



Global shortwave aerosol direct radiative forcing from MODIS measurements for mineral dust, marine aerosol, biomass-burning and industrial pollution.

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What is needed?

1. The aerosol optical thickness for each type
2. The aerosol optical properties for each type
3. The surface albedo

Step 1

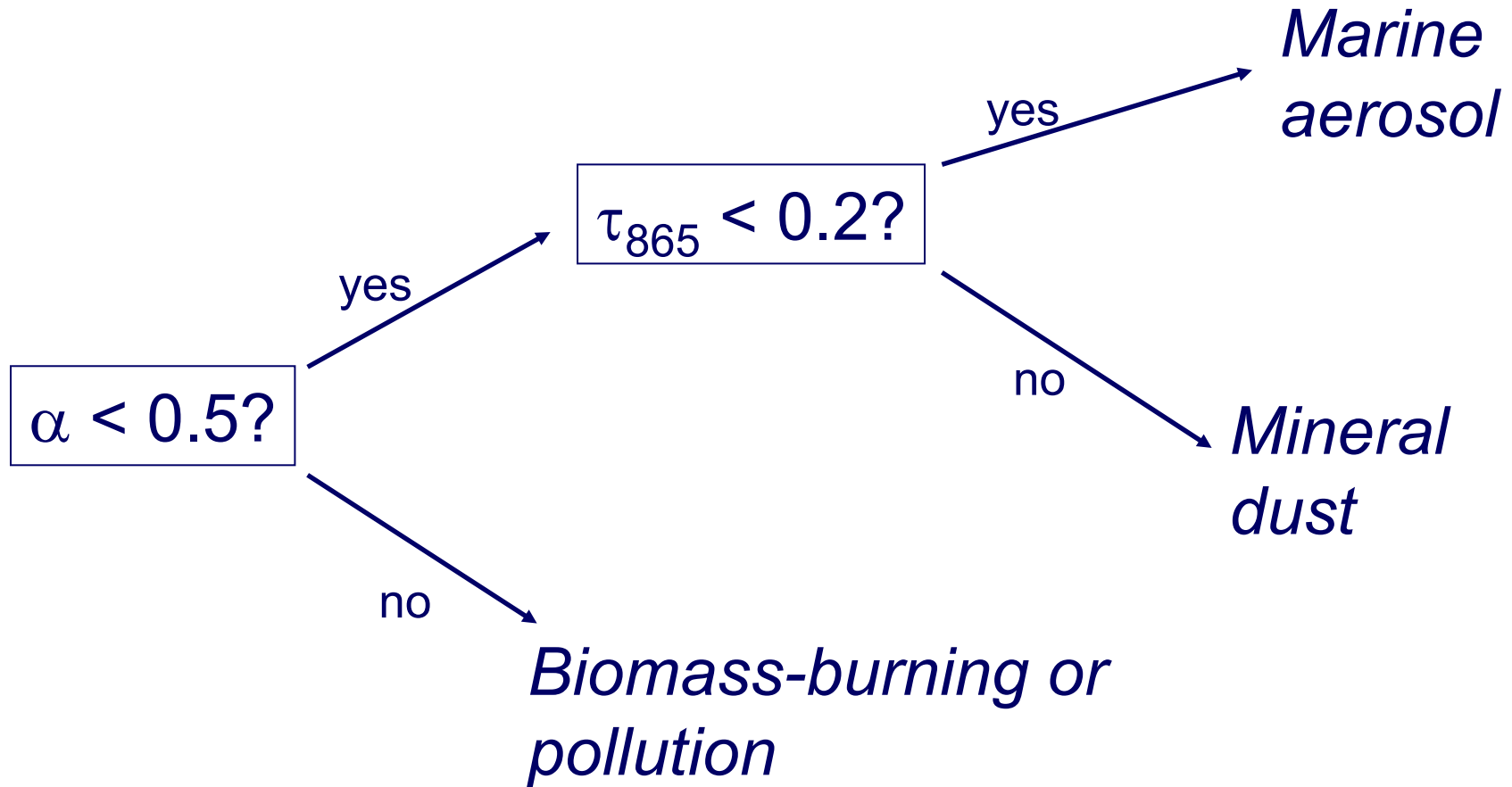
Get the optical thickness for mineral dust, marine aerosol and anthropogenic aerosols

Solve $\tau_{\text{total}} = \tau_{\text{dust}} + \tau_{\text{marine}} + \tau_{\text{biomass/pollution}}$

Help wanted!

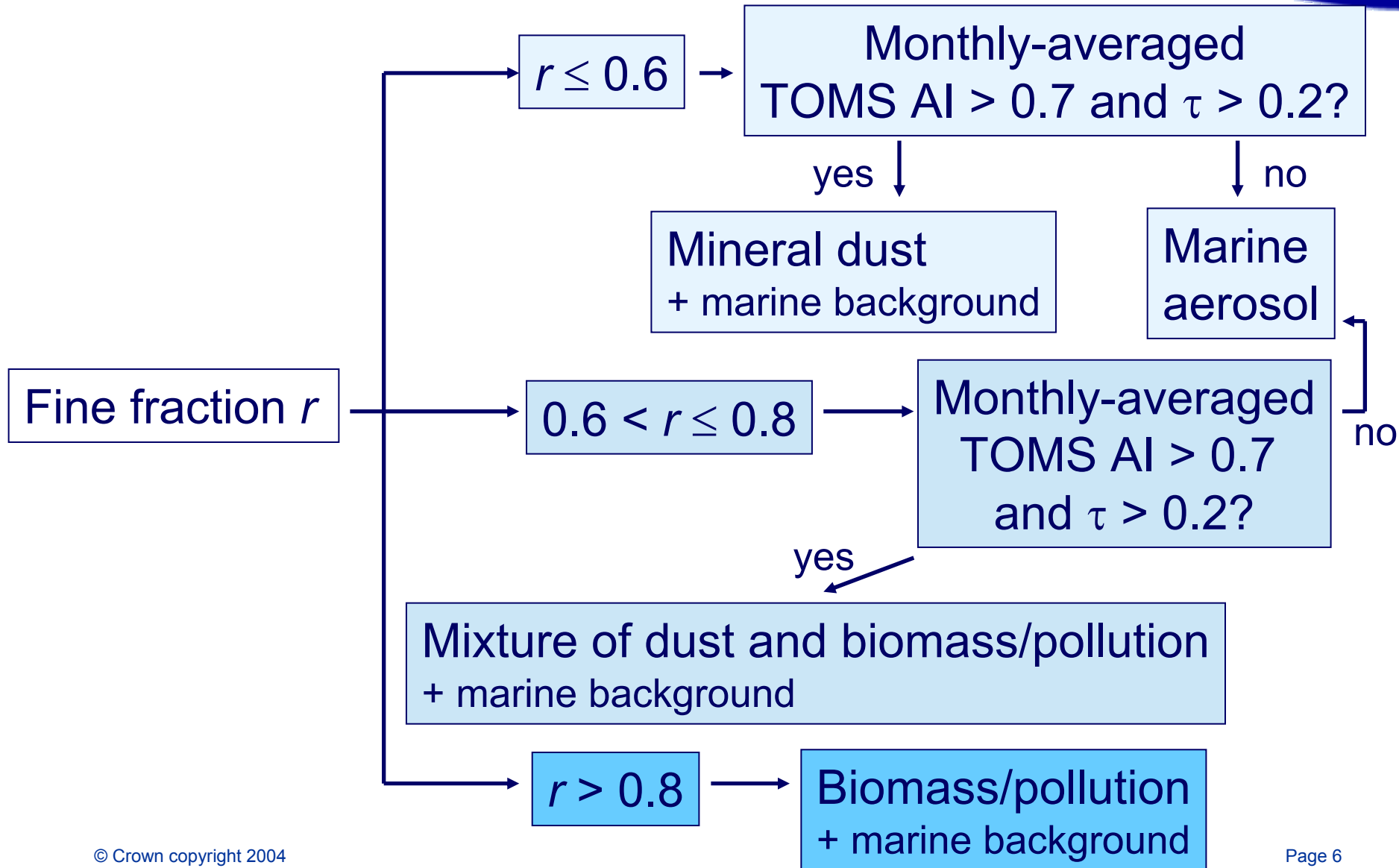
- Ångström exponent α
spectral dependence of the extinction
- Fine fraction r
fraction of the OT due to the accumulation-mode particles
- Surface wind speeds
give a rough estimate of the marine aerosol OT
- TOMS aerosol index
detects UV-absorbing aerosols (i.e. dust and biomass-burning)

