Radiative forcing of black carbon and mineral dust deposited to snowpack

Mark Flanner

September 30, 2010 AeroCom Meeting, Oxford



Outline

1 Calling all BC and dust deposition fields!

2 Background

- 3 AeroCom Experiments
- 4 Future Directions

Snow darkening from BC and other "snowsols"

Snow darkening from BC and other "snowsols"



Part-per-billion levels of BC significantly reduce snow albedo because:

Snow darkening from BC and other "snowsols"



Part-per-billion levels of BC significantly reduce snow albedo because:

 $\bullet\,$ Black carbon visible absorptivity is $\sim 10^5$ greater than ice

Snow darkening from BC and other "snowsols"



Part-per-billion levels of BC significantly reduce snow albedo because:

- ullet Black carbon visible absorptivity is $\sim 10^5$ greater than ice
- Snow scatters visible radiation efficiently via refraction
 - A typical reflected green photon undergoes \sim 1000 scattering events before emerging from the top of snowpack

Snow darkening from BC and other "snowsols"



Part-per-billion levels of BC significantly reduce snow albedo because:

- ullet Black carbon visible absorptivity is $\sim 10^5$ greater than ice
- Snow scatters visible radiation efficiently via refraction
 - $\bullet~$ A typical reflected green photon undergoes $\sim 1000~{\rm scattering}$ events before emerging from the top of snowpack
- Longer persistence in near-surface snow than atmosphere.

Albedo perturbation from impurities



Albedo perturbation from impurities



• Simulate it yourself at: http://snow.engin.umich.edu

Springtime uniqueness



• Solar energy incident on snowpack peaks in March-May

Springtime uniqueness



- Solar energy incident on snowpack peaks in March–May
- This is also the season of maximum albedo feedback strength: d(albedo)/dT (Hall and Qu, 2006, Fernandes et al., 2009)

Global-scale studies on snow darkening

The Snow, Ice, and Aerosol Radiative (SNICAR) model, coupled with the NCAR CAM/CLM model

- 5 spectral bands
- 5 vertical snow layers
- Snow aging/microphysics model (Flanner and Zender, 2006)
- Particle removal with meltwater
- BC optical properties from Bond and Bergstrom (2006)

Global-scale studies on snow darkening

The Snow, Ice, and Aerosol Radiative (SNICAR) model, coupled with the NCAR CAM/CLM model

- 5 spectral bands
- 5 vertical snow layers
- Snow aging/microphysics model (Flanner and Zender, 2006)
- Particle removal with meltwater
- BC optical properties from *Bond and Bergstrom* (2006)
- Coupled atmosphere-land aerosol treatment (via deposition) (e.g., Rasch et al, Mahowald et al, Liu et al) ...

Global-scale studies on snow darkening

The Snow, Ice, and Aerosol Radiative (SNICAR) model, coupled with the NCAR CAM/CLM model

- 5 spectral bands
- 5 vertical snow layers
- Snow aging/microphysics model (Flanner and Zender, 2006)
- Particle removal with meltwater
- BC optical properties from Bond and Bergstrom (2006)
- Coupled atmosphere-land aerosol treatment (via deposition) (e.g., *Rasch et al, Mahowald et al, Liu et al*) ... **or**:
- Drive CLM-offline with aerosol deposition fields from external sources (e.g., AeroCom)

Spatial/temporal characteristics of BC/snow forcing



 Forcing operates mostly in local springtime, when and where there is large snow cover exposed to intense insolation, coincident with peak snowmelt

Spatial/temporal characteristics of BC/snow forcing



- Forcing operates mostly in local springtime, when and where there is large snow cover exposed to intense insolation, coincident with peak snowmelt
- Global forcing is dominated by fossil fuel and biofuel sources of BC, but strong biomass burning events can dominate Arctic forcing

Spatial/temporal characteristics of BC/snow forcing



- Forcing operates mostly in local springtime, when and where there is large snow cover exposed to intense insolation, coincident with peak snowmelt
- Global forcing is dominated by fossil fuel and biofuel sources of BC, but strong biomass burning events can dominate Arctic forcing
- Global-mean forcing (including snow and sea-ice):
 ~ 0.03 0.06 W m⁻² (Koch et al, Rypdal et al, Hansen et al, Jacobson)



