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Constraining Indirect Effects on the Process Level

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Model physics and chemistry in the GFDL AM3 Model (to be used for AR5)

Aerosol-Liquid Cloud Interactions

A prognostic scheme of cloud droplet number concentration (Ming et al., 2007) with an explicit treatment of aerosol activation at cloud base (Ming et al., 2006).

Convection Parameterization

Move from the <u>relaxed Arakawa-Schubert (RAS)</u> in AM2 to the <u>Donner deep convection scheme</u> (Donner, 1993) and the <u>University of Washington (UW) shallow</u> <u>convection scheme</u> (Bretherton et al., 2003). By providing in-plume updraft velocity, the latter two are ideal for implementing aerosol/cloud microphysics. •Online aerosol transport and tropospheric and stratospheric chemistry

Anthropogenic aerosol radiative flux perturbation (RFP, W m⁻²) at TOA from preindustrial to present-day

	AM3 (to be used for AR5)	AM2 (used for AR4)
Direct effects – Sulfate and organic carbon	0 (assuming	-1.3 (external mixing)
Direct effects - Black carbon	internal mixing of sulfate and black carbon)	0.5 (external mixing)
Indirect effects	-1.3	Not included

Aerosol direct effects Importance of atmospheric absorption

TOA – all-sky $0 W m^{-2}$ TOA – clear-sky -1.0 $W m^{-2}$





Surface – all-sky -1.3 W m⁻²

-30. -10. -7.5 -5.0 -3.0 -1.0 0.0 1.0 3.0 5.0 7.5 10. 30.

A sanity check on aerosol absorption

Comparison with AERONET measurements of co-albedo



AM2

AM3

1st and 2nd aerosol indirect effects

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A preliminary simulation of the 20th century climate using the IPCC emissions

Global Annual-Mean SURFACE TEMPERATURE (K) (referenced to 1881-1920 average)



An approach to narrowing down the uncertainties in aerosol indirect effects

Dissecting indirect effects on the process level



A theoretical constraint on cloud susceptibility

Revisiting Twomey's Cloud Susceptibility:

$$\frac{\partial R}{\partial N}\Big|_{l} = \frac{R(1-R)}{3N}$$

In discrete form, the <u>relative cloud susceptibility</u> can be written as:

$$\frac{\Delta R}{\left(\frac{\Delta N}{N}\right)} \bigg|_{l} = \frac{R(1-R)}{3} \longrightarrow \frac{\Delta R}{\left(\frac{\Delta N}{N}\right)} \bigg|_{l,\max} = \frac{1}{12} = 8.3\%$$
when $R = 0.5$

Comparison of AM3-simulated cloud albedo susceptibility with satellite data

Difference in cloud albedo (x1000) caused by a uniform 10% increase in droplet number in July.



Oreopoulos and Platnick (2008)

Dissecting 2nd Indirect Effect on the Process Level Advection PI/PD Dry **Dry/Wet** Emissions Aerosols Removal **Evaporation** Activation Cloud Cloud Liquic **Droplets** Cloud **Albedo** 2nd IE

How Droplet Number Affects Cloud Liquid?

The governing equation for cloud liquid can be simplified into

$$\frac{\partial l}{\partial t} = S - \frac{\partial l}{\partial t}_{auto} = S - a \frac{l^m}{N^n}$$

Sources/Sinks except autoconversion; assumed to be constant w.r.t. droplet number (?).

At steady state,
$$\frac{\partial l}{\partial t} = 0$$

Then,
 $\ln\left(1 + \frac{\Delta l}{l}\right) = \frac{n}{m}\ln\left(1 + \frac{\Delta N}{N}\right)$

In this model configuration, n/m is less than 1/7 (0.14).

Model-Simulated Dependence of Cloud Liquid on Droplet Number



Aerosols are important for understanding regional climate change

Zonal-mean responses to <u>aerosol direct and indirect</u> <u>effects</u> simulated with a slab ocean model

Surface temperature (K)

Precipitation (mm day-1)



Ming and Ramaswamy (2009)

How does the tropical heat engine response to aerosol forcing, and why? From the viewpoint of atmospheric energy transport, the response gives rise to a cross-equatorial heat flux from SH to NH.



Concluding remarks

•In AM3, a prognostic scheme of droplet number establishes a physical link between aerosols and clouds;

•Theories, models and measurements are used to better constrain indirect effects;

•Aerosol-induced circulation changes need to be studied more thoroughly.