The POLDER-1 algorithm over clear-sky oceans



Bellouin et al., *GRL*, 2003

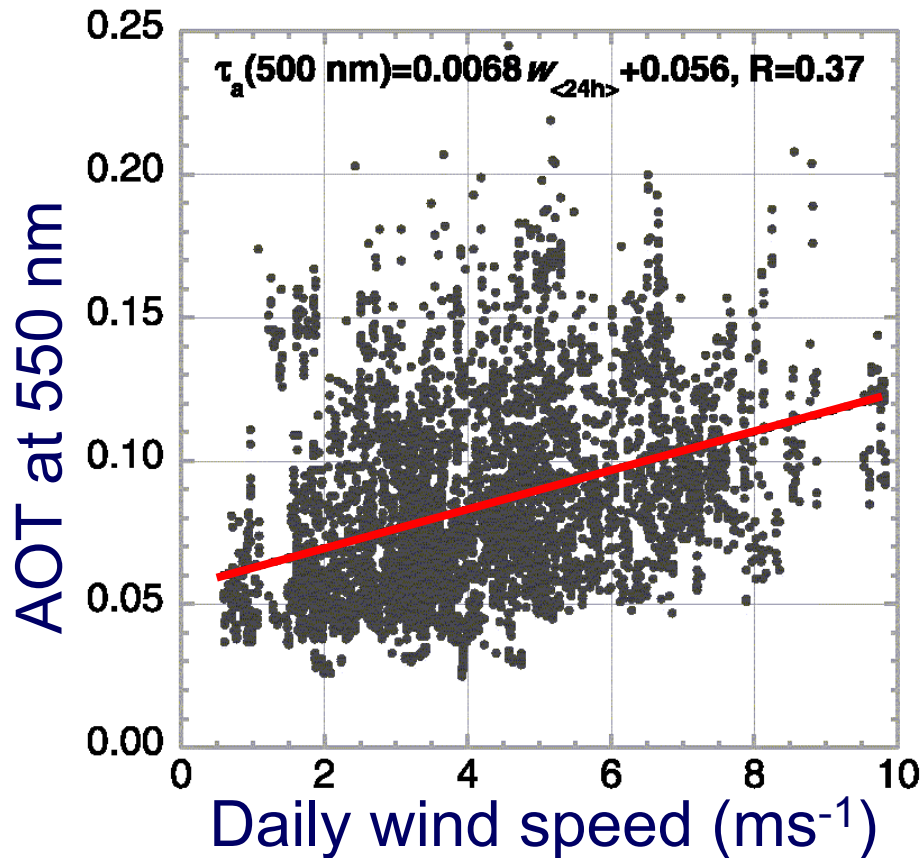
The MODIS algorithm over clear-sky oceans



Measurements from the Met Office C-130 Osborne and Haywood, *Atmos. Res.*, 2004

Experiment	Aerosol type	r
SHADE	Mineral dust	0.67
SAFARI 2000	Aged biomass-burning (over ocean)	0.97
	Fresh biomass-burning (over land)	0.95
TARFOX	Industrial pollution	0.88
ACE-2	Industrial pollution, mixed with marine aerosol	0.60
---	Marine aerosol	0.16

Get a *sensible* estimate of the marine aerosol OT when dust or biomass/pollution is identified.



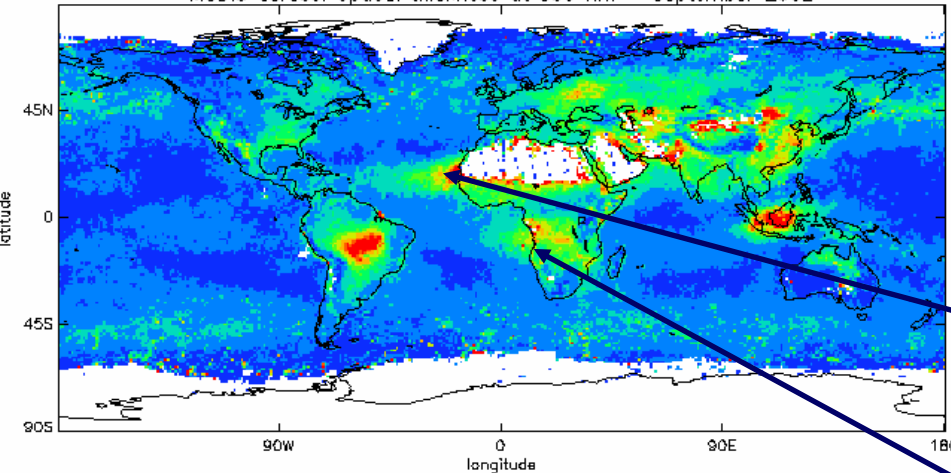
Linear relationship from Smirnov et al., *JGR*, 2003

In the algorithm, wind speeds are provided by SSM/I.