• Efficacy (Hansen et al., 2005):

$$\left[\frac{\Delta T_s/F}{\Delta T_s(CO_2)/F(CO_2)}\right]$$

(2)

Efficacy

• Efficacy (Hansen et al., 2005):

$$\left[\frac{\Delta T_s/F}{\Delta T_s(CO_2)/F(CO_2)}\right]$$
(2)

 $\bullet\,$ Our equilibrium climate experiments indicate that BC/snow forcing has efficacy of 3 ± 1

Efficacy

• Efficacy (Hansen et al., 2005):

$$\left[\frac{\Delta T_s/F}{\Delta T_s(CO_2)/F(CO_2)}\right]$$
(2)

- $\bullet\,$ Our equilibrium climate experiments indicate that BC/snow forcing has efficacy of 3 ± 1
- Reason 1: All of the forcing energy is deposited directly in the cryosphere, a component of the Earth System responsible for powerful albedo feedback

The importance of snow grain size

- Snow exhibits large variability in grain size $(30 < r_e < 2000 \,\mu\text{m})$, $r_e \propto (\text{specific surface area})^{-1}$
- Grain size determines pure snow albedo, depth profile of absorption, and the magnitude of perturbation by impurities



Springtime forcing from BC and dust



- Springtime snow-averaged surface forcings (*Flanner et al.*, 2009)
 - Eurasia: $+3.9 \text{ W m}^{-2}$ (1.0 W m⁻² from dust)
 - North America: $+1.2 \text{ W m}^{-2}$ (0.2 W m⁻² from dust)

Springtime forcing from BC and dust



- Springtime snow-averaged surface forcings (*Flanner et al.*, 2009)
 - Eurasia: $+3.9 \text{ W m}^{-2}$ (1.0 W m⁻² from dust)
 - North America: $+1.2 \text{ W m}^{-2}$ (0.2 W m⁻² from dust)
- BC emissions from Asia increased from
 - $\sim 1.6-2.6\,\text{Tg/yr}$ during
 - 1980-2000 (Bond et al., 2007)

Sources of uncertainty

Perturbed physics experiments to characterize forcing uncertainty (*Flanner et al.*, 2007)

Table: Range in global-mean BC/snow radiative forcing resulting from reasonable ranges of the following factors: (*Flanner et al.*, 2007)

	Low	High
BC Emissions	-46%	+100%
Snow Aging	-42%	+58%
Melt Scavenging	-31%	+8%
Optical Properties	-12%	+12%
Snow Cover Fraction	-17%	+8%
Absorption by Dust (skewed)	$\sim \pm$	20%

Sources of uncertainty

Perturbed physics experiments to characterize forcing uncertainty (*Flanner et al.*, 2007)

Table: Range in global-mean BC/snow radiative forcing resulting from reasonable ranges of the following factors: (*Flanner et al.*, 2007)

	Low	High
BC Emissions	-46%	+100%
Snow Aging	-42%	+58%
Melt Scavenging	-31%	+8%
Optical Properties	-12%	+12%
Snow Cover Fraction	-17%	+8%
Absorption by Dust (skewed)	$\sim \pm$	20%

Something important is missing here

Sources of uncertainty

Perturbed physics experiments to characterize forcing uncertainty (*Flanner et al.*, 2007)

Table: Range in global-mean BC/snow radiative forcing resulting from reasonable ranges of the following factors: (*Flanner et al.*, 2007)

	Low	High
BC Emissions	-46%	+100%
Snow Aging	-42%	+58%
Melt Scavenging	-31%	+8%
Optical Properties	-12%	+12%
Snow Cover Fraction	-17%	+8%
Absorption by Dust (skewed)	$\sim \pm$	20%

Something important is missing here ... transport and deposition

Methods and Experiments

 Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)
- "Offline" land model forced with atmospheric boundary conditions

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)
- "Offline" land model forced with atmospheric boundary conditions
- Snowpack physics are fully active (meltwater removal of aerosols, snowpack growth/decay and layer division/combination of masses),

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)
- "Offline" land model forced with atmospheric boundary conditions
- Snowpack physics are fully active (meltwater removal of aerosols, snowpack growth/decay and layer division/combination of masses), but evolution is constrained by fixed atmospheric state

