vs. **South Asia**  
vs. **East Asia** in transient/time slice experiments

# Key Issues

- ❑ Observations (Poor coverage and limited data)
- ❑ Scale dependency (Spatial and temporal, e.g., interannual vs. decadal, spatial heterogeneity in patterns of change)
- ❑ Characteristics (e.g., extremes vs. seasonal)
- ❑ Timing (e.g., seasonal cycle - onset vs. withdrawal)
- ❑ Upstream/downstream effects (remote effects/teleconnections)
- ❑ Challenges in simulations (overall large biases, incremental improvement in CMIP5; mean state vs. changes?)
- ❑ Forcing uncertainties (On top of GHG: Aerosols? Natural variability? Linearity? Land use?)