The MODIS algorithm: Data for September 2002



MODIS aerosol optical thickness at 550 nm – September 2002



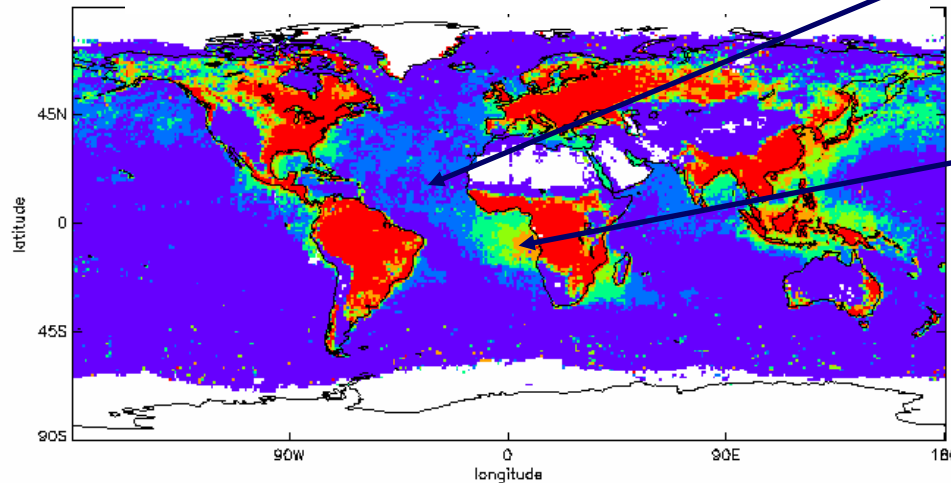
MODIS total AOT at 550 nm

Dust event

Biomass-burning event



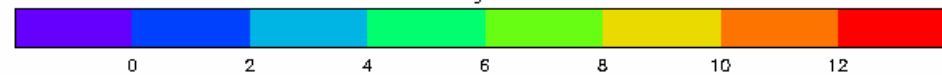
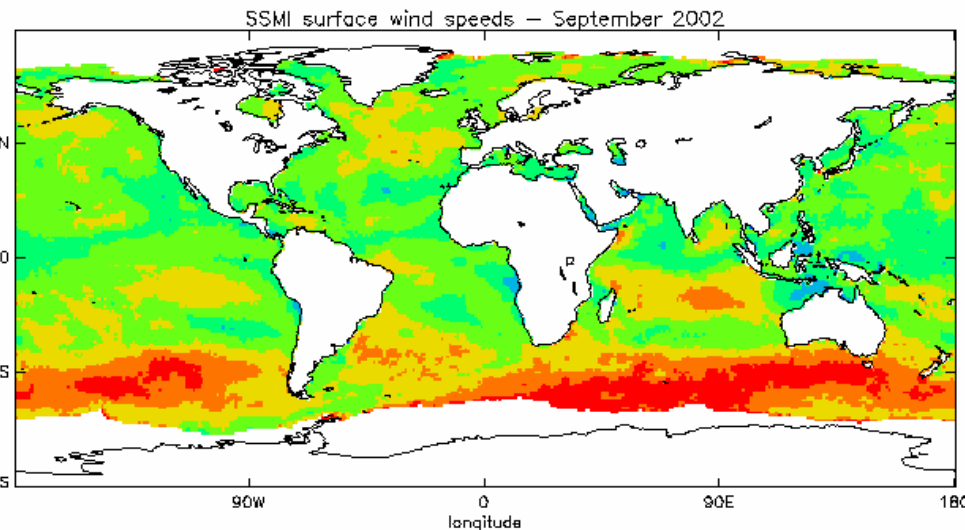
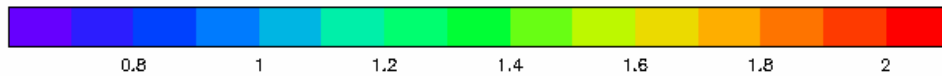
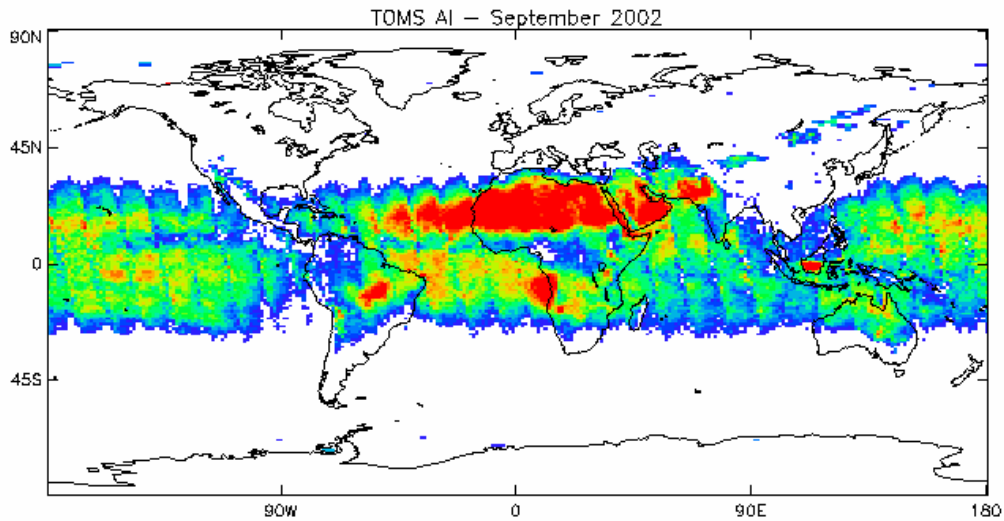
0 0.2 0.4 0.6 0.8



MODIS fine fraction

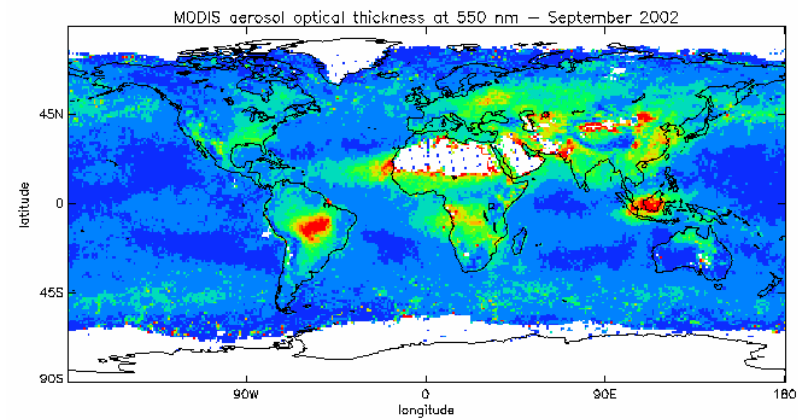
0.5 0.6 0.7 0.8 0.9

The MODIS algorithm: Data for September 2002



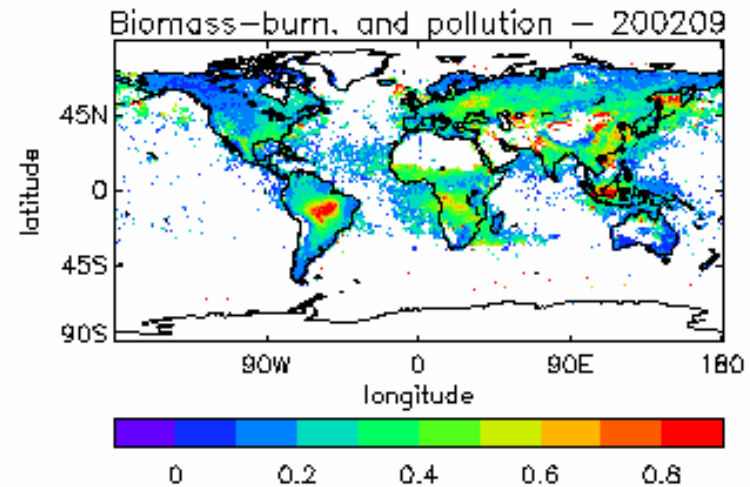
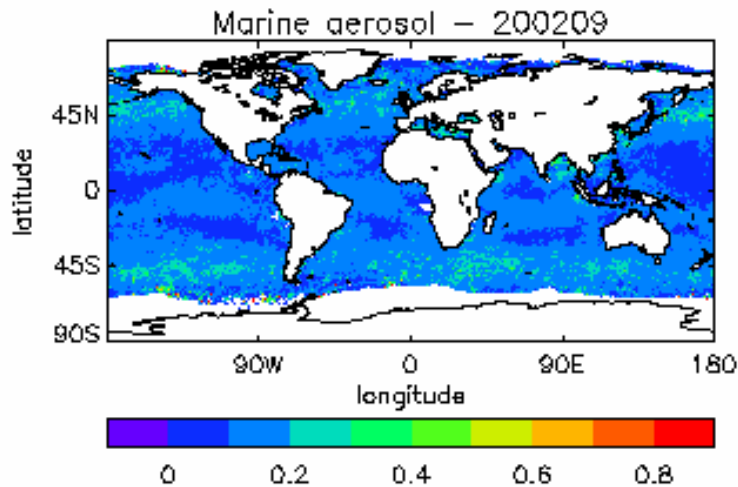
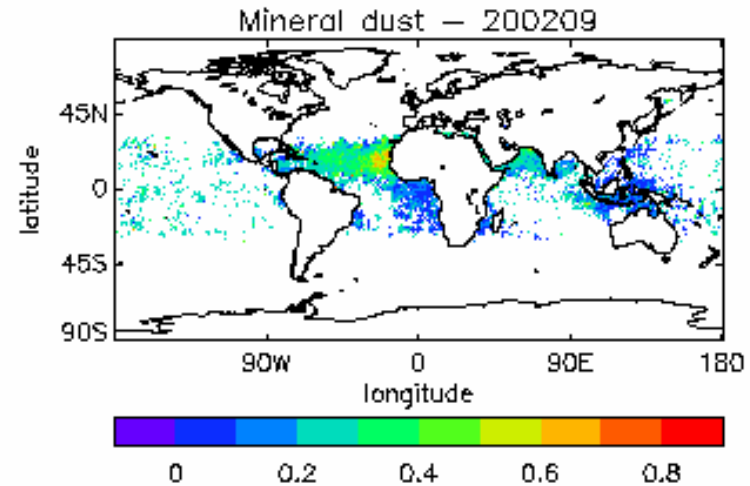
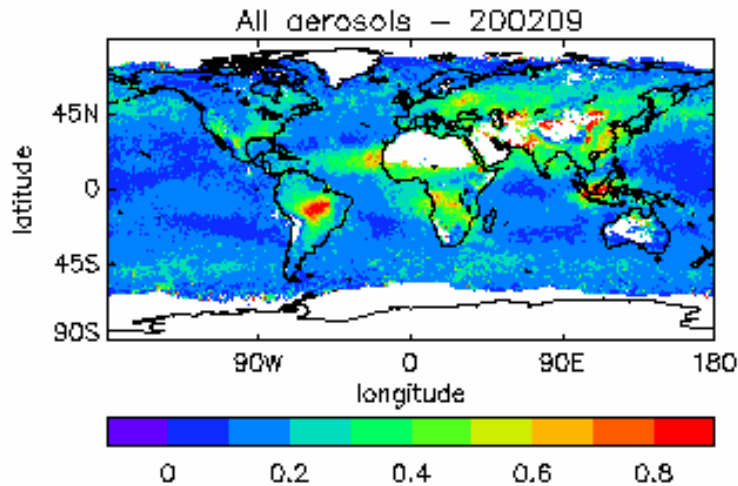
← TOMS Aerosol Index