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)
- "Offline" land model forced with atmospheric boundary conditions
- Snowpack physics are fully active (meltwater removal of aerosols, snowpack growth/decay and layer division/combination of masses), but evolution is constrained by fixed atmospheric state
- Model averaging period: 1995–2004 with 5 years spinup, applying annually-repeating deposition

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)
- "Offline" land model forced with atmospheric boundary conditions
- Snowpack physics are fully active (meltwater removal of aerosols, snowpack growth/decay and layer division/combination of masses), but evolution is constrained by fixed atmospheric state
- Model averaging period: 1995–2004 with 5 years spinup, applying annually-repeating deposition
- Forcing from black carbon and mineral dust

- Apply monthly-resolved aerosol deposition fields from present-day AeroCom experiments to drive the NCAR CLM/SNICAR snow model and quantify land-snow radiative forcing
- Wet and dry deposition fields partitioned into hydrophilic and hydrophobic snowpack components (unique optical properties)
- "Offline" land model forced with atmospheric boundary conditions
- Snowpack physics are fully active (meltwater removal of aerosols, snowpack growth/decay and layer division/combination of masses), but evolution is constrained by fixed atmospheric state
- Model averaging period: 1995–2004 with 5 years spinup, applying annually-repeating deposition
- Forcing from black carbon and mineral dust
- Here, examine AeroCom:
 - Phase I "B" experiments (identical emissions)
 - Phase II "A2CTRL" experiments

BC Deposition in A2CTRL



Dust Deposition in A2CTRL



14/25

Phase I (B) Global forcing

Table: Global annual-mean radiative forcing of BC and mineral dust in land-based snowpack $[W\,m^{-2}]$

Model	BC	Mineral dust
ARQM	0.022	0.010
GISS	0.015	0.007
LOA	0.023	0.006
LSCE	0.023	0.007
MATCH	0.022	0.008
TM5	0.025	0.005
UIO-CTM	0.021	0.007
UIO-GCM	0.021	N/A
ULAQ	0.027	0.006
UMI	0.021	0.008
Mean	0.022	0.007
σ	14%	30%

Phase II (A2CTRL) Global forcing

Table: Global annual-mean radiative forcing of BC and mineral dust in land-based snowpack $[W\,m^{-2}]$

Model	BC	Mineral dust
CAM4-BAM	0.023	0.026
CAM4-Oslo	0.023	0.006
CAM-Oslo	0.023	0.006
HadGEM2-ES	0.027	0.002
MPIHAM-V2	0.022	0.011
Mean	0.024	0.010
σ	8%	90%

BC/Snow Forcing in Phase I B



17 / 25

BC/Snow Forcing in A2CTRL



Dust/Snow Forcing in A2CTRL



Seasonal Cycle of BC/Snow Forcing



Figure: Phase I B

Phase II A2CTRL

• Peak forcing in March or April

Seasonal Cycle of Dust/Snow Forcing



Spring BC/Snow Forcing in A2CTRL



Spring Dust/Snow Forcing in A2CTRL



 Collect more A2CTRL and Hindcast data! (BC, dust, and POM wet and dry deposition fields)

• Collect more A2CTRL and Hindcast data! (BC, dust, and POM wet and dry deposition fields) Contact: flanner@umich.edu

- Collect more A2CTRL and Hindcast data! (BC, dust, and POM wet and dry deposition fields) Contact: flanner@umich.edu
- Comparison with observations (e.g., Doherty et al., 2010)

- Collect more A2CTRL and Hindcast data! (BC, dust, and POM wet and dry deposition fields) Contact: flanner@umich.edu
- Comparison with observations (e.g., Doherty et al., 2010)
- Prescribe observed snow cover in CLM/CESM

- Collect more A2CTRL and Hindcast data! (BC, dust, and POM wet and dry deposition fields) Contact: flanner@umich.edu
- Comparison with observations (e.g., Doherty et al., 2010)
- Prescribe observed snow cover in CLM/CESM
- Include sea-ice forcing (which requires a different model)

- Collect more A2CTRL and Hindcast data! (BC, dust, and POM wet and dry deposition fields) Contact: flanner@umich.edu
- Comparison with observations (e.g., Doherty et al., 2010)
- Prescribe *observed* snow cover in CLM/CESM
- Include sea-ice forcing (which requires a different model)
- Quantify pre-industrial snow forcing from AeroCom experiments

Questions?

• Thanks to Michael, Philip, Stefan, and Mian.

