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Constraining the aerosol influence on liquid water path

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# Aerosols and liquid clouds

Aerosol modify the properties of liquid clouds

The albedo of a liquid cloud scene is a function of:

- Cloud fraction (CF)
- Liquid water path (LWP)
- Droplet number concentration (N<sub>d</sub>)



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- The AOD-LWP relationship is also confounded by humidity
- How do we separate the aerosol effect?



# How do we identify aerosol effects?



A number of different methods have been used:

- Controlling for confounders
- Exogenous perturbations (experiments)
- Mediating variables

### Mediating variables



Gryspeerdt et al., JGR, 2016

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Can we use this method for identifying the aerosol-LWP relationship?

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Note that we can't test this by setting N<sub>d</sub> constant

$$\left. \frac{dN_d}{dAI} = 0 
ightarrow \left. \frac{dLWP}{dAI} \right|_{causal} = 0$$

But it should also be able to "predict" the anthropogenic LWP change ( $\Delta$ LWP<sub>actual</sub>)

ECHAM6 HAM

$$\Delta LWP_{diag} = \left. \frac{dLWP}{dAI} \right|_{PD} \times \Delta AI$$

Actual **ΔLWP** 

Diag. ∆LWP (AI)







ECHAM6 HAM

CAM5.3

CAM5.3 MG2

CAM5.3 CLUBB-MG2

$$\Delta LWP_{diag} = \left. \frac{dLWP}{dAI} \right|_{PD} \times \Delta AI$$

Actual ΔLWP



50





# Calculating $\Delta N_d$



► Anthropogenic △N<sub>d</sub> can be accurately calculated from the present day CCN-N<sub>d</sub> relationship

$$\Delta N_d = \left. \frac{dN_d}{dCCN} \right|_{PD} \times \Delta CCN$$

 Use conditional probabilities to conserve the non-linearity (e.g. P(N<sub>d</sub>|CCN))

Gryspeerdt et al, PNAS, 2017





Actual ∆LWP













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  - Systematic errors in cloud retrievals (e.g. sub-adiabaticity)
- Feedbacks
  - Aerosol dependent entrainment of dry air
  - (Aerosol processing)

# Satellite relationships



The N<sub>d</sub>-LWP relationship in MODIS data is strongly non-linear

- Increase in LWP with increasing N<sub>d</sub> at low N<sub>d</sub>
- Decrease in LWP at high N<sub>d</sub>

#### Satellite relationships



## Correlated errors?

Both N<sub>d</sub> ( $\propto \tau_c^{0.5} r_e^{-2.5}$ ) and LWP ( $\propto \tau_c r_e$ ) are calculated from r<sub>e</sub> and  $\tau_c$ 

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Similar results using MODIS and AMSR-E (microwave) LWP

- Retrieval errors don't dominate the relationship
- Although they still play a role

What radiative forcing does this suggest?

- Changes in N<sub>d</sub> only (RFaci)
  - ▶ ≈-0.30 Wm<sup>-2</sup>



InAOD-Albedo Sensitivty

-3.0 -2.0 -1.0 0.0 1.0 2.0 3.0

- Changes in N<sub>d</sub> only (RFaci)
  - $\approx$ -0.30 Wm<sup>-2</sup>
- Changes in CF (mediated by N<sub>d</sub>)
  - ► ≈-0.50 Wm<sup>-2</sup> (Gryspeerdt et al., 2016)



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- Changes in LWP (mediated by N<sub>d</sub>)
  - ► ≈+0.27 Wm<sup>-2</sup> (upper bound)
- Total (liquid clouds only)
  - ▶ ≈-0.51 Wm<sup>-2</sup>



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Conditions where the  $N_d$  is varied exogenously provide an opportunity to test this

- Volcanoes
- Shiptracks

# Volcanic emissions



 In 2007, N<sub>d</sub>-LWP similar in regions A and B

- Region A shows weaker LWP decrease with N<sub>d</sub> in 2008
- Suggests little LWP change at high N<sub>d</sub>

# Shiptracks



- A similar effect is observed in shiptracks
- LWP increases in shiptracks where the "shiptrack" N<sub>d</sub> is low
- No change in LWP at high "shiptrack" N<sub>d</sub>

Thanks to Matt Christensen for shiptrack data!

# Scale dependence



The N<sub>d</sub>-LWP relationship is scale dependent

 Must reach temporal/spatial scale where N<sub>d</sub> is varied by aerosol

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- What about ice processes... ?