MODIS τ_{550} →

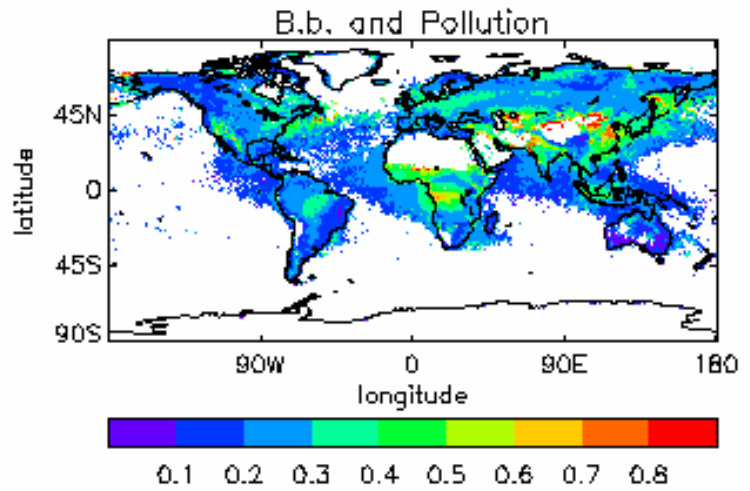
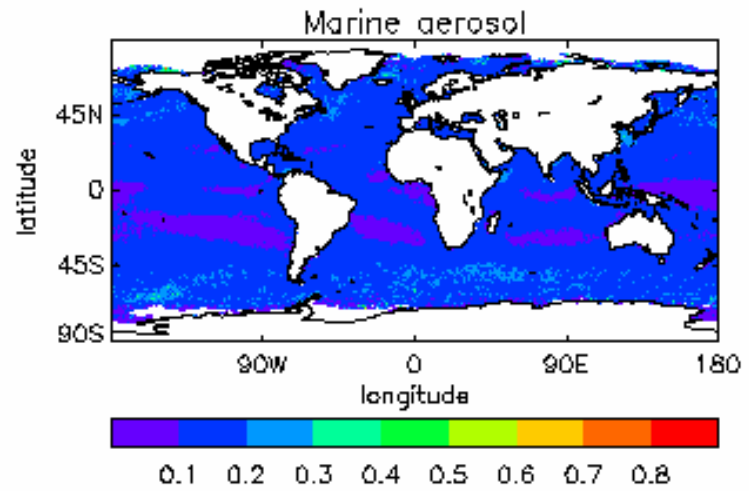
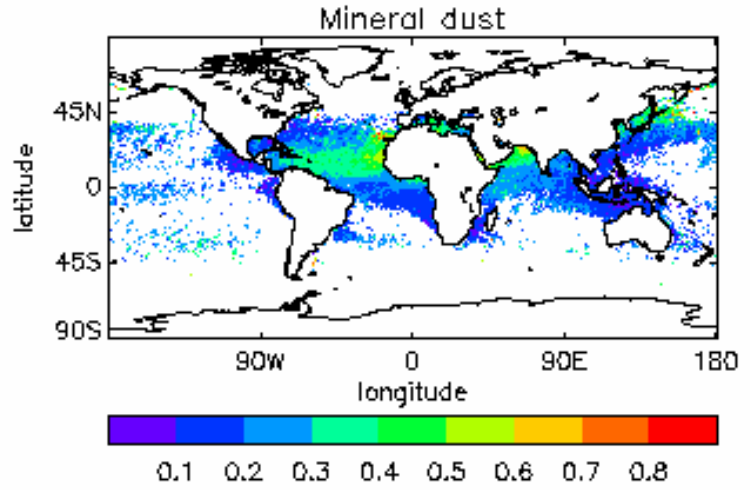
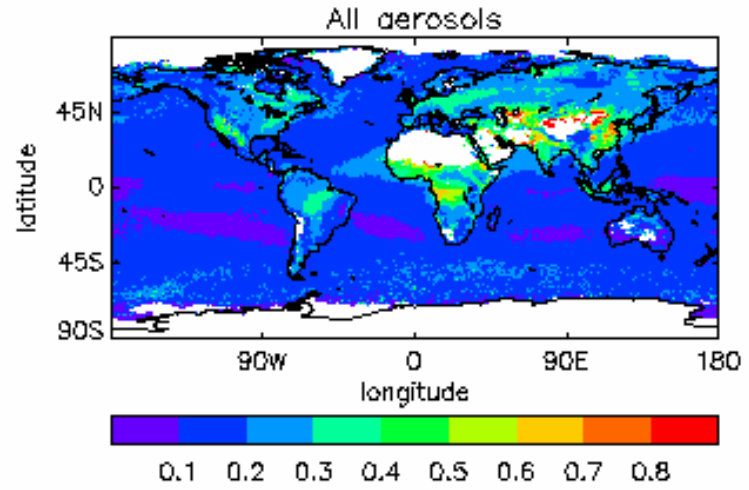


← SSM/I wind speeds

The MODIS algorithm: Results for September 2002



Distributions of optical thicknesses for 2002



Step 2

**Estimate the radiative forcing
from the discriminated optical
thicknesses**

From the optical thickness to the radiative forcing



Aerosol model:
- size distribution
- refractive index
(from AERONET)

Mie theory, 24λ

Aerosol optical thickness
of a given aerosol type
(from MODIS or POLDER)

$\tau(\lambda)$, $\omega_0(\lambda)$, $P(\Theta, \lambda)$

**radiative transfer
code**

Surface albedo
(VIS & near IR)

DARF ΔF

- instantaneous or daily-averaged
- at the TOA or surface

Aerosol optical properties

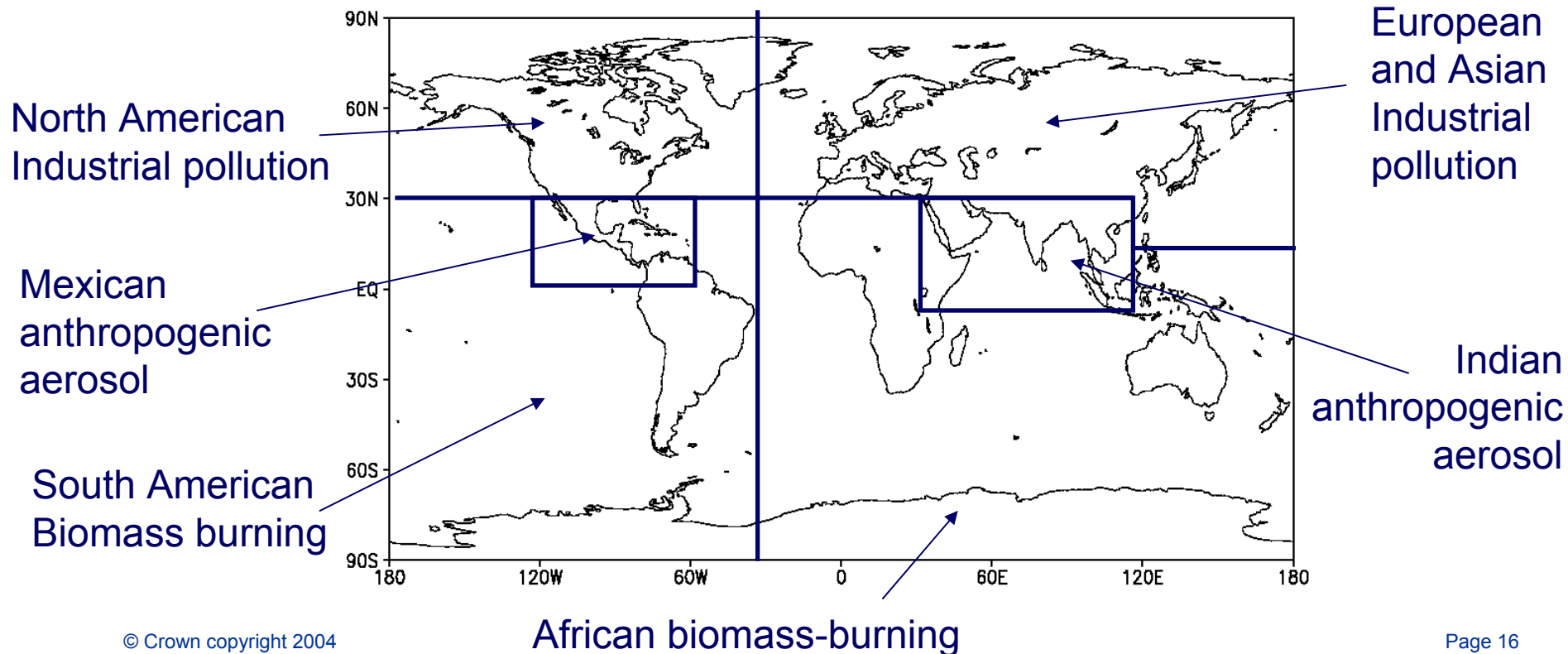
Dubovik et al., JAS, 2002

Aerosol	AERONET site	ω_0 at 550 nm
Dust	Cape Verde	0.97
Marine aerosol	Hawaii	0.98 (0.99)
Industrial pollution	Greenbelt, USA	0.97
Industrial pollution	Créteil, France	0.93
Industrial pollution, biomass-burning	Mexico City, Mexico	0.88
Industrial pollution, biomass-burning	Maldives (INDOEX)	0.89
Biomass-burning	Brazil	0.90
Biomass-burning	Zambia	0.86

Biomass-burning and pollution properties

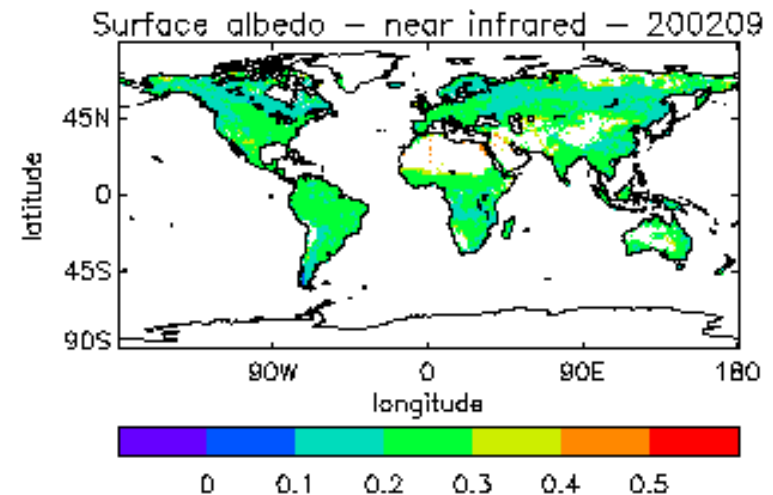
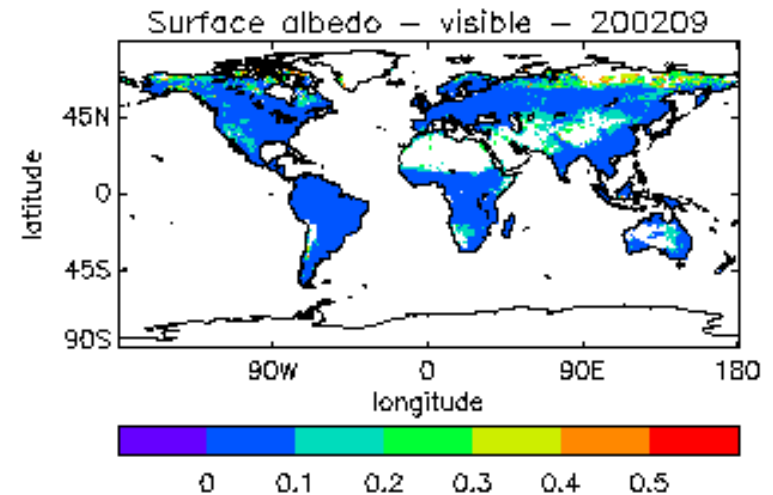


- The optical thickness is derived in the same way for all biomass-burning and pollution aerosols.
- But optical properties differ according to geographic location, using regional boxes.

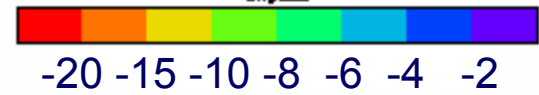
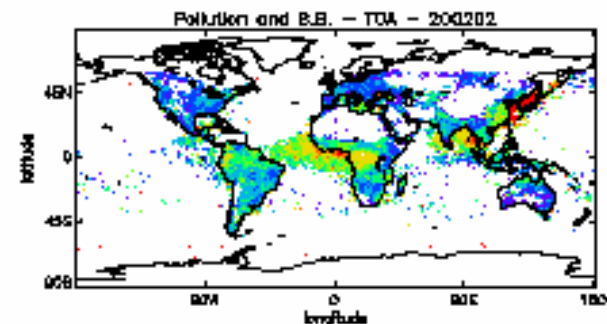
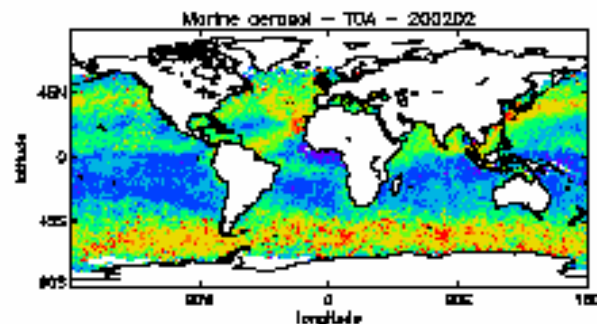
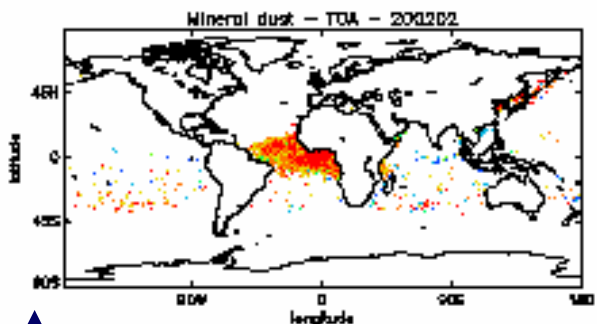


Over ocean, the albedo is computed using *Cox and Munk* [1954]

Over land, the albedo is derived from MODIS measurements (products MOD43B3, *Schaaf et al.*, 2002) and corrected for aerosol effects.



MODIS: Monthly average for February 2002

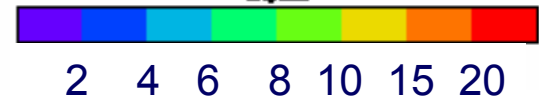
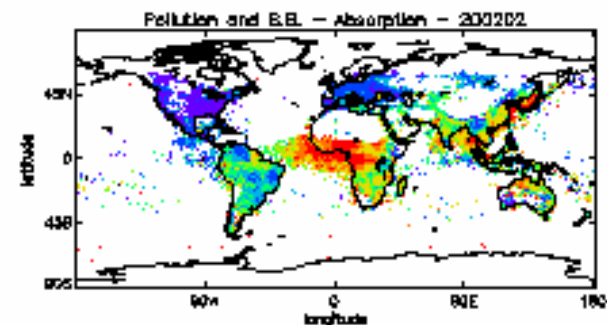
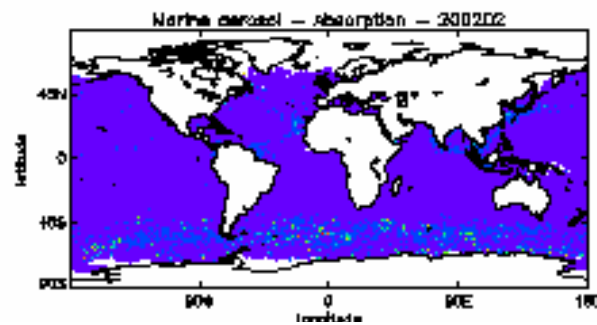
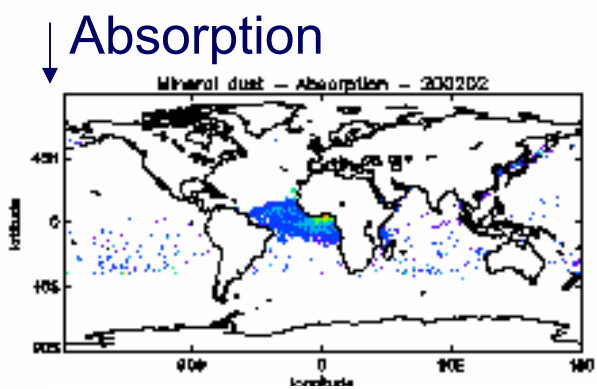


↑
Top of atmosphere

Mineral dust

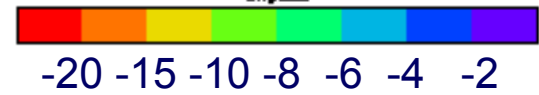
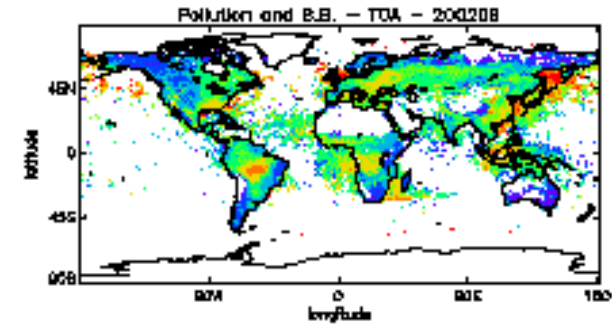
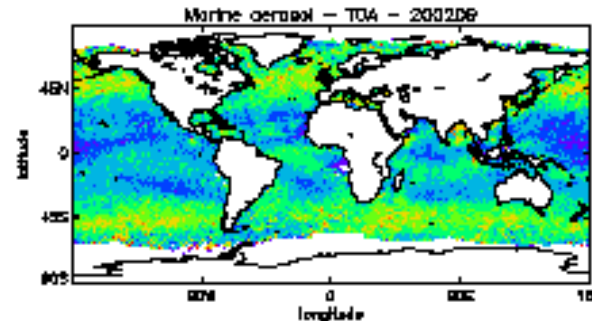
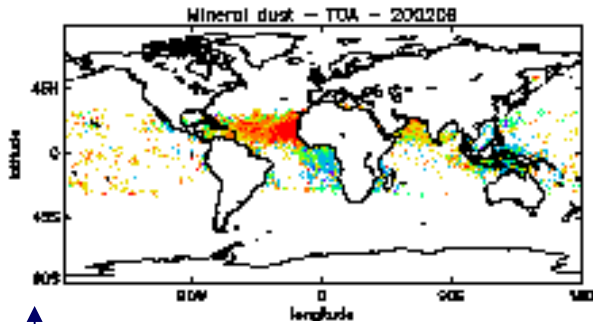
Marine aerosol

Biomass+Poll.



↓
Absorption

MODIS: Monthly averages for September 2002



Top of atmosphere

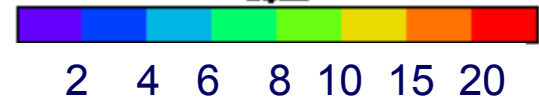
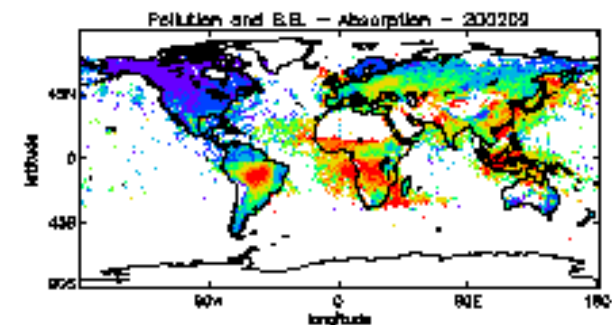
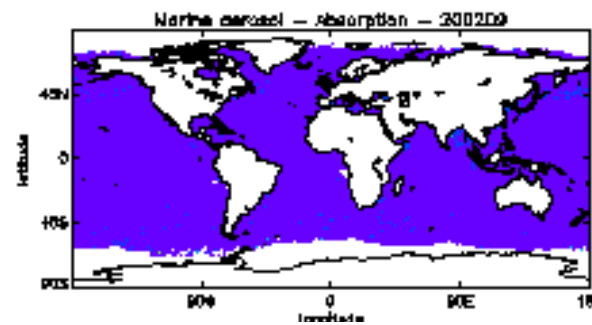
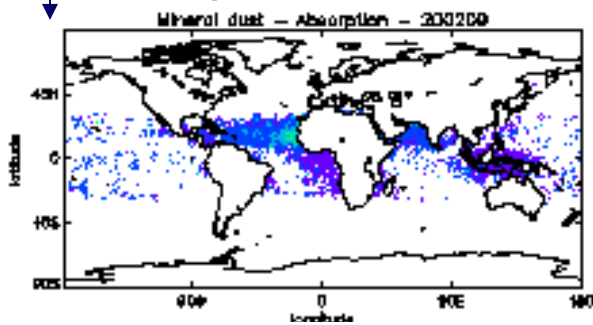
Mineral dust

Marine aerosol

Biomass+Poll.



Absorption



MODIS: Global clear-sky averages for 2002



Mineral dust	TOA (Wm ⁻²)	Surface (Wm ⁻²)	Abs. (Wm ⁻²)	τ (550 nm)	E (Wm ⁻² / unit τ)
Global	-0.48	-0.57	0.09	0.009	-56
Ocean	-0.71	-0.85	0.14	0.013	-56
Marine aerosol					
Global	-3.61	-4.25	0.64	0.076	-47
Ocean	-5.35	-6.31	0.95	0.113	-47
bb + poll					
Global	-2.39	-5.43	3.04	0.093	-26
Ocean (+)	-0.52	-1.18	0.66	0.014	-37
Land (-)	-6.18	-14.06	7.88	0.255	-24

(+) lower bound of anthropogenic RF (-) upper bound of anthropogenic RF

- Our algorithm applied to MODIS data does a good job distributing the total optical thickness to mineral dust, marine aerosol, and biomass-burning and pollution aerosols.
- Choosing realistic aerosol properties from AERONET measurements improves the confidence in the estimated radiative forcings.
- Paper to be submitted soon.

Radiative forcing efficiency

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Contributions from

Nicolas Bellouin

Shekar Reddy

Jim Haywood

The concept of radiative forcing efficiency (RFE) has been introduced to decouple uncertainties on aerosol burden/OD from uncertainties in other inputs and RT and to allow intercomparaison.

Clear-sky RFE of a particular aerosol type depends on:

- aerosol single scattering albedo & aerosol upscattering
- surface albedo
- diurnal and seasonal distribution of SZA at a particular location / region
- histogram of AOD (for a given average AOD).
- + small uncertainty on RT scheme (assuming RT is done properly!)

All-sky RFE depends additionnally on:

- vertical distribution of aerosol and cloud
- cloud fraction.
- + it may be more sensitive to the RT scheme used.

Moreover RFE will depend critically on

- RH growth factor if reported by unit of dry mass (sulfate, OM, sea-salt)
- radius cut size if reported by unit of mass for sea-salt and dust

Clear-sky and all-sky TOA SW RFE from our GCM calculations:

	clear-sky	all-sky
sulfate	RFE = -235	and -145 W (g sulfate) ⁻¹
	per mass of sulfate, but also includes ammonium & water fairly constant since B&A [1995], on the low side? fairly constant for different SRES sulfate distributions	
BC	RFE = +1200	and +1400 W (g BC) ⁻¹
	BC single scattering albedo = 0.2 BC density is low (1 g cm ⁻³)	
OM	RFE = -132	and -87 W (g OM) ⁻¹
	slightly absorbing, less hygroscopic than sulfate	

Needs to be intercompared in AEROCOM B & PRE
Weighted by the sophistication of the RT procedure.

Global RF and RFE from MODIS/AERONET aerosol properties and RT calculations

Mineral dust	TOA (Wm^{-2})	Surface (Wm^{-2})	Abs. (Wm^{-2})	τ (550 nm)	E (Wm^{-2} / unit τ)
Global	-0.48	-0.57	0.09	0.009	-56
Ocean	-0.71	-0.85	0.14	0.013	-56
Marine aerosol					
Global	-3.61	-4.25	0.64	0.076	-47
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Clear-sky TOA SW RFE (@550 nm) from the GCM calculations:

				“MODIS”
dust ocean	RFE = -21	W m ⁻² tau ⁻¹	vs	-56 W m ⁻² tau ⁻¹
sea-salt ocean	RFE = -25	W m ⁻² tau ⁻¹	vs	-47 W m ⁻² tau ⁻¹
anthropogenic ocean	RFE = -12	W m ⁻² tau ⁻¹	vs	-37 W m ⁻² tau ⁻¹
anthropogenic land	RFE = -10	W m ⁻² tau ⁻¹	vs	-24 W m ⁻² tau ⁻¹
anthropogenic globe	RFE = -11	W m ⁻² tau ⁻¹	vs	-26 W m ⁻² tau ⁻¹

BUT GCM clear-sky \diamond MODIS clear-sky (sampling issue) !

==> sample MODIS clear-sky in model nudged 2002 run

==> intercompare in AEROCOM B&PRE to see if LMDZ is an outlier

* RT scheme ?

* aerosol SSA ?

* surface albedo ?

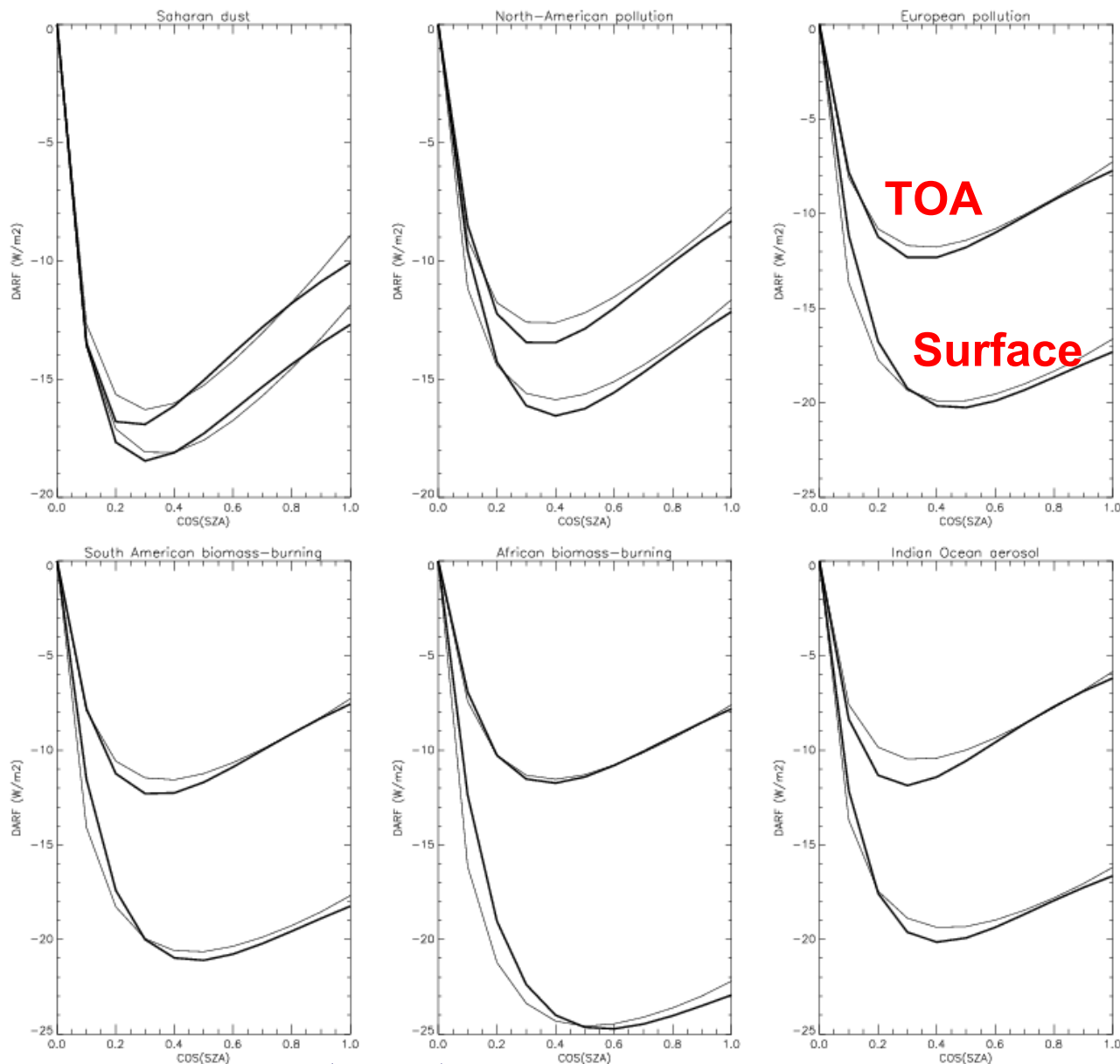
* our GCM dust and sea-salt calculations are done in the presence of other (absorbing) aerosols, which shifts RFs to less negative values.

Shortwave 24-stream 24-waveband versus 2-stream 2-waveband RT codes

Aerosol optical depth = 0.1 Surface albedo = 0.0

7 aerosol models

DARF (Wm^{-2})

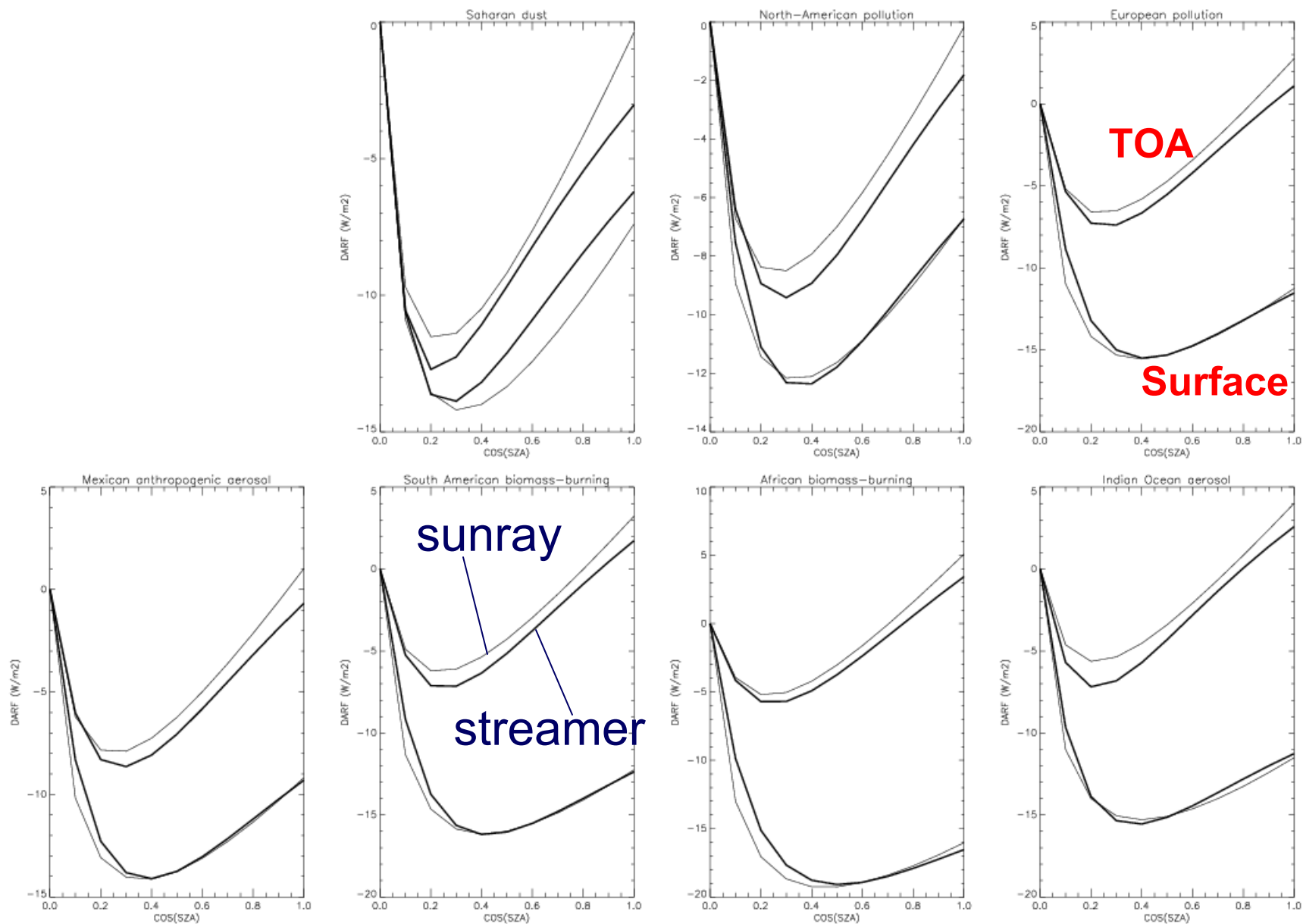


cos (SZA)

DARF (Wm^{-2})

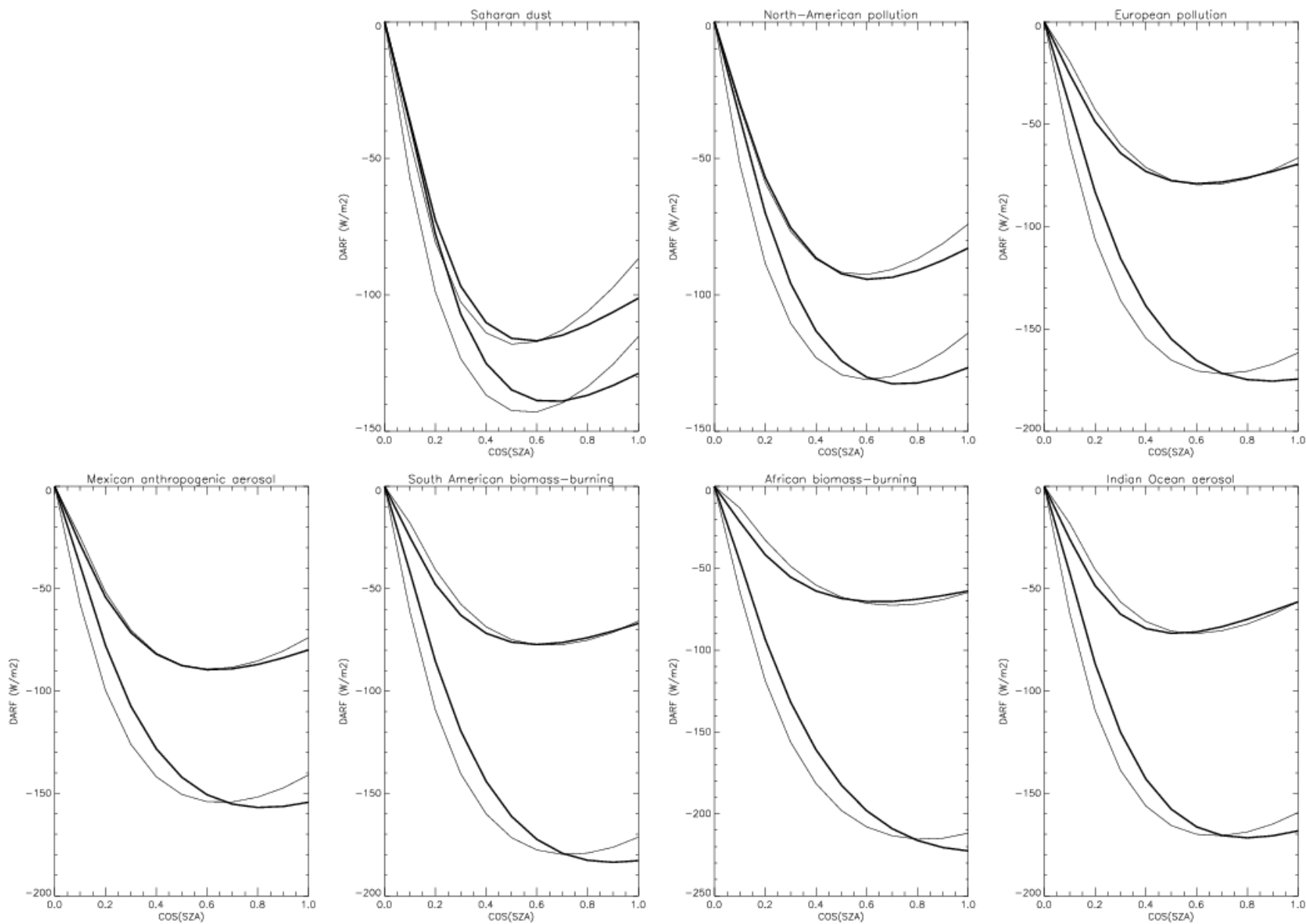
Broadband 24-stream 24-waveband versus 2-stream 2-waveband RT codes

Aerosol optical depth = 0.1 Surface albedo = 0.2



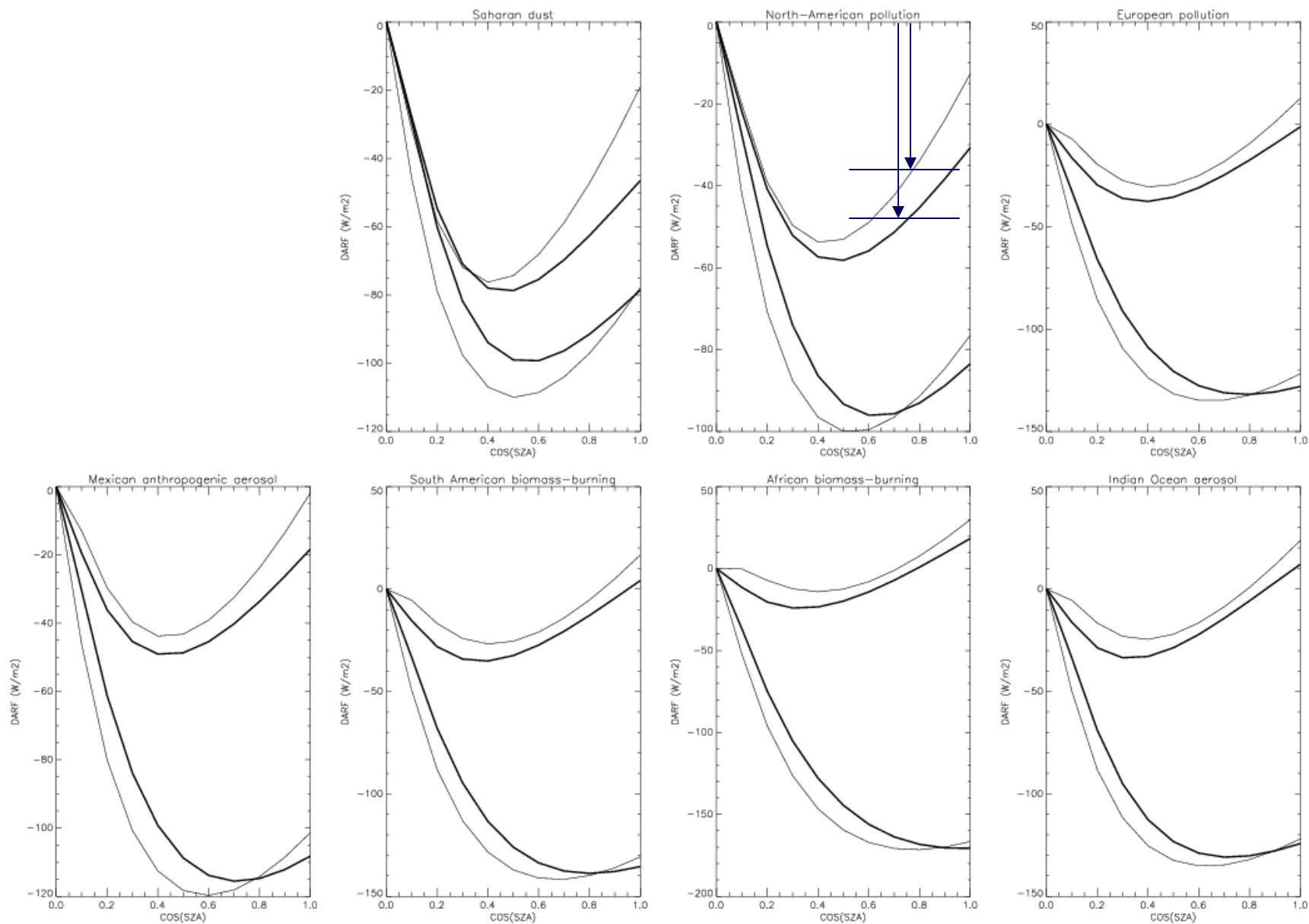
Broadband 24-stream 24-waveband versus 2-stream 2-waveband RT codes

Aerosol optical depth = 1.0 Surface albedo = 0.0



Broadband 24-stream 24-waveband versus 2-stream 2-waveband RT codes

Aerosol optical depth = 1.0 Surface albedo = 0.2



The direct aerosol RF is not as linear as we may think!

$$\Delta F_{\text{dust}} + \Delta F_{\text{bb}} \not\approx \Delta F_{\text{dust+bb}}$$

Implication is that $F_{\text{dust+bb}} - F_{\text{bb}} > F_{\text{dust}} - F_0$